

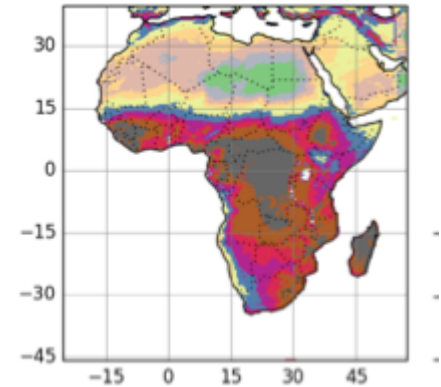
Regional scale soil-moisture feedbacks in a convection-permitting simulation over Africa

Sonja Folwell, Chris Taylor, Rachel Stratton (UK Met. Office)

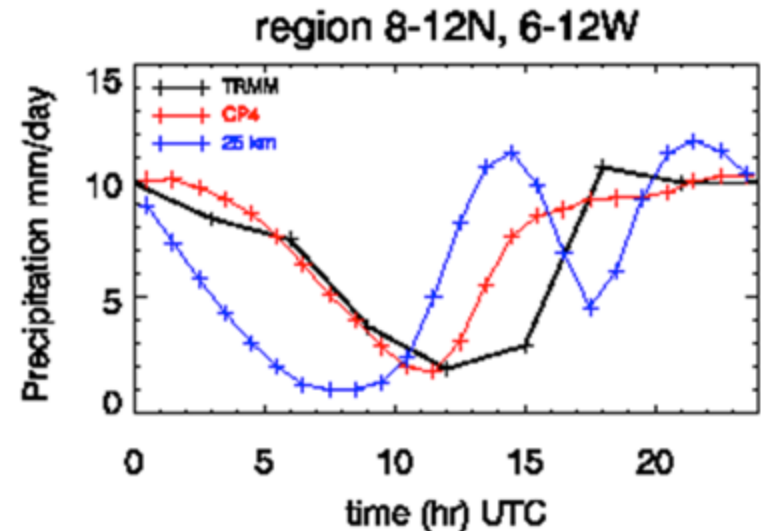
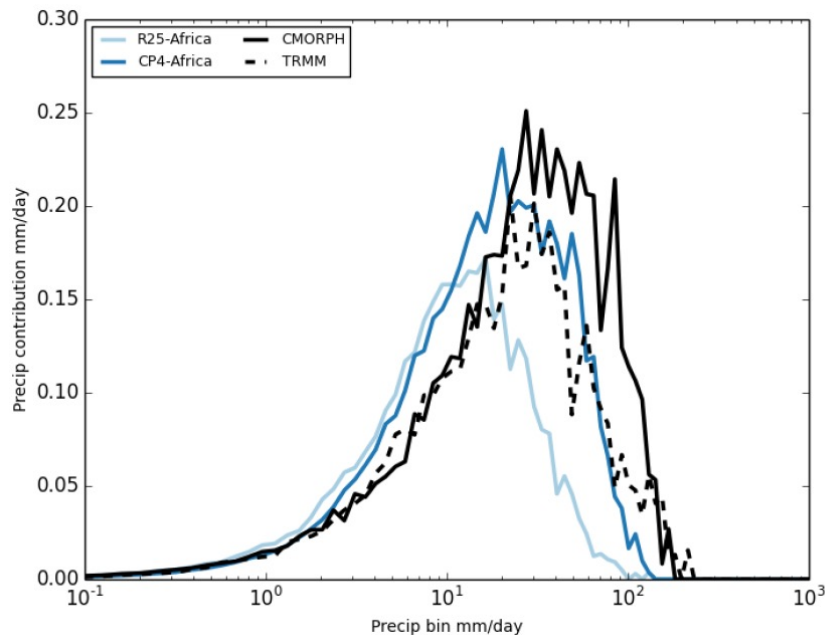
Experimental setup for CP4-Africa

Two model experiments using the **MetUM**, both forced with observed SST (1997 to 2006), and forced by a common global run (25 km):

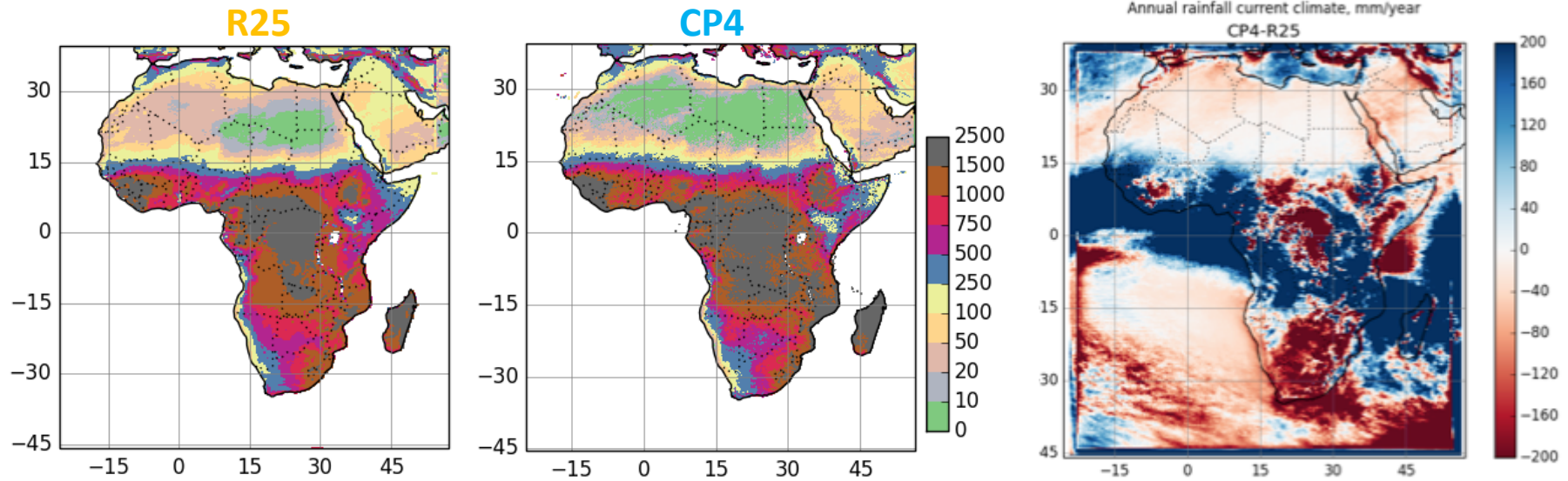
- 25 km parameterised convection (**R25**)
- 4.5 km convection permitting (**CP4**), convection switched off



Stratton et al. J Clim, 2018



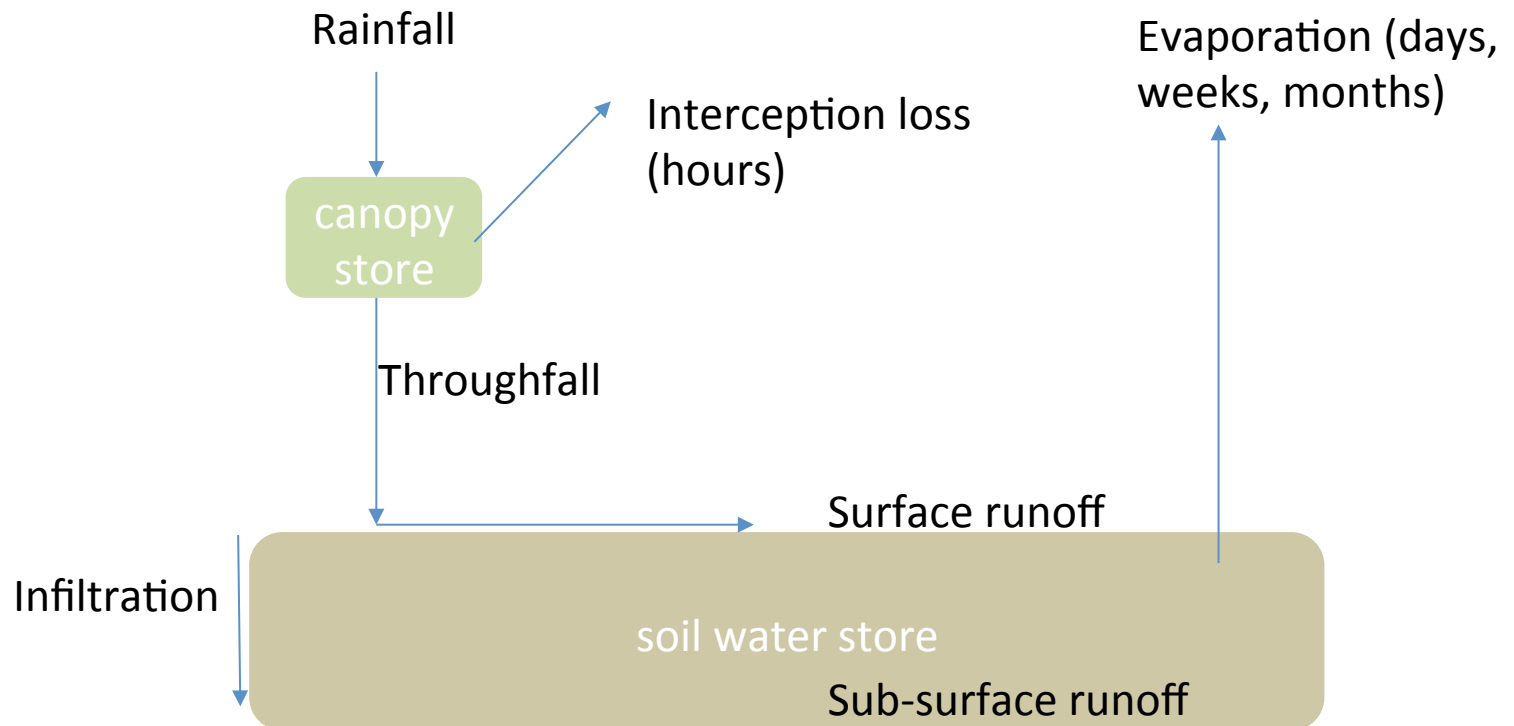
Annual rainfall



Annual rainfall over land: 611 mm/year vs. 672 mm/year

Surface water partitioning

- Are surface fluxes sensitive to the whether the model is run in a parameterised or convection permitting mode?
- Changing the rainfall intensity, quantity so what is the net effect of all these changes?

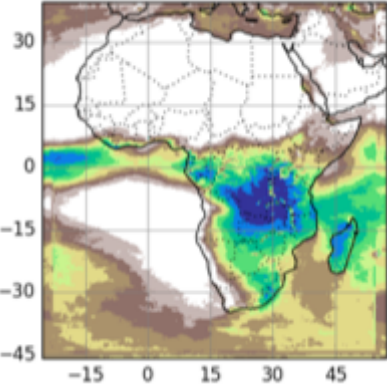


CP4: Fewer events

Fraction of days with >0.5 mm rain, DJF

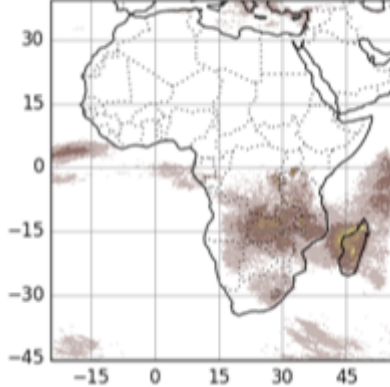
R25

25km



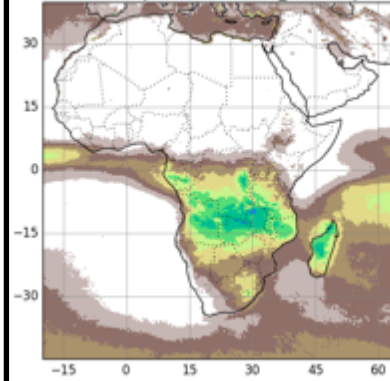
CP4

4km



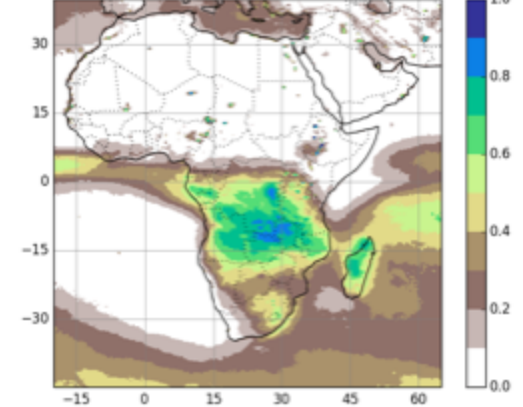
TRMM

Frac. days >0.5 mm, TRMM 3B42, DJF



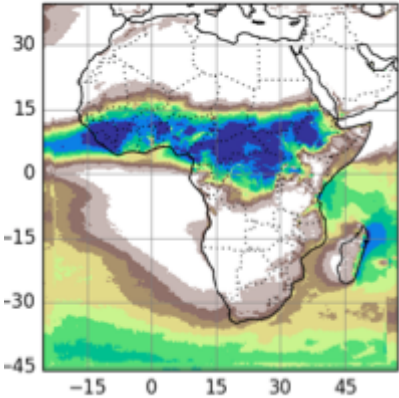
CMORPH

Frac. days >0.5 mm, CMORPH, DJF

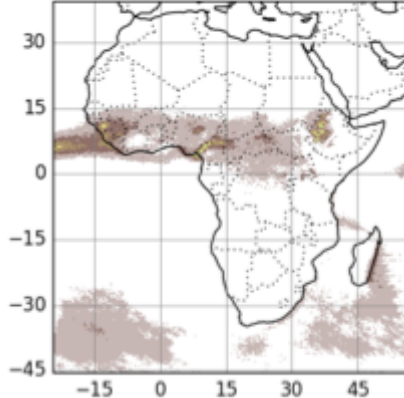


Fraction of days with >0.5 mm rain, JJA

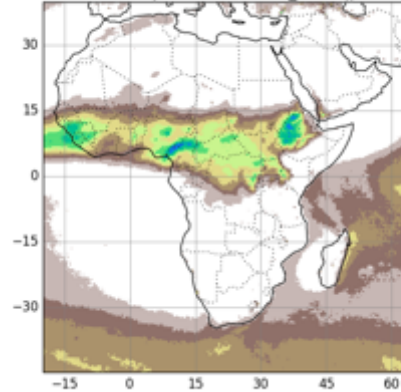
25km



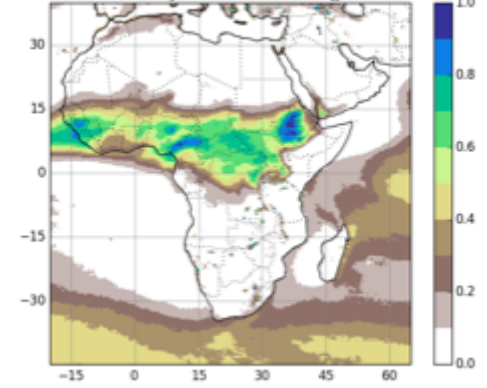
4km



Frac. days >0.5 mm, TRMM 3B42, JJA

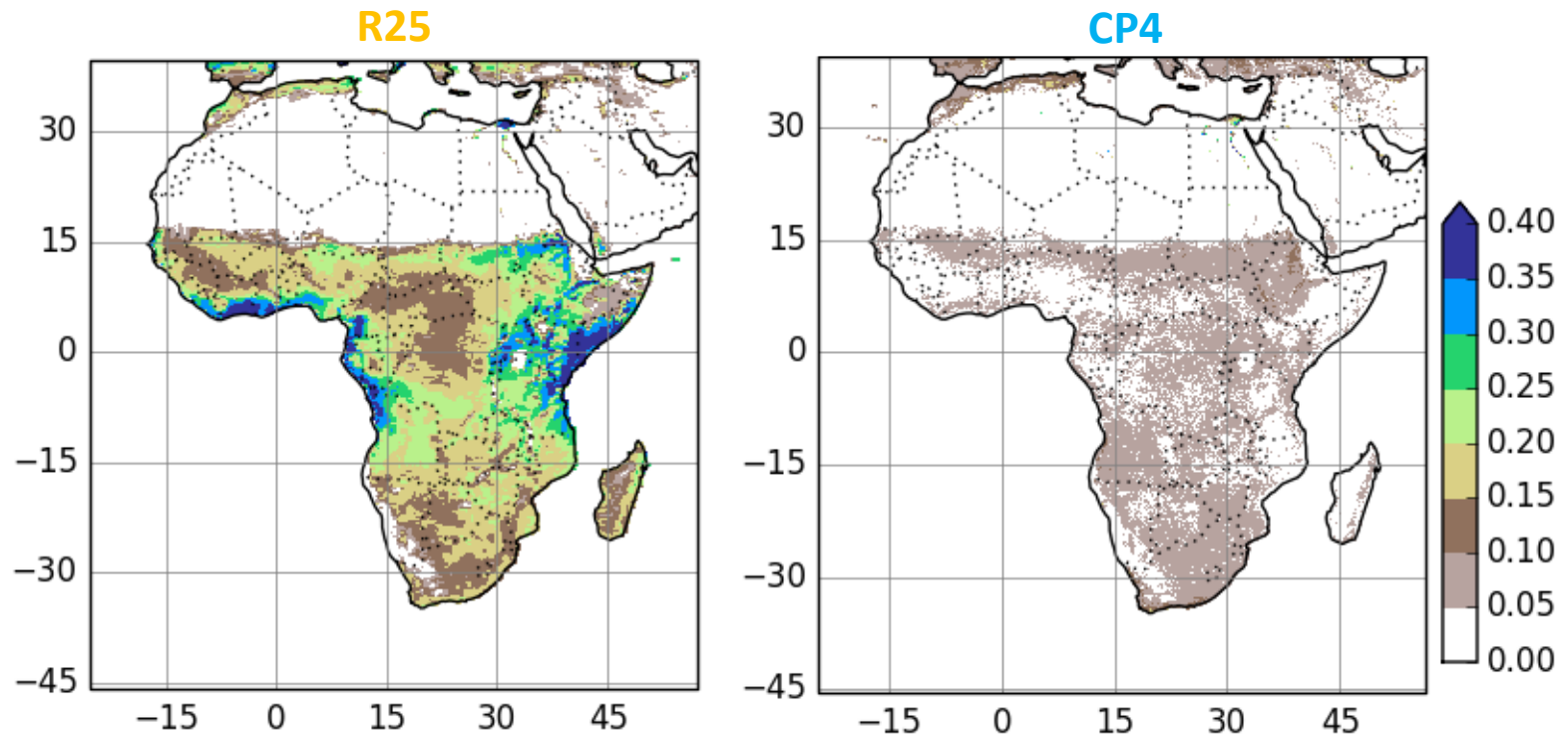


Frac. days >0.5 mm, CMORPH, JJA



Fractional canopy interception loss

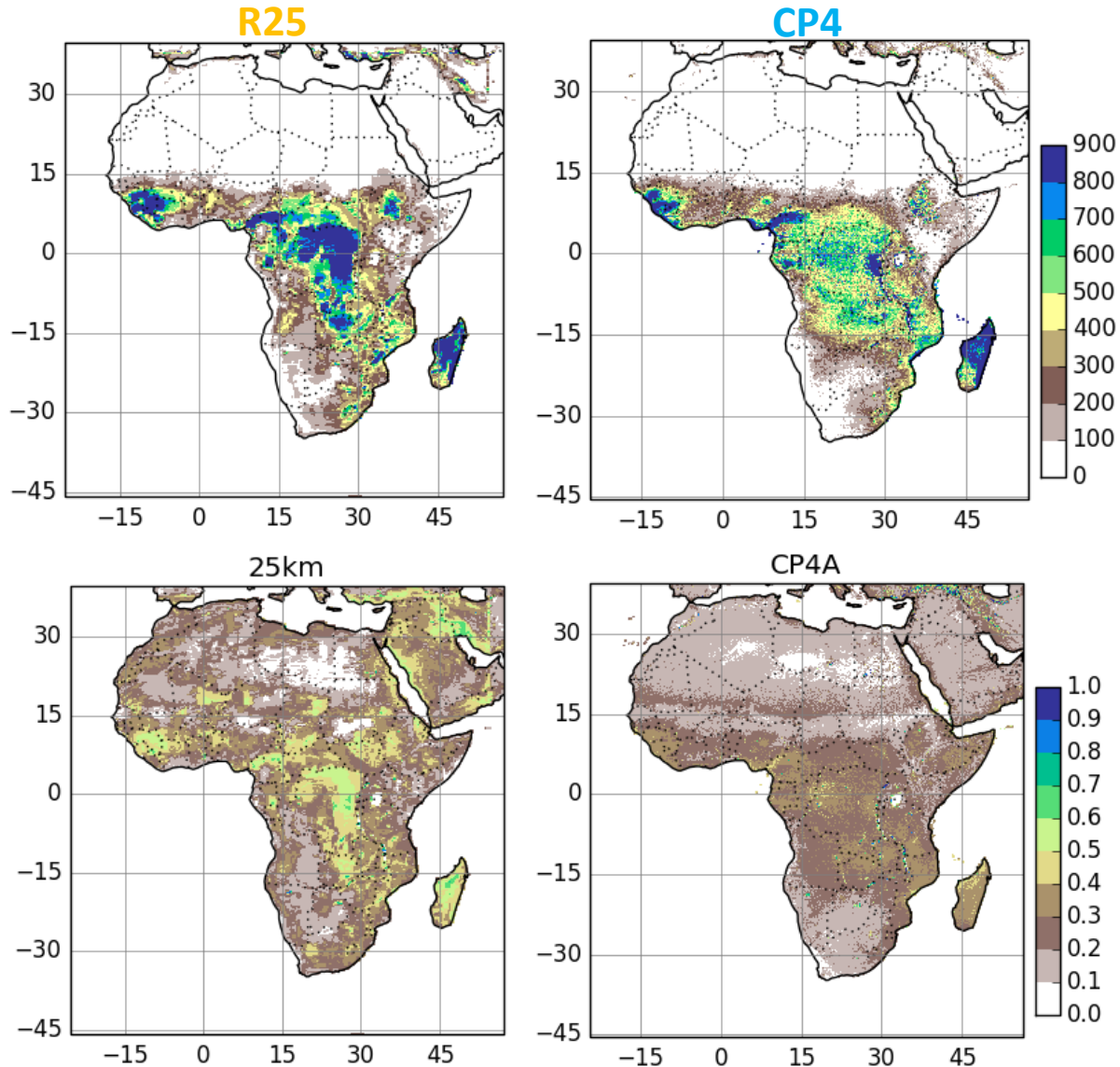
Water intercepted and returned to the atmosphere over very short timescales, E_c
Fraction of annual rainfall intercepted by the canopy (E_c/Rain)



Observations of interception loss over tropical rain forests (Amazonia) $\sim 15\%$
Tai Park in Ivory Coast = 9.2% (Aug to Dec)

Surface runoff

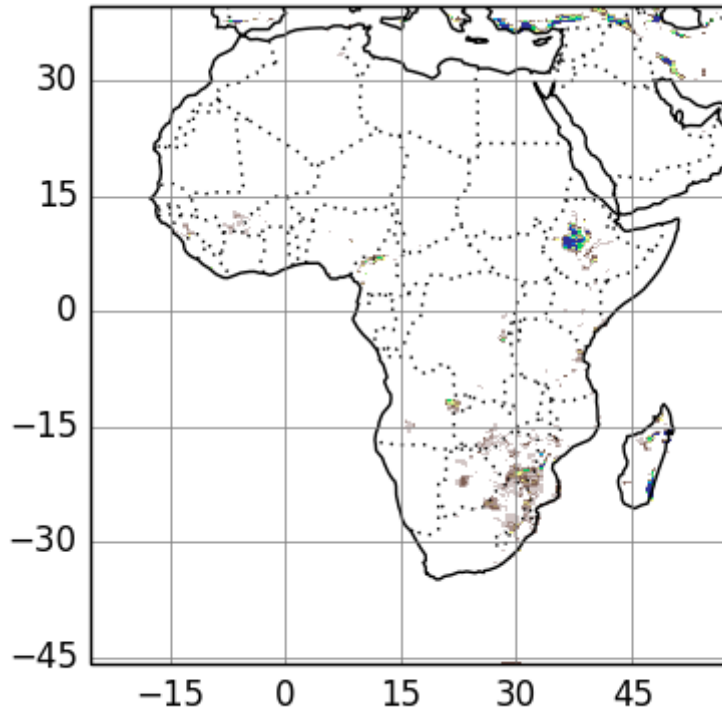
Annual surface runoff, 214 mm/year, 189 mm/yr



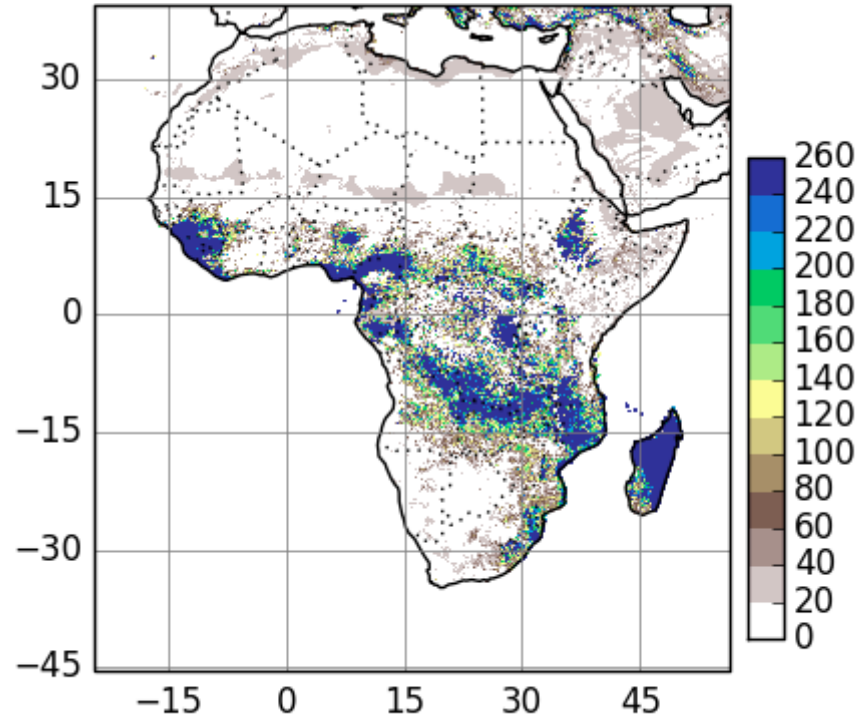
Sub-surface runoff

Annual sub-surface runoff, 6 mm/year, 79 mm/year

R25



CP4



JULES sub-grid rainfall

JULES assumes a sub-grid distribution of rainfall

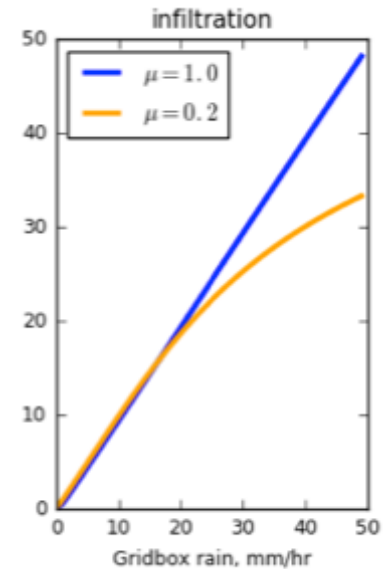
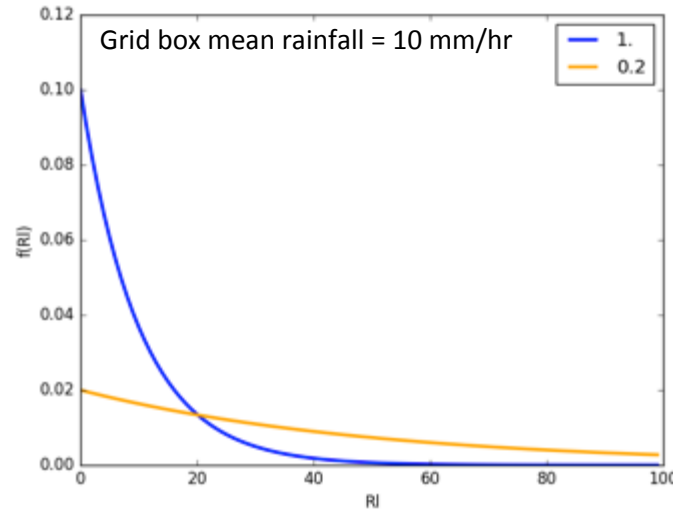
$$f(R \downarrow l) = (\mu/R) \exp(-\mu R \downarrow l / R)$$

Gridbox mean throughfall, $T \downarrow$

$$T \downarrow f = R(1 - C/C \downarrow m) \exp(-\mu C \downarrow m / R)$$

Gridbox mean infiltration excess runoff, Y

$$Y = \begin{cases} RC/C \downarrow m \exp(-\mu KC \downarrow m / RC) + R(1 - C/C \downarrow m) \exp(-\mu C \downarrow m / R \Delta t), & K \Delta t \leq C \\ CR \cdot \exp(-\mu(K \Delta t + C \downarrow m - C) / R \Delta t), & \Delta t > C \end{cases}$$



In **coupled** mode:

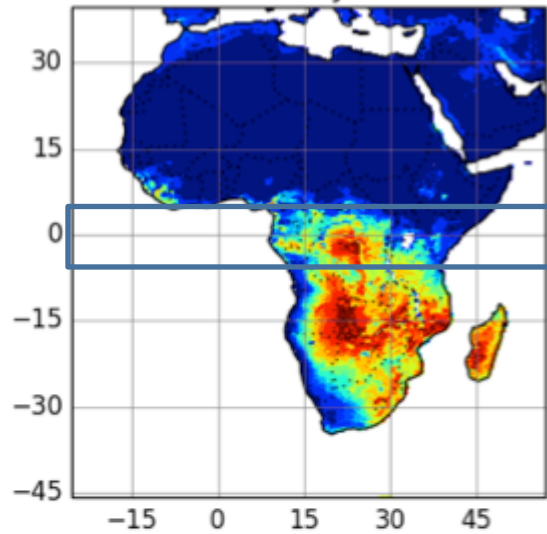
R25 model -> convective rainfall; rainfall area fraction ~ 0.2.

-> large scale rainfall; 0. < rainfall area fraction < 1.

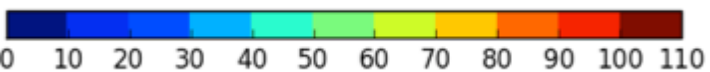
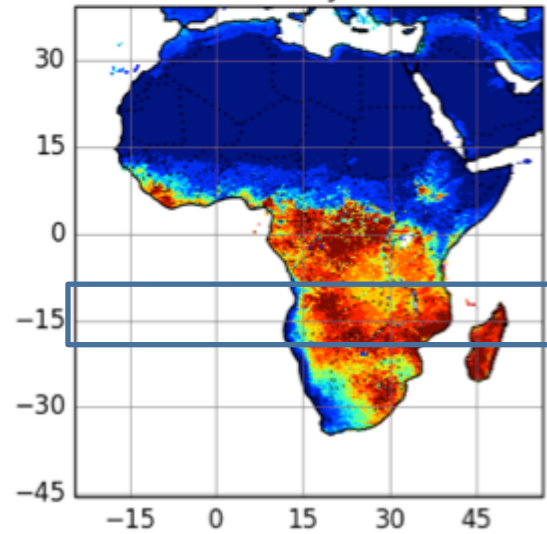
CP4 model -> large scale rain; rain area fraction = 1

Seasonal evaporation

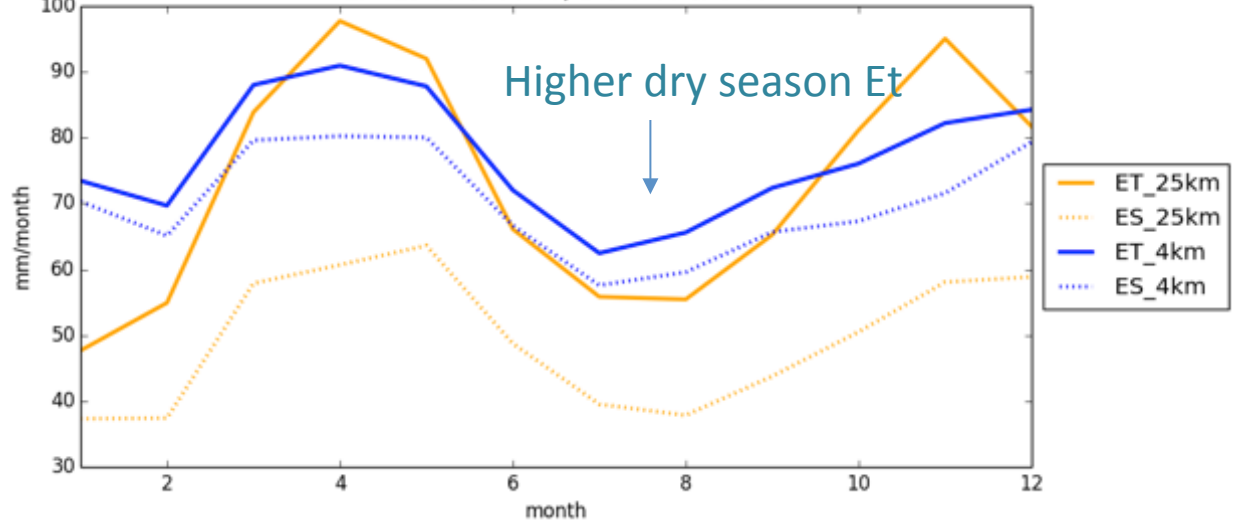
25km-Jan



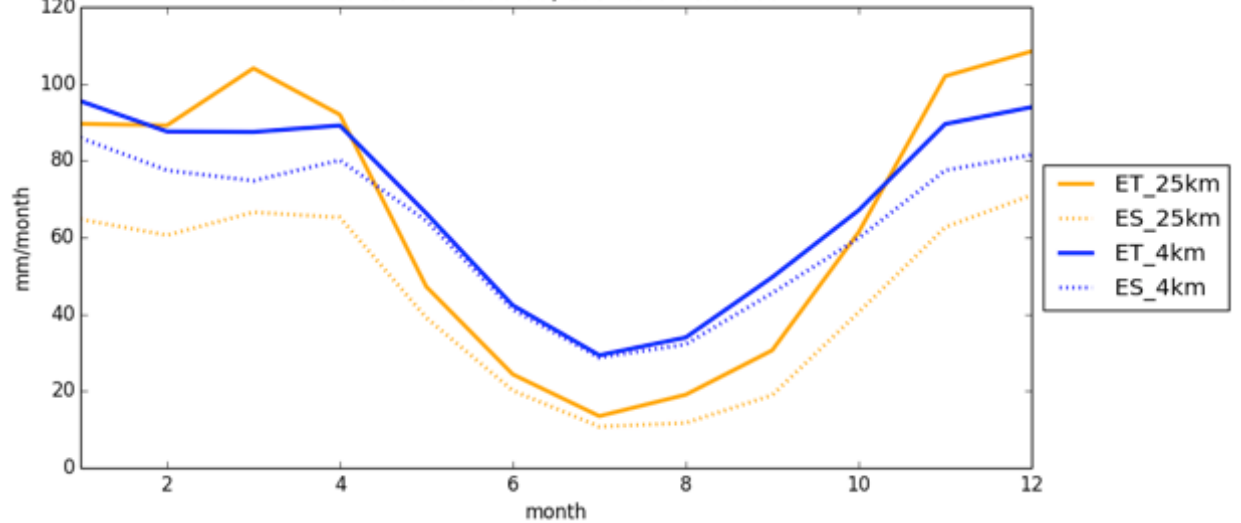
CP4A-Jan



Lon-15.00 to 50.00, Lat -5.00 to 5.00

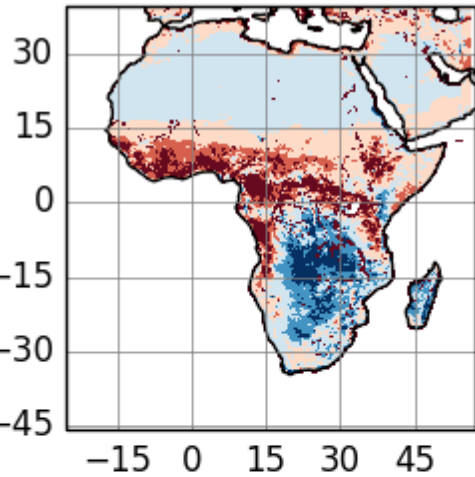


Lon-15.00 to 50.00, Lat -15.00 to -5.00



Fluxes: DJF

CP4 - R25



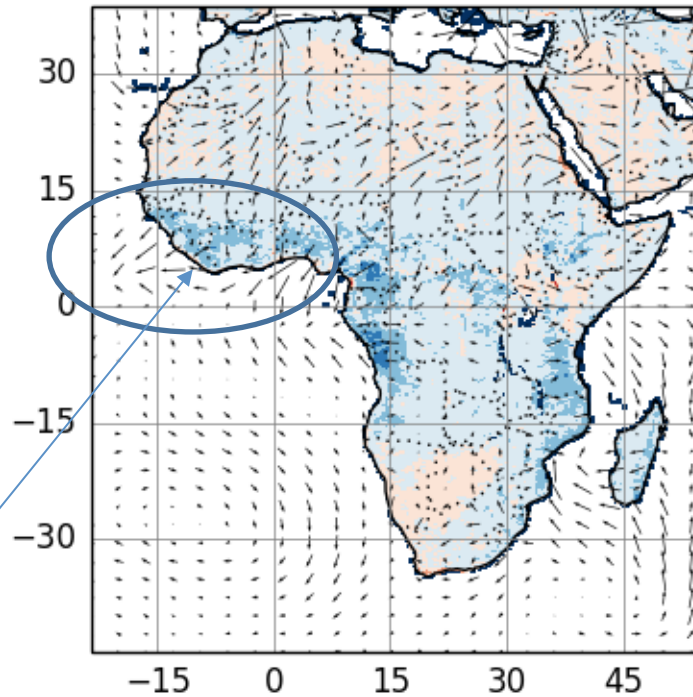
LH, Wm^{-2}



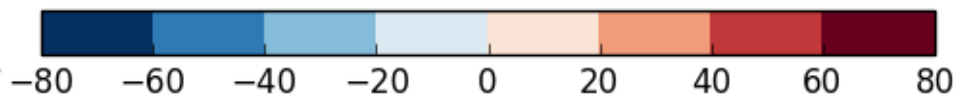
Tair 2-3K cooler

Dry season in the northern hemisphere

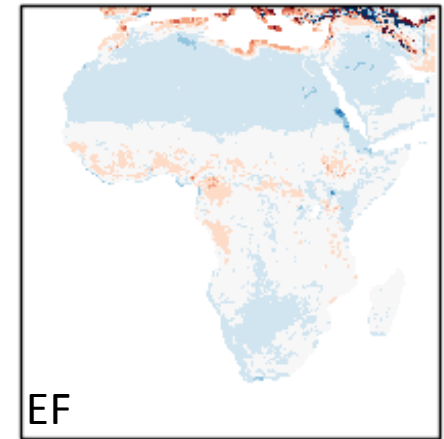
CP4 - R25



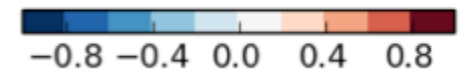
SH, Wm^{-2}



CP4 - R25



EF



— $\frac{2m}{s}$

Summary

- Do we see a difference in the land surface response in these idealised experiments?
- Yes we see a large difference in canopy interception losses:
 - CP4 is too low
 - R25 too high along west coast and east Africa.
- We see an impact on dry season evaporation, both in West Africa and across central Africa with additional $\sim 30 \text{ Wm}^{-2}$
- There is also a model structural effect due to the sub-grid distribution of rainfall which is something for development.
- Canopy evaporation has masked the effects of the land surface in the rainy season.
- Thank you!

