



# OBSERVATIONAL EVIDENCE FOR SOIL MOISTURE PATTERNS AFFECTING DEEP CONVECTION IN MATURE MESOSCALE CONVECTIVE SYSTEMS (MCS)

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2: National Centre for Earth Observation, UK.

# Surface sensitivity of MCS: afternoon-initiation case

Widely evaluated in models / observations:



Light mean wind



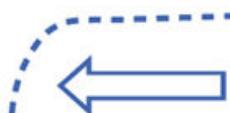
Convection is fostered on the dry side of convergence zones of land-surface boundaries

Max. CAPE, max.  $\Theta_e$ , convergence line

Deep, weak current



Shallow, strong current



+ ~2 hours



Cool, moist soil

Warm, dry soil

~10 km

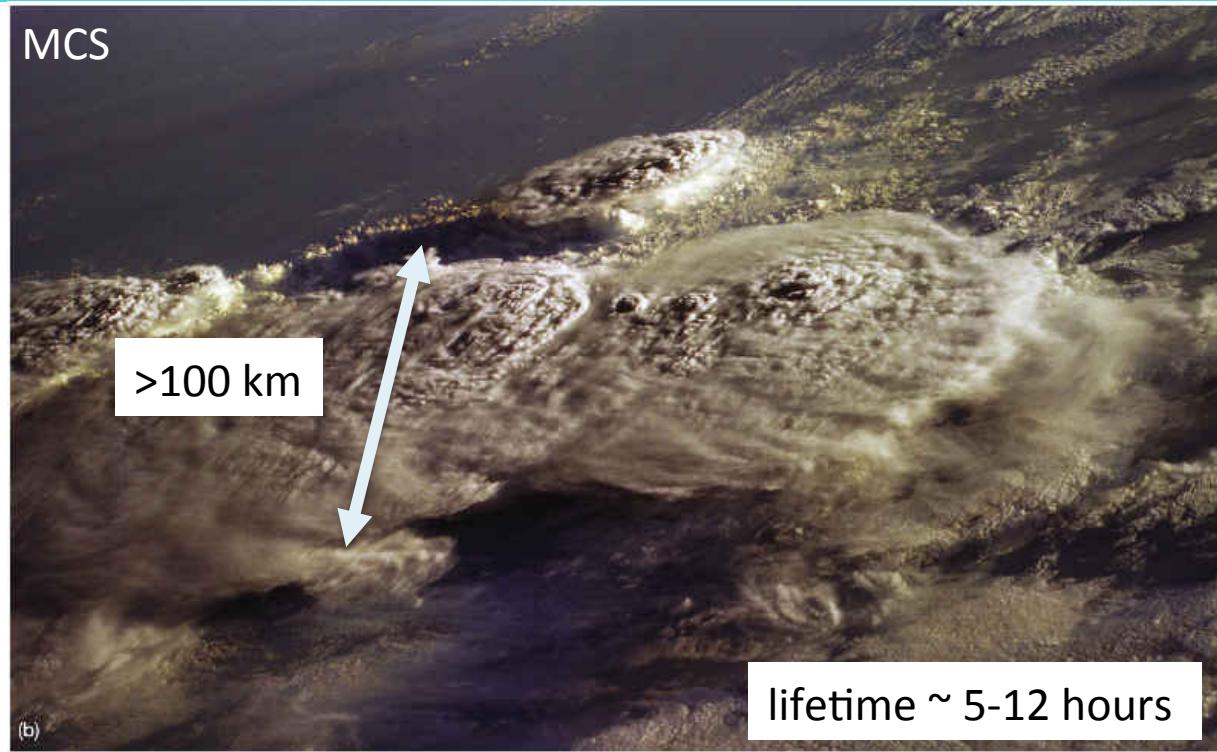
Cool, moist soil

Fully organised MCS

Taylor et al., Nat. Geosci. (2011)

# Surface sensitivity of MCS: the messy mature stage

MCS



ISS, NASA

- Low-level jets and elevated triggering
- Storm-induced gravity waves and cold pool dynamics

## Surface sensitivity?

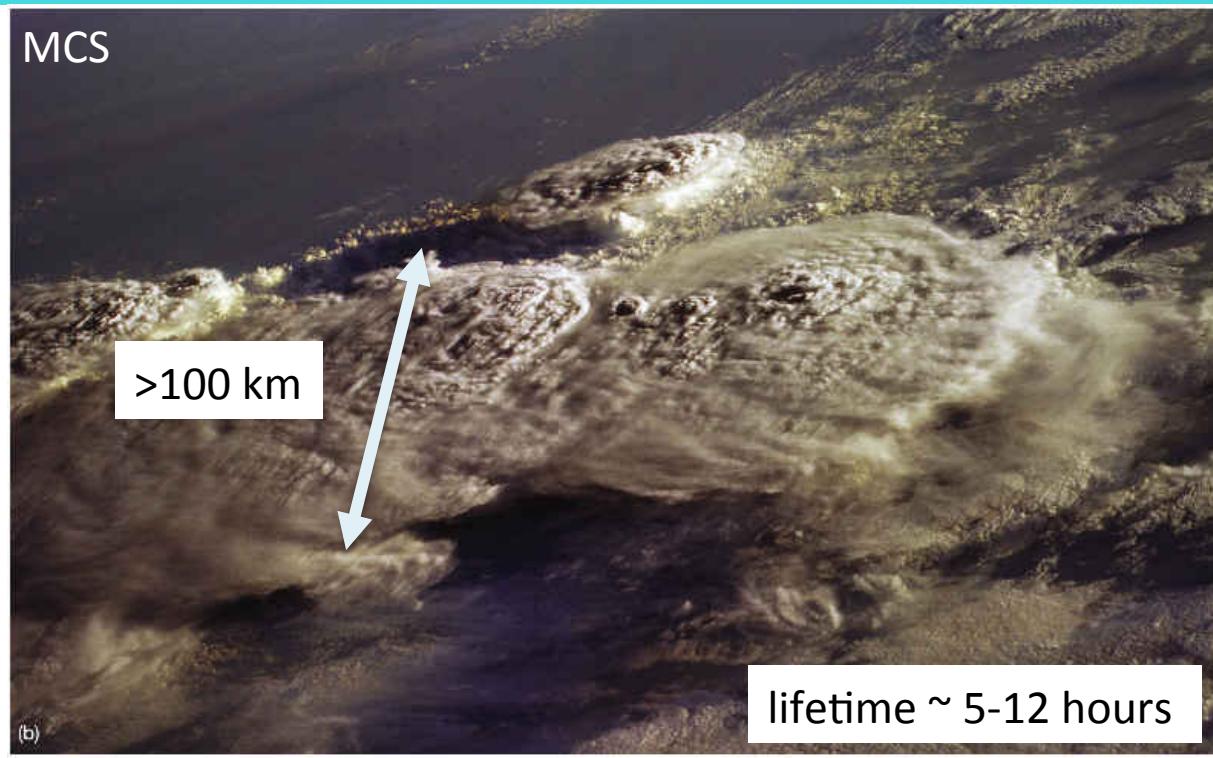
Models suggest:

- Intensification over wet soils / moist boundary layer (CAPE)

- Do soil moisture conditions affect MCS propagation and can we observe it?
- If yes, at what scales?

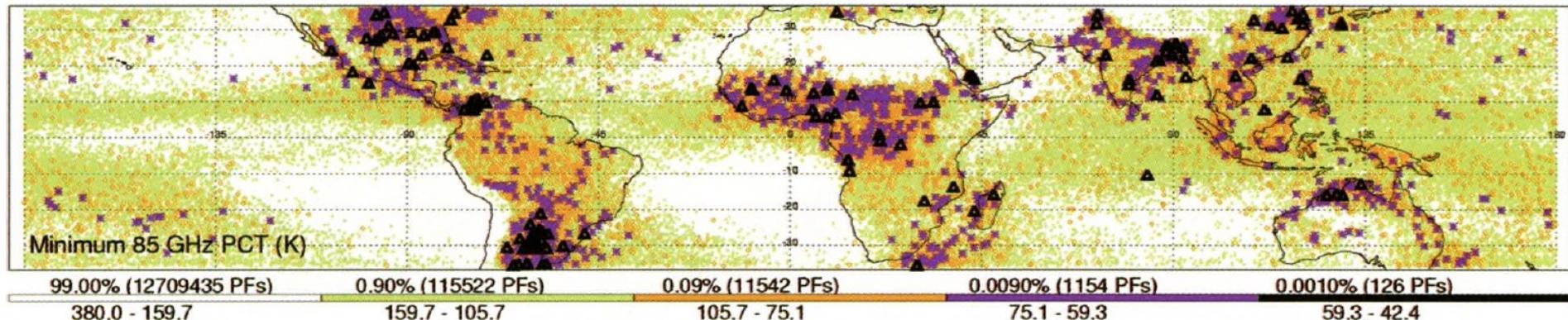
# Surface sensitivity of MCS: the messy mature stage

MCS



ISS, NASA

Mesoscale convective systems: some of the most intense storms on Earth

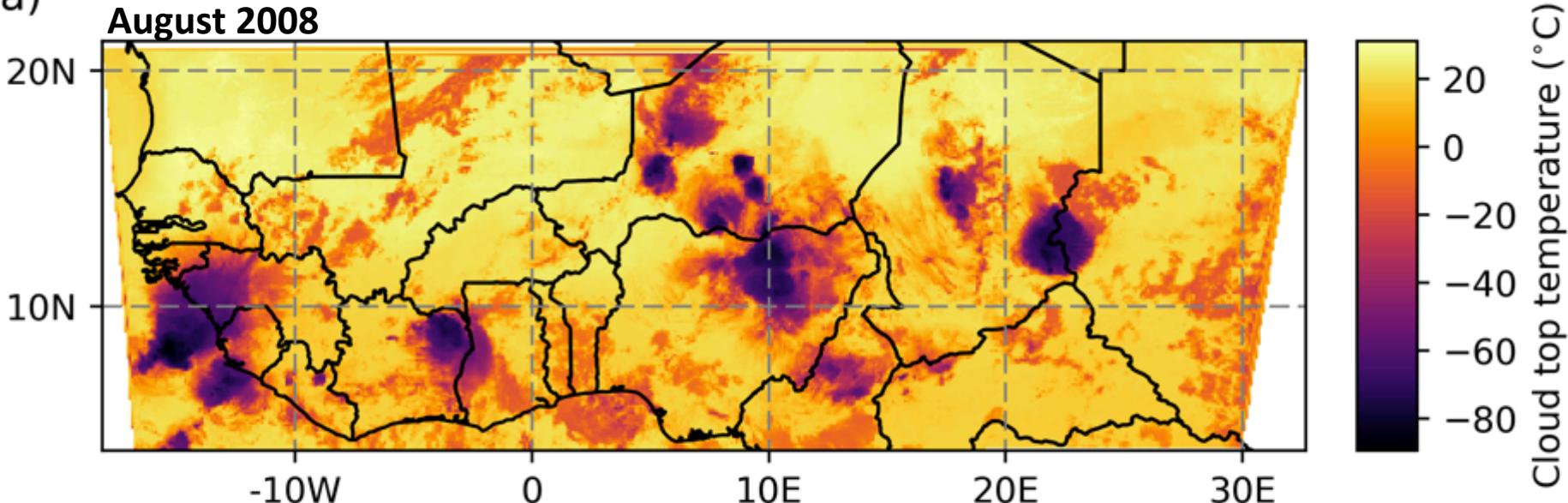


# Observing MCS during the West African monsoon

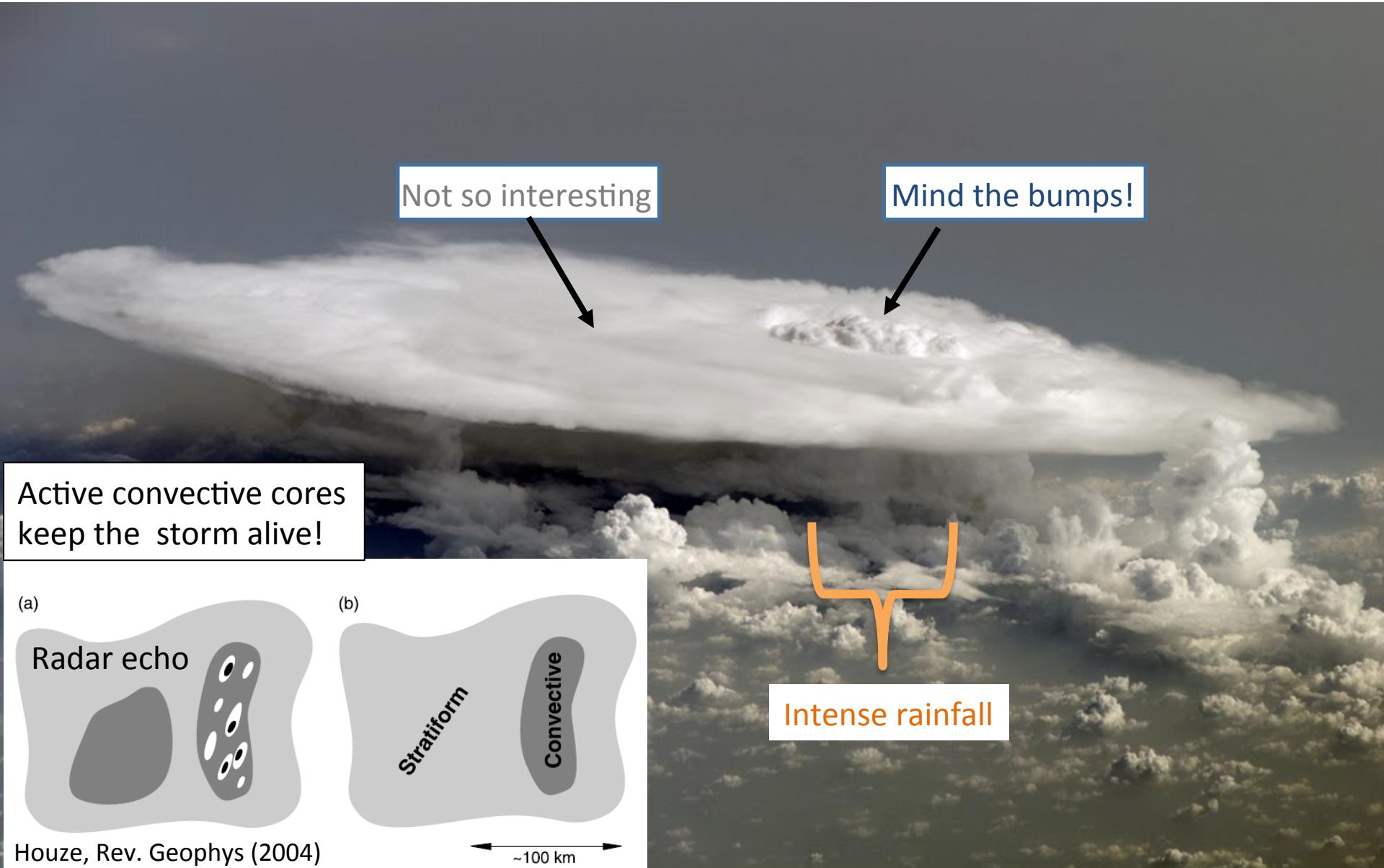
Meteosat Second Generation cloud top temperatures (brightness temperatures < -40C)  
**Used here:** Jun-Sep 2006-2010, hourly intervals, 5km

MCS > 15,000 km<sup>2</sup>, -> **90% of extreme rainfall occurrence and 50-90% of annual rain**

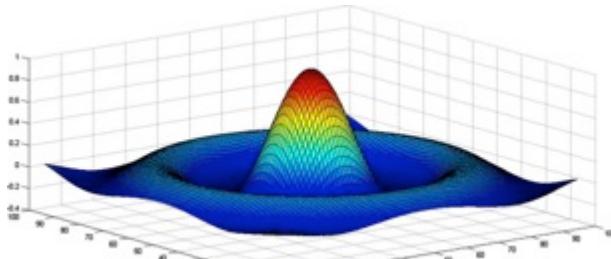
a)



# Deep convection: connecting lower & upper troposphere



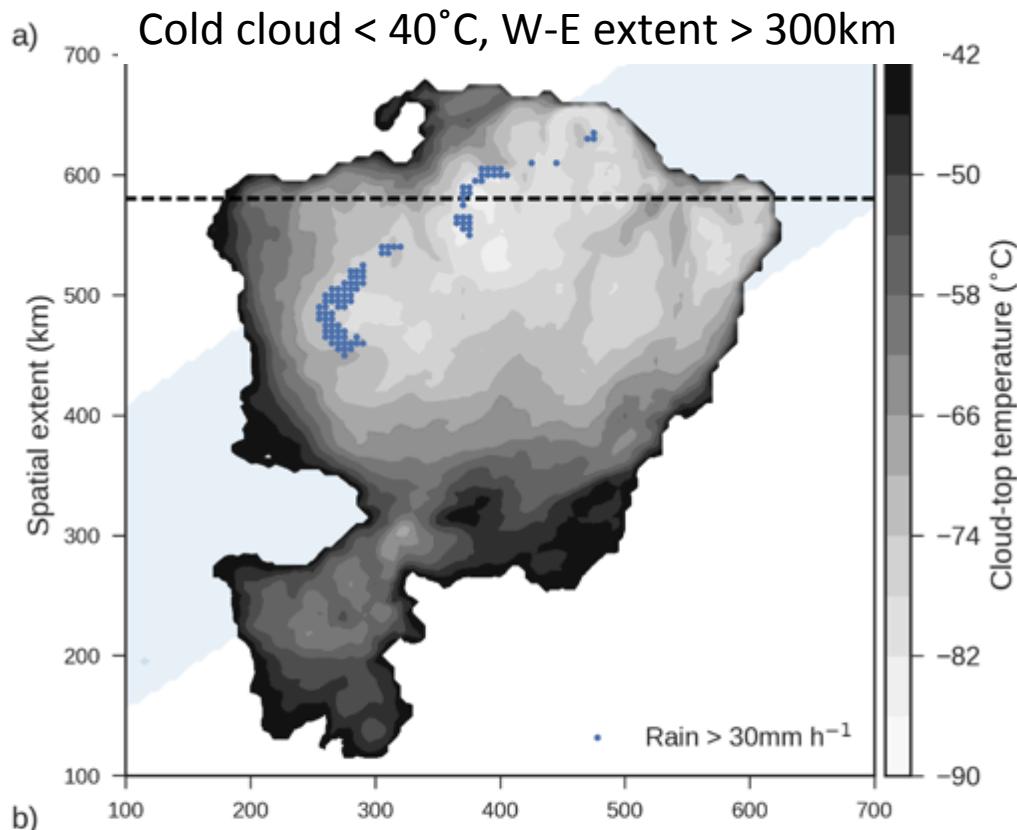
# Detecting convective cores from cloud top temperatures



2D Mexican hat wavelet

We identify the centre point of such cold “convective cores” by performing a **wavelet scale decomposition of cloud top temperatures**

(Klein et al., JGR-A, 2018)

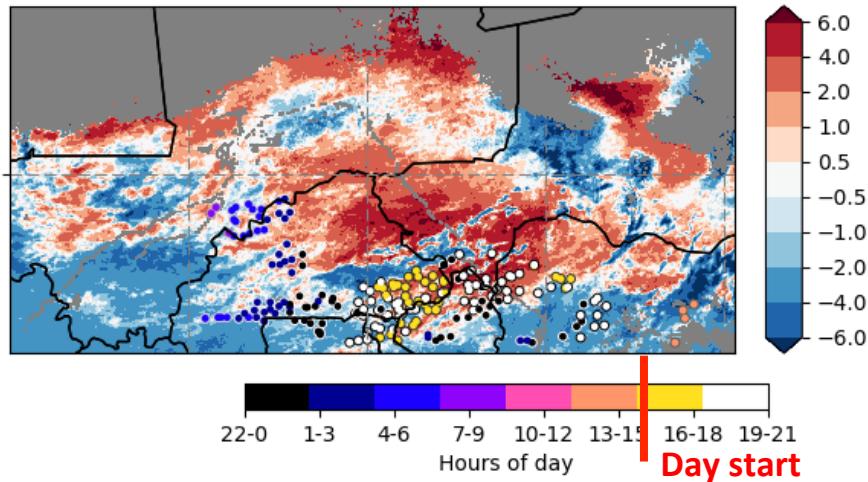


# Matching up convective cores and wet/dry surface conditions at different scales

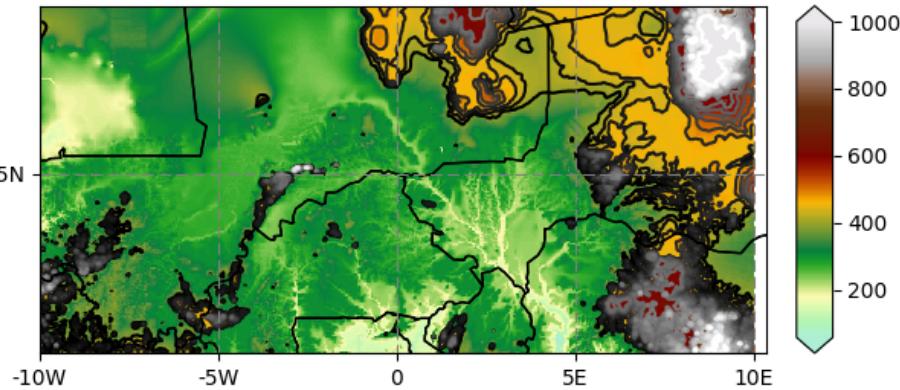
Scale decomposition of Meteosat land-surface temperature anomalies (LSTA):  
17-0500UTC convective cores and preceding LSTA (06-1600UTC)

-> LSTA is a good proxy for soil moisture in the semi-arid Sahel

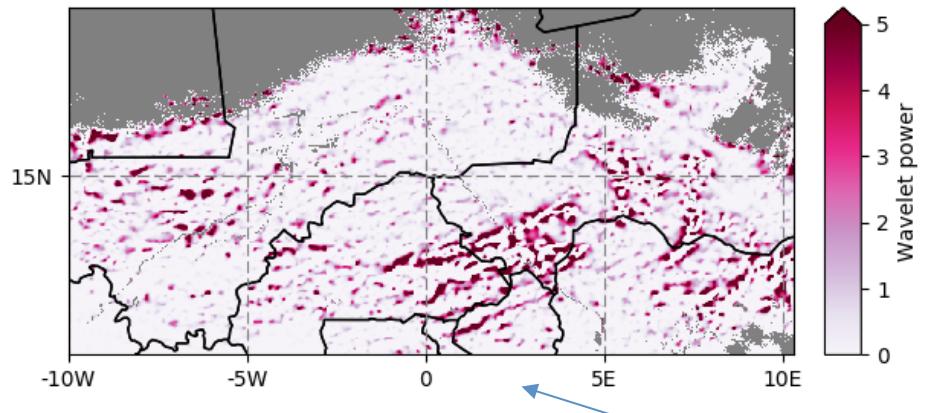
26-06-2008: 0630-1600UTC LSTA & afternoon/nighttime cores



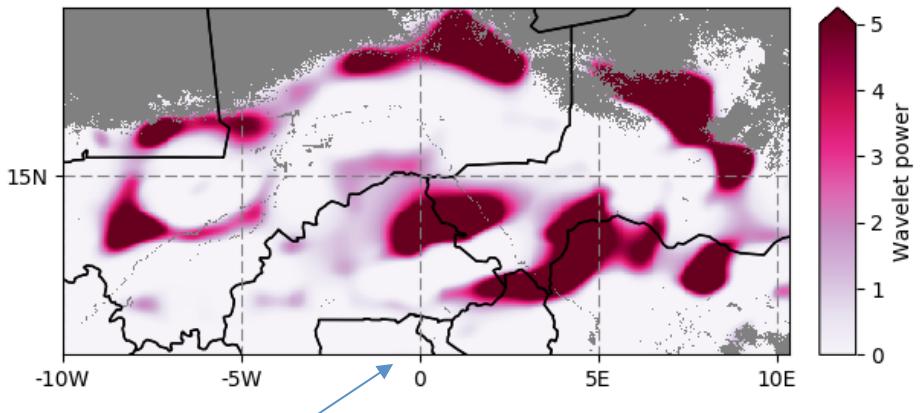
Domain topography, total height difference: 342m



Surface scales 9 - 30km



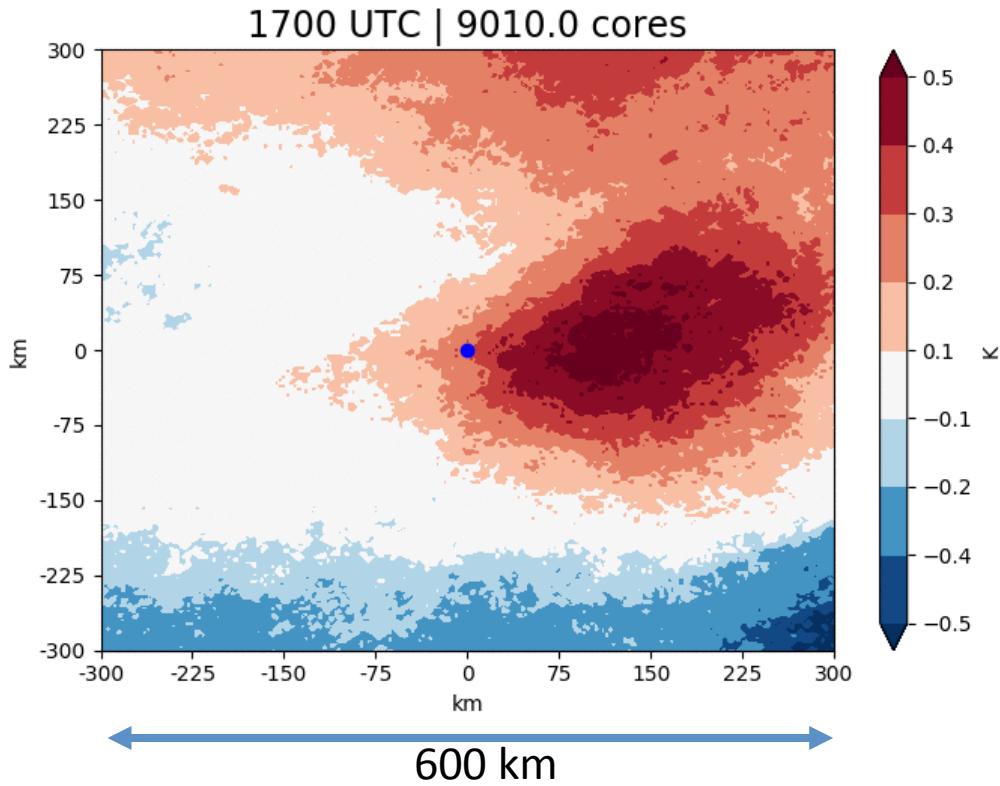
Surface scales 80 - 250km



Positive amplitudes only

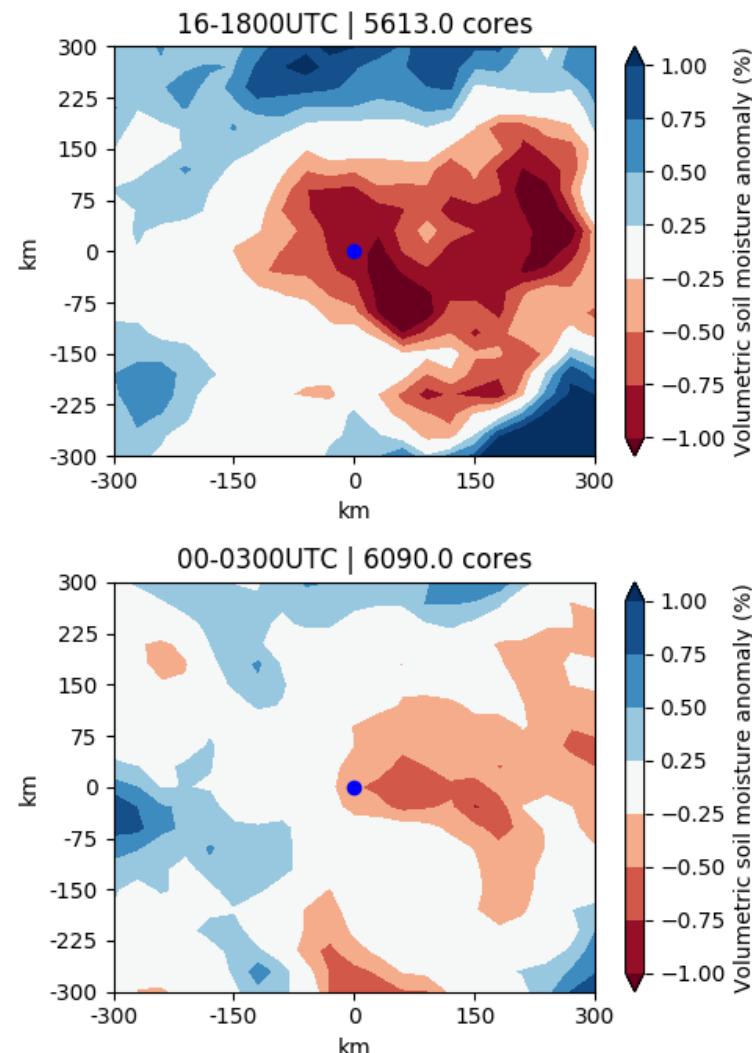
# Composites of surface conditions centred on convective cores

## Seasonal land surface temperature anomalies



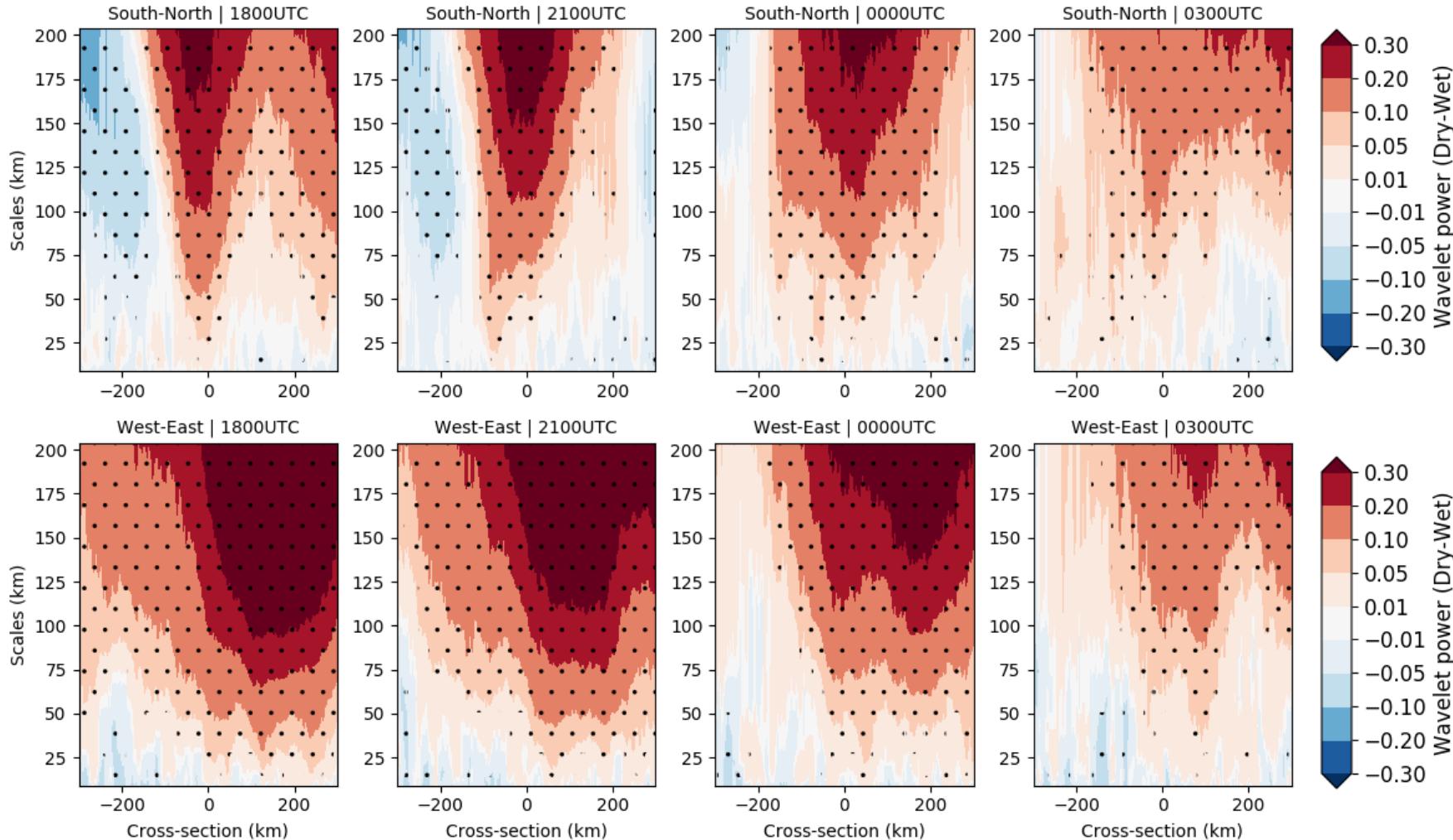
- ‘stationary’ ~2-3 hours warm upstream feature
- Persistent dry signal into nighttime hours

## Regional soil moisture anomalies (AMSR-E, $0.25^\circ$ ) centred on intense rainfall ( $> 20$ mm, TRMM radar)



# Scales of surface features throughout the day

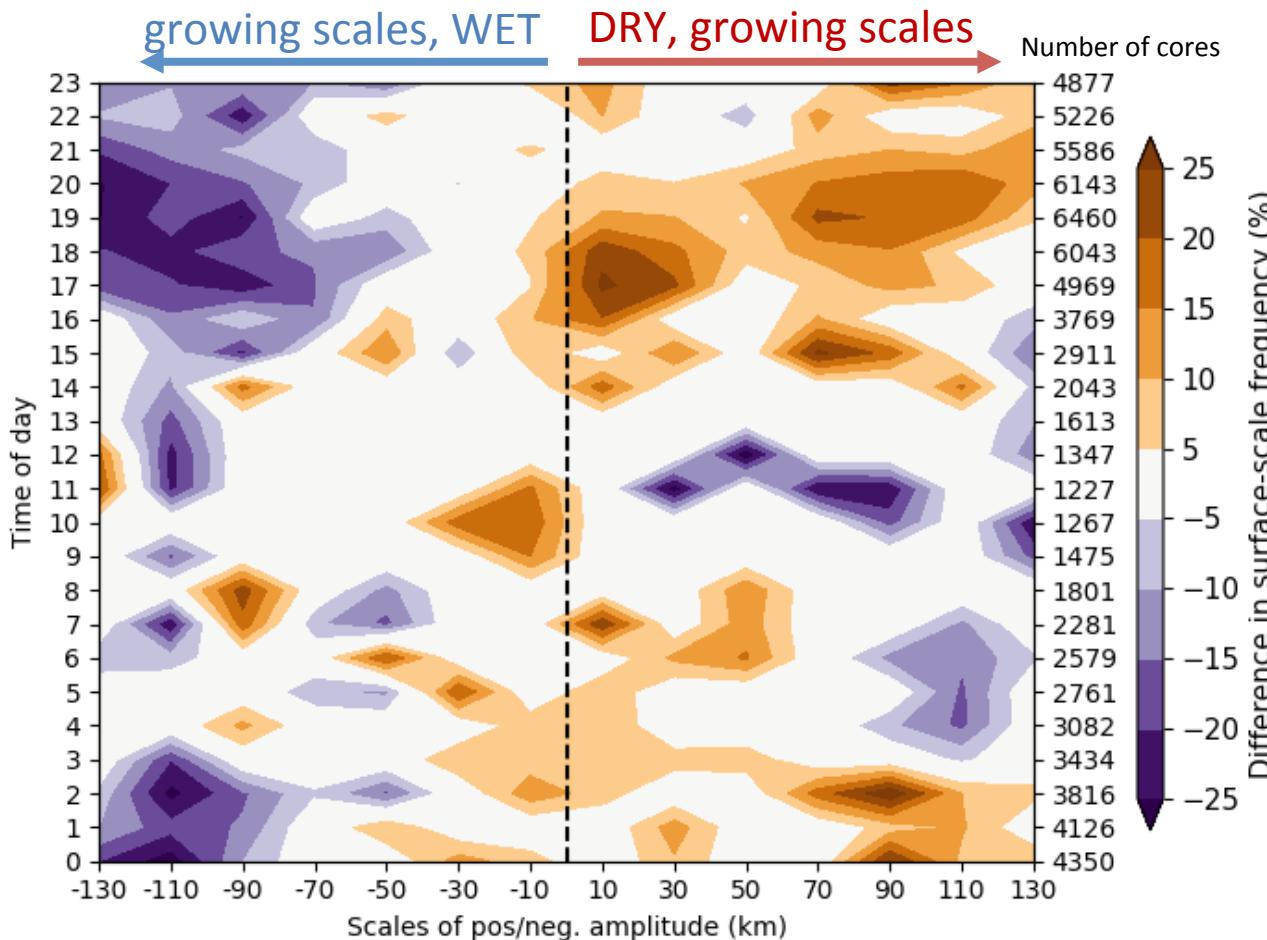
Difference of wavelet powers for dry and wet LST anomalies centred on convective cores:



- Strong meridional temperature gradient down to ~25km as part of the larger W-E feature  
-> propagation north of preceding-day storm tracks?
- Minimum significant scale of the warm feature is increasing through the night

# The effect of surface features at different scales on the frequency of convective cores over the diurnal cycle

Difference in frequency of co-located scales in comparison to the total scale distribution:



**Shown scales represent the scale of maximum wavelet power for the surface patch at the core location**

10%-value: 1 in 10 convective cores could be affected

- Afternoon maximum similar to storm initiation case
- Large wet features > 80km decrease core frequency, even for nocturnal MCS
- Surface scale increase is consistent with increasing influence of advection

# Summary

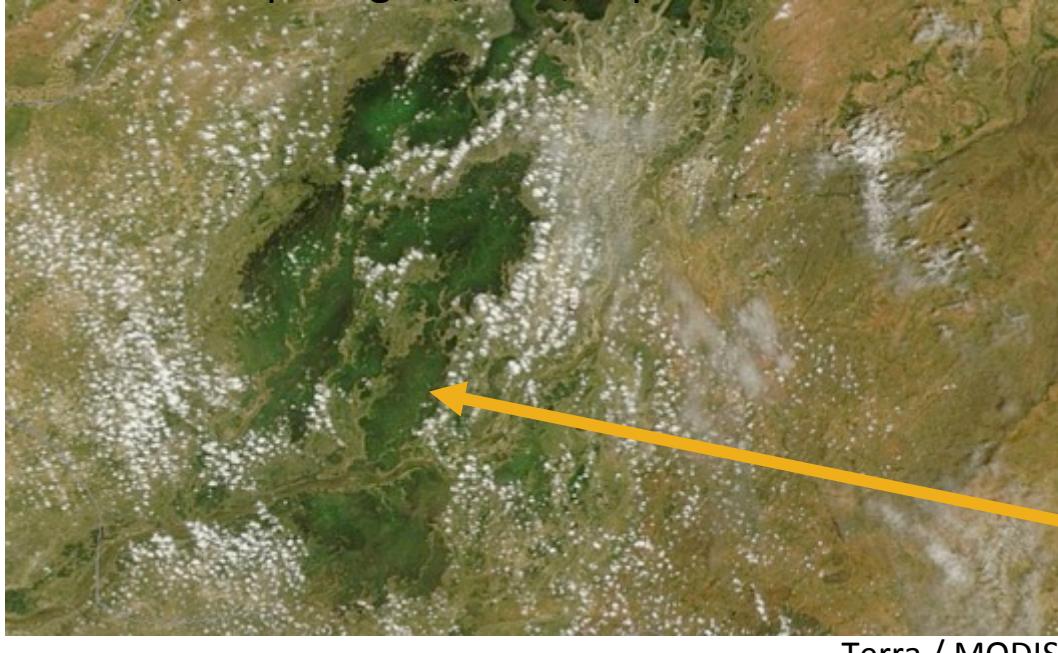
- We used a wavelet scale decomposition on cloud and surface features to evaluate sensitivities of convective cores within mature MCSs to surface conditions
- A dry-patch enhancement dominates even for nocturnal MCSs until ~0200UTC and is strongest in the afternoon
- The enhancement over dry patches follows a diurnal cycle and might reflect the smoothing of surface-induced variability in the planetary boundary layer
- No sign of a co-located or upstream positive wet feedback
- Potentially useful for forecasts of MCS propagation tracks
- What do models say?

# Questions?

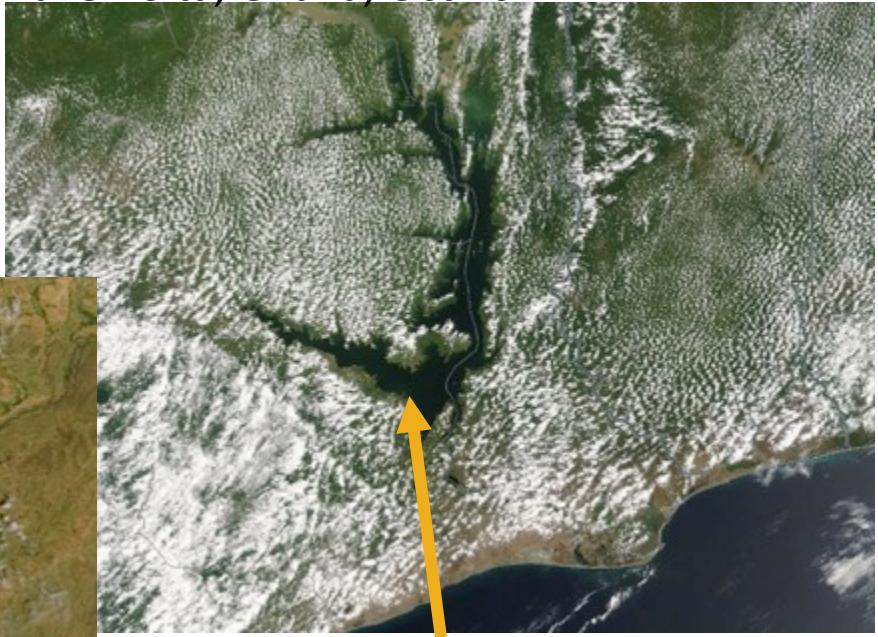
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# The surface can affect where clouds form

Wetland, Mopti region, Mali, Sep 2013



Lake Volta, Ghana, Oct 2014

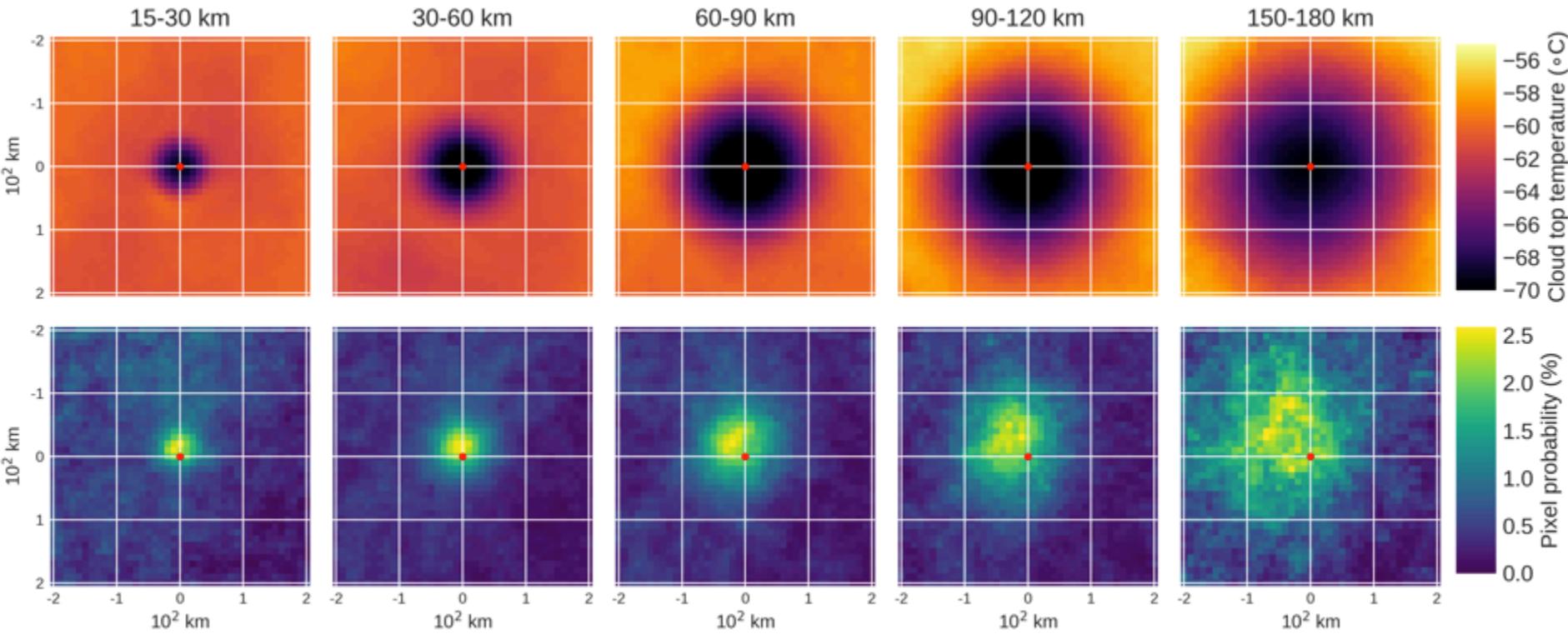


Terra / MODIS

- Cooler surface
- Small sensible heat flux
- Shallower boundary layer

# Scales of ‘convective cores’ and co-located TRMM PR rainfall

Composites centred on convective cores (<-50°C): temperature & heavy rain probability



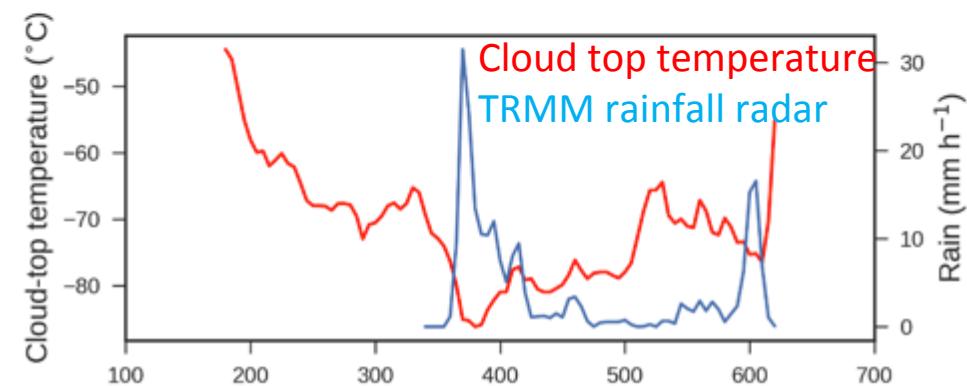
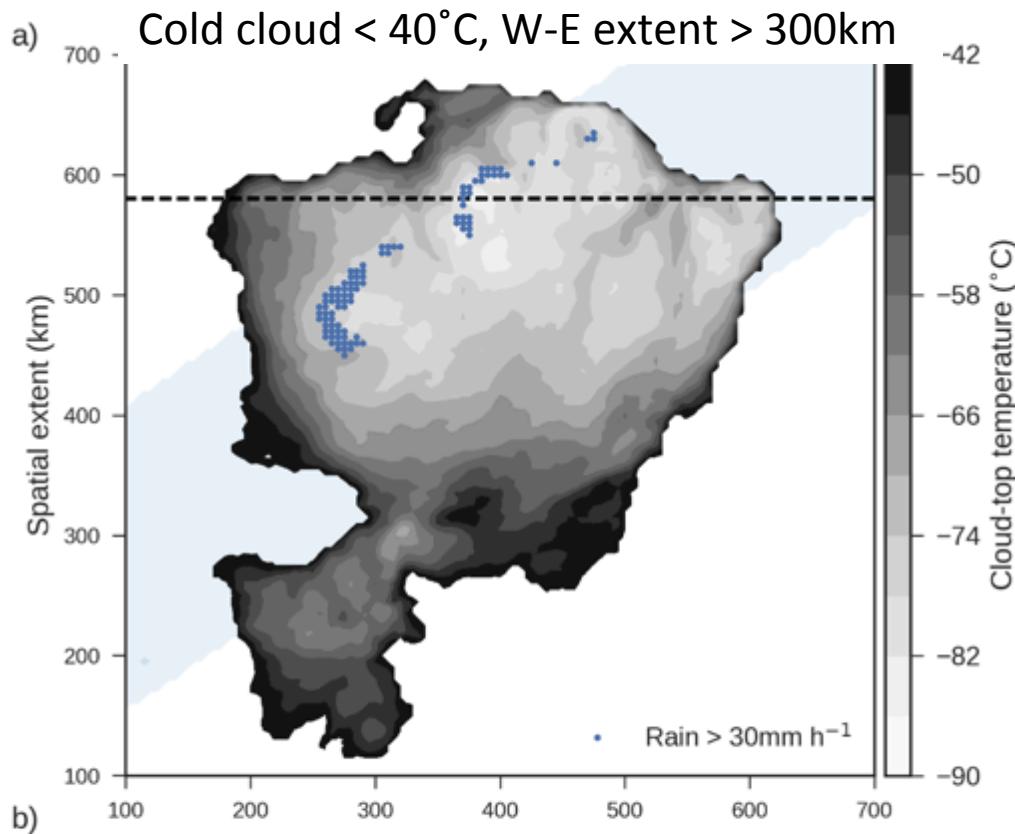
In 15km radius surrounding centre points for scales  $\leq 60\text{km}$ :

- 1 in 3 sub-cloud features coincide with convective rain ( $> 8 \text{ mm/h}$ )
- 8.3% are associated with heavy rain (99<sup>th</sup> centile,  $> 30 \text{ mm/h}$ )

(Klein et al., JGR-A, 2018)

# Detecting convective cores from cloud top temperatures

- Cloud-top temperatures indicate the height of the cloud top
- Coldest temperatures are often co-located with strong updrafts and intense convective rainfall



# Thank you for your attention!



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# Sensitivities of nocturnal MCSs?

- Idealised simulations suggest preference for wet soils / moist boundary layer
- Effect of low-level jets at nighttime and elevated triggering (no contact to PBL anymore)
- 

Birth and decay processes of convective cells cause storm size to grow

Average 12-14m/s propagation speed -  $\sim 43\text{km/h} \rightarrow 150\text{km} \sim 3\text{h distance}$   
 $15,000\text{km}^2 = 70\text{km radius}$

$10\text{m/s} = 36\text{km/h} \rightarrow 150\text{km} \sim 4\text{h distance}$

MCS grow to  $15,000\text{km}^2$  within  $\sim 4\text{hours}$ , do we see initiation point?

What about at midnight and after?