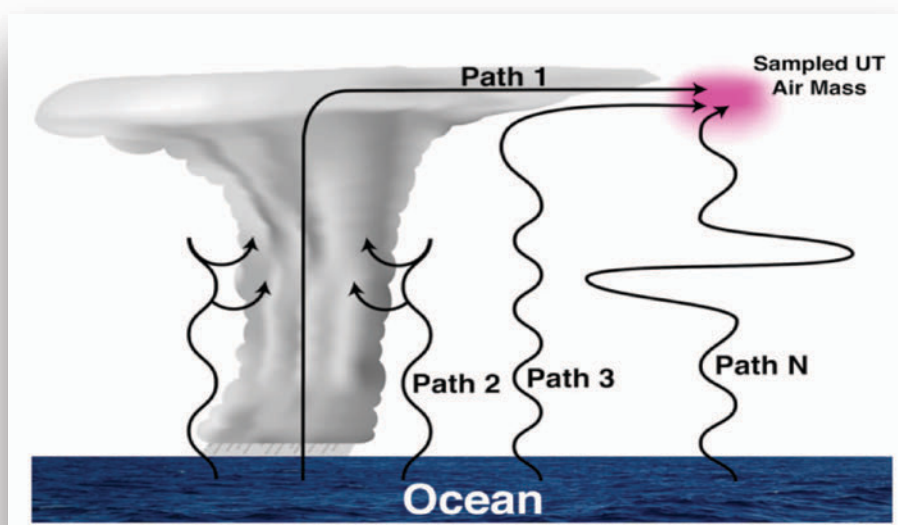
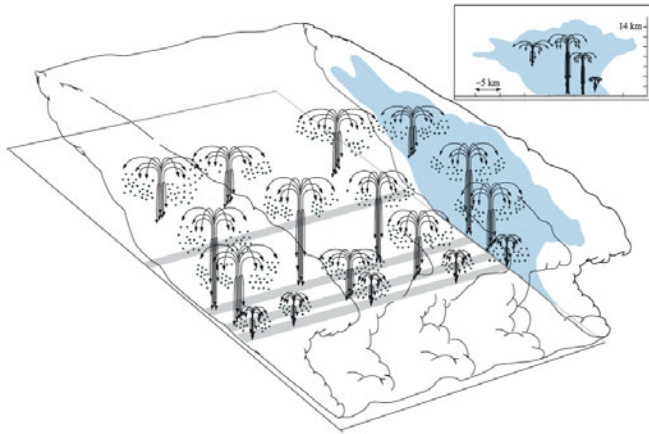


## *On the Use of Chemical Tracers to Diagnose Convective Transport Pathways from the PBL to the UT/LS*

**Z. Johnny Luo**<sup>1</sup>, L. L. Pan<sup>2</sup>, E. L. Atlas<sup>3</sup>, S. M. Chelpon<sup>1</sup>, S. B. Honomichl<sup>2</sup>, E. C. Apel<sup>2</sup>,  
R. S. Hornbrook<sup>2</sup>, and S. R. Hall<sup>2</sup>

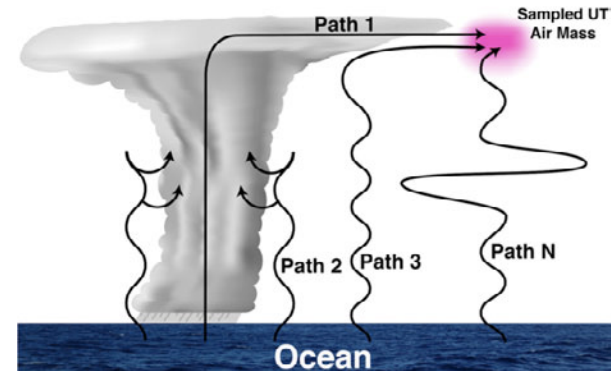
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### Travel time from the PBL to the UT

- Convective cores (10 m/s):  
 $15 \text{ (km)} / 10 \text{ (m/s)} = 25 \text{ min}$
- Outside of the core/inside clouds (1 cm/s):  
 $15 \text{ (km)} / 1 \text{ (cm/s)} = 17 \text{ days}$
- Large-scale ascent (1 mm/s):  
 $15 \text{ (km)} / 1 \text{ (mm/s)} = 174 \text{ days}$

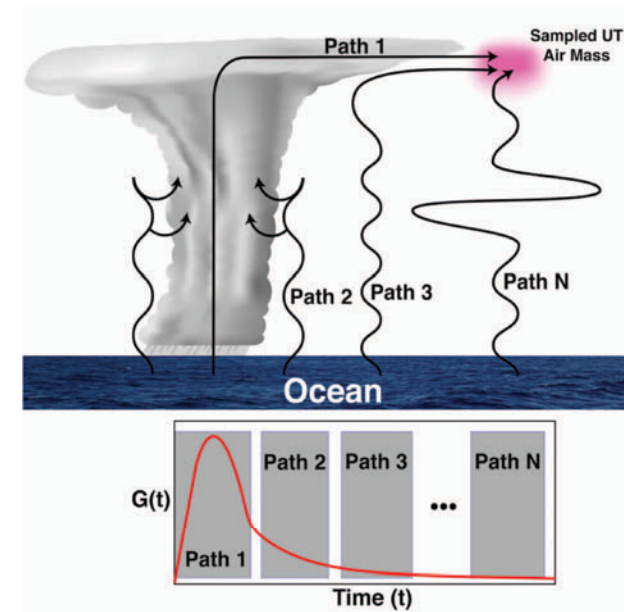


### Objectives of the study:

- quantify the contributions of these different transport pathways using tracer measurements, and
- develop observation-based metric for evaluating model representation of convective transport.

## Outlines

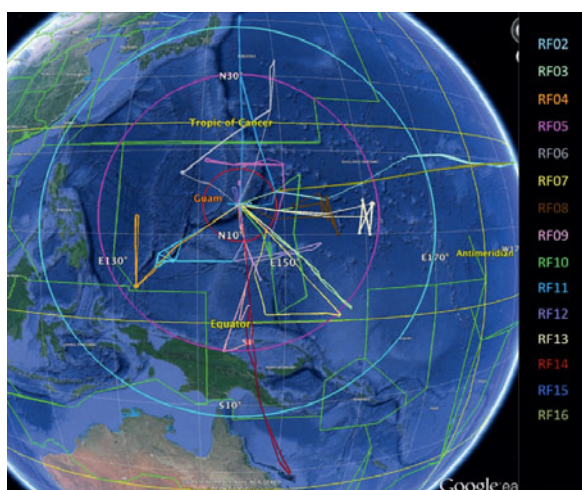
1. Behavior of chemical tracers from a recent field campaign - CONTRAST
2. A metric for quantifying contributions of different transport paths – Transit Time Spectrum
3. Summary & implications



**The idea:** UT air mass is a "cocktail" coming from different transport paths. Because each path has a characteristic transit time (from PBL to UT), we can use a suite of tracers with different lifetimes to probe the relative contributions of these paths.



## The Convective Transport of Active Species in the Tropics (CONTRAST) Experiment



### Scientific Objectives:

- Characterizing the influence of deep convection on the chemical composition and the photochemical budget of  $O_3$  at the level of convective outflow
- Evaluate the budget of organic and inorganic bromine and iodine in the TTL
- Investigate transport pathways from the oceanic surface to the tropopause



ID	Instrument	Observed Trace Gas	Atmospheric Lifetime [Days]
1	TOGA	Acetaldehyde	0.8
2	TOGA	Dimethyl Sulfide (DMS)	1
3	TOGA	nButane	1
4	TOGA	Tetrachloroethylene	1
5	TOGA	Formaldehyde	1.4
6	TOGA	Ethyl Benzene	2
7	TOGA	Methyl Ethyl Ketone (MEK)	3
8	TOGA	Isopentane	3.2
9	TOGA	Methyl Iodide	4
10	TOGA	Isobutane	6
11	TOGA	Benzene	10
12	TOGA	Propane	10.5
13	TOGA	Methanol	12
14	TOGA	Bromoform	15
15	TOGA	Ethyl Nitrate	15
16	TOGA	Chlorobenzene	20
17	TOGA	Methyl Vinyl Ketone (MVK)	30
18	TOGA	Dibromomethane	94
19	TOGA	Dichloromethane	109
20	TOGA	Methyl Chloroform	109
21	TOGA	Chloroform	112
22	TOGA	Acetonitrile	174
23	TOGA	Methyl Bromide	292
24	TOGA	Carbon Tetrachloride	9490

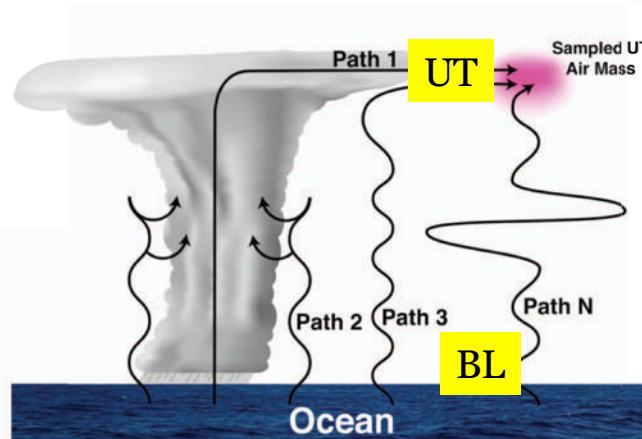
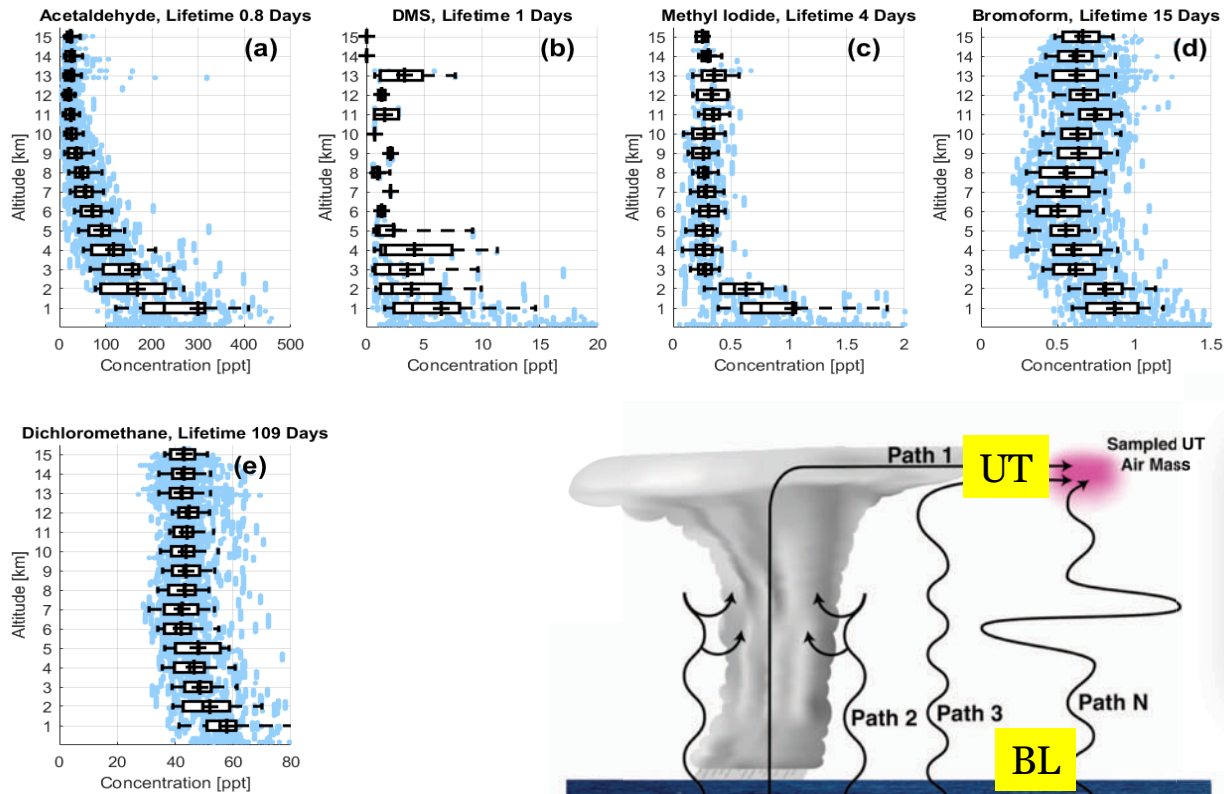


We define a ratio:

$$\mu_i^* \equiv \frac{m_i(UT)}{m_i(BL)}$$

, where i is the tracer index

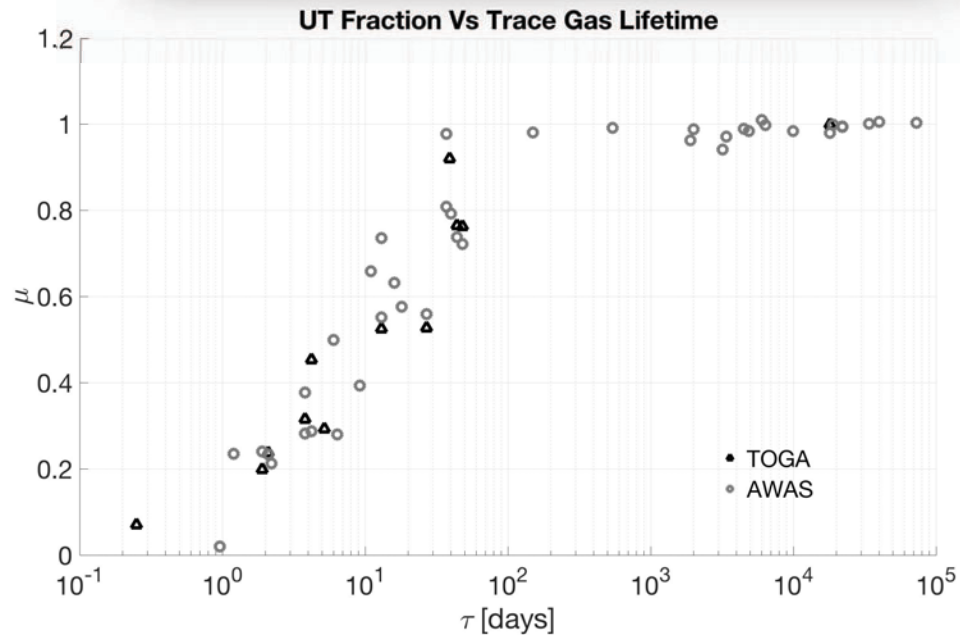
TOGA Trace Gas Measurements, Vertical Profiles



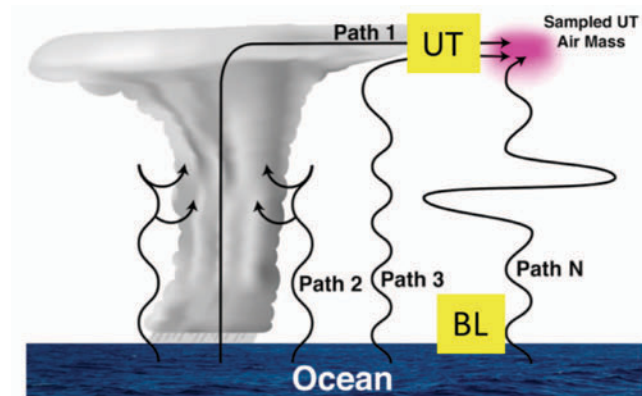
Sampled UT air mass is a "cocktail" from different transport paths

We define a ratio:  $\mu_i^* \equiv \frac{m_i(UT)}{m_i(BL)}$ , where i is the tracer index

42 tracers observed during the CONTRAST field campaign over the TWP WP

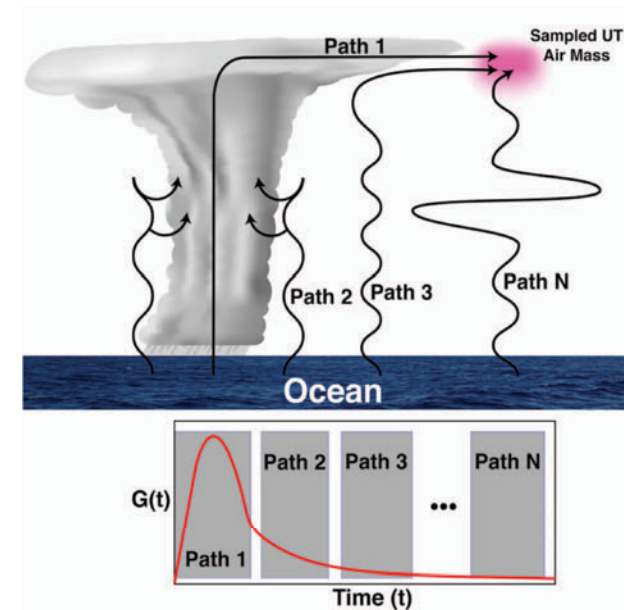


How to extract some quantitative information about transport paths using the  $\mu^*$ - $LT$  relation?



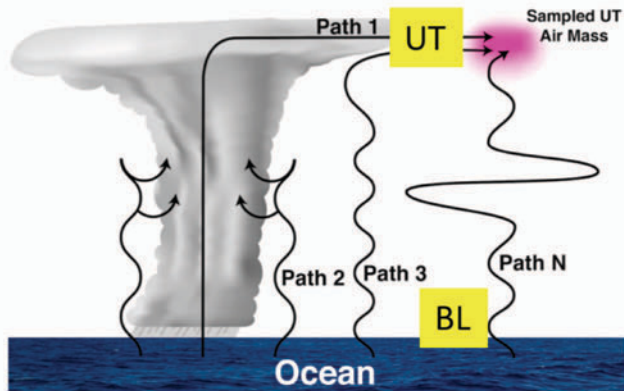
## Outlines

1. Behavior of chemical tracers from a recent field campaign - CONTRAST
2. A metric for quantifying contributions of different transport paths – Transit Time Spectrum
3. Summary & implications



**The idea:** because different paths have different characteristic time scale, we may use a suite of tracers with different lifetime to probe these paths. Short-lived tracers are only sensitive to fast paths, while long-lived tracers can go through all paths.





$G(t)$  is essentially a ***weighting Function*** of the transport paths (or a formula for the UT “cocktail”).

Assumptions:

- 1) Tracer mixing ratio decays exponentially with time following its lifetime ( $\tau_i$ )
- 2) MBL is the source of the transport and UT is the destination.

Along a single path

$$m_i(UT) = m_i(BL)e^{-t/\tau_i}$$

Along a multitude of paths

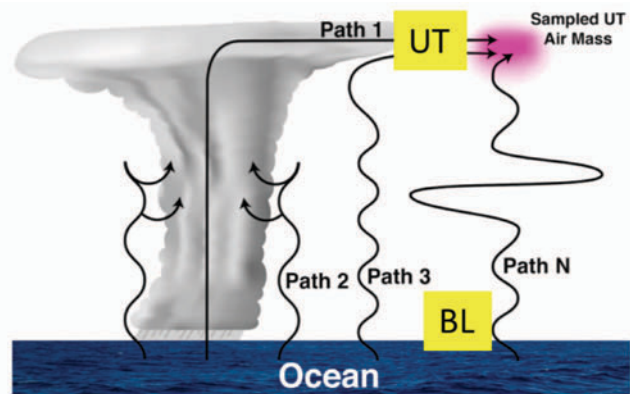
(each  $t$  represents a path)

$$m_i(UT) = m_i(BL) \int_0^\infty e^{-t/\tau_i} G(t) dt$$

$$\mu_i^* = m_i(UT) / m_i(BL)$$

$$= \int_0^\infty e^{-t/\tau_i} G(t) dt \dots\dots\dots (1)$$

$G(t)$  is a weight with which different transport paths (with different transit time  $t$ ) contribute to the sampled UT air mass. We call  $G(t)$  “***transit time spectrum***”.



Along a single path

$$m_i(UT) = m_i(BL)e^{-t/\tau_i}$$

Along a multitude of paths

(each t represents a path)

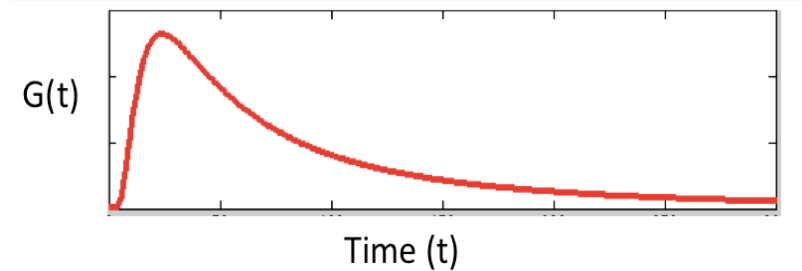
$$m_i(UT) = m_i(BL) \int_0^{\infty} e^{-t/\tau_i} G(t) dt$$

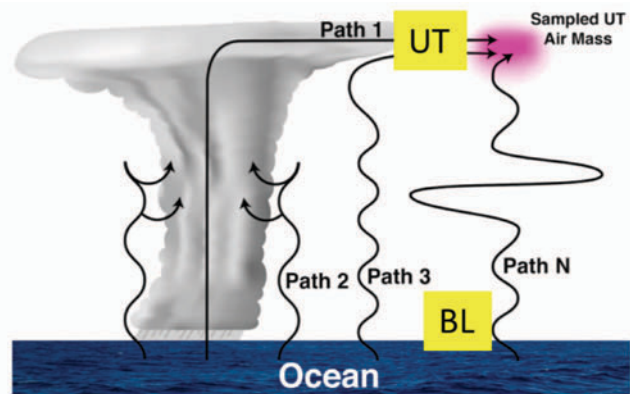
$$\mu_i^* = m_i(UT) / m_i(BL)$$

$$= \int_0^{\infty} e^{-t/\tau_i} G(t) dt \dots\dots\dots (1)$$

Equation (1) basically is the **Laplace transform** of G(t).

In principal, we could perform Inverse Laplace Transform to back out G(t). (Note that Eq (1) defines a matrix of i and t.)





Along a single path

$$m_i(UT) = m_i(BL)e^{-t/\tau_i}$$

Along a multitude of paths  
(each t represents a path)

$$m_i(UT) = m_i(BL) \int_0^{\infty} e^{-t/\tau_i} G(t) dt$$

$$\mu_i^* = m_i(UT) / m_i(BL)$$

$$= \int_0^{\infty} e^{-t/\tau_i} G(t) dt \dots\dots\dots (1)$$

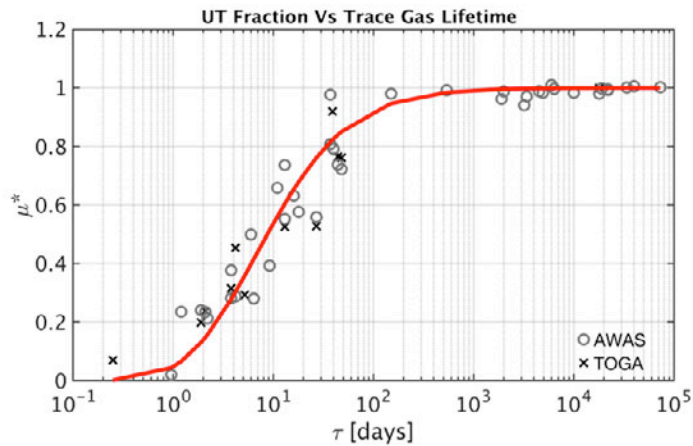
Equation (1) basically is the **Laplace transform** of G(t).

In principal, we could perform Inverse Laplace Transform to back out G(t).

We use some a priori knowledge to constrain the retrieval of G(t): assume the transport follows a *vertical diffusion model*, in which case G(t) has an analytical solution

$$G(t) = \frac{z}{2\sqrt{\pi Kt^3}} \exp\left(\frac{z}{2H} - \frac{Kt}{4H^2} - \frac{z^2}{4Kt}\right)$$

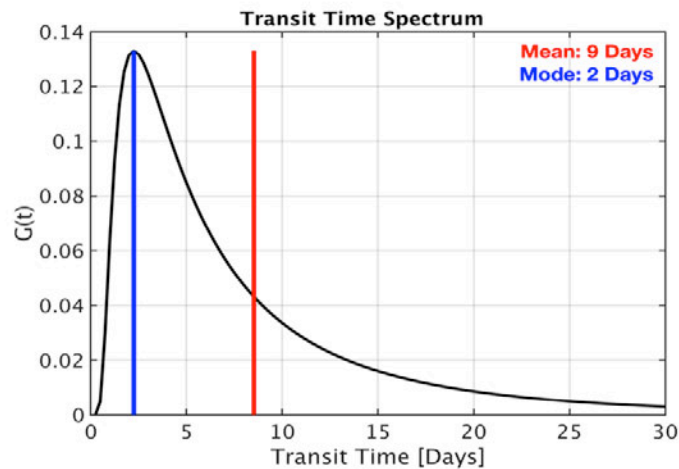
, z is height, H is scale height, K is the diffusion coefficient, and t is time



$$G(t) = \frac{z}{2\sqrt{\pi Kt^3}} \exp\left(\frac{z}{2H} - \frac{Kt}{4H^2} - \frac{z^2}{4Kt}\right)$$

Following Schoberl et al. (2005), we minimizing F using least square fit

$$F = \sum_{i=1}^N (\mu_i - \mu_i^*)^2$$

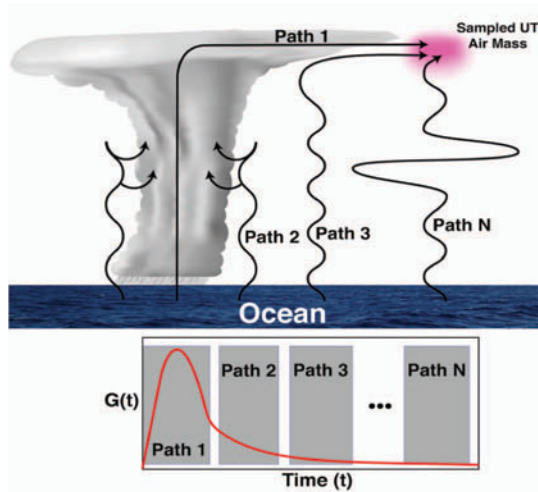


For the UT (10-13km)

Mode: 2 days

Mean: 9 days

$$\bar{T} = \int_0^{\infty} tG(t) dt \quad (\text{red line})$$

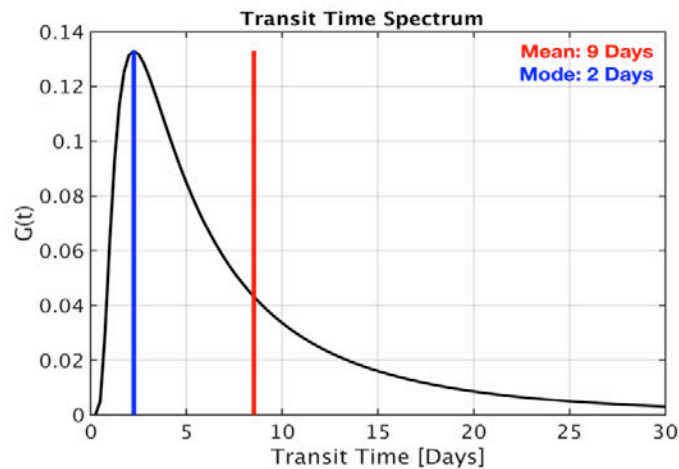


## A consistency check of the result

Masunaga and Luo (2016) estimated the convective mass flux ( $M_c$ ) is about  $0.005$ - $0.015 \text{ kg m}^{-2} \text{ s}^{-1}$  over the TWP region.

$$M_c = \rho_{\text{air}} W_{\text{mean}} \rightarrow W_{\text{mean}} \approx 0.8 - 2.4 \text{ cm/s}$$

$$T_{\text{transit}} = 13 \text{ km} / W_{\text{mean}} = 6 - 19 \text{ days}$$



For the UT (10-13km)

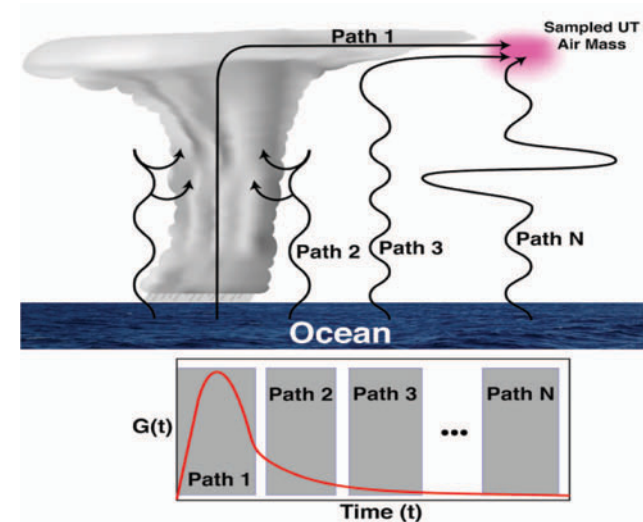
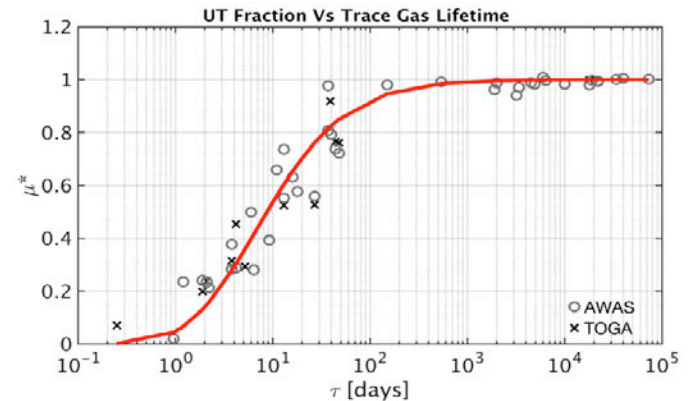
Mode: 2 days

Mean: 9 days

$$\bar{T} = \int_0^{\infty} tG(t) dt \quad (\text{red line})$$

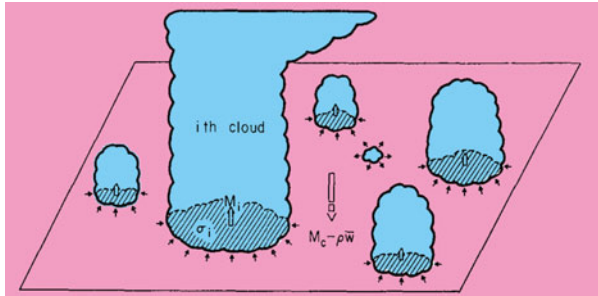
## Summary

- Chemical Tracers with different lifetimes can be used to probe convective transport pathways because an essential property of the transport paths is the characteristic transit time from the source (PBL) to the destination (UT)
- We demonstrated this concept by using a unique set of experimental data and a simple model, and defined an observation-based metric - **Transit Time Spectrum or  $G(t)$**  – for quantifying contributions from different transport paths.
- The mean convective transport time scale derived from tracers (9 days) is broadly comparable with the estimate based on convective mass flux.

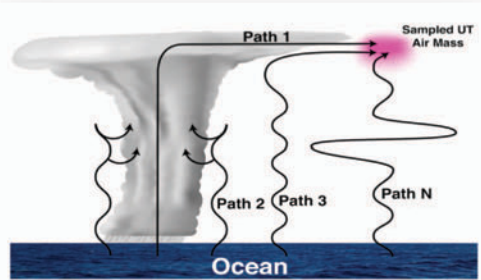
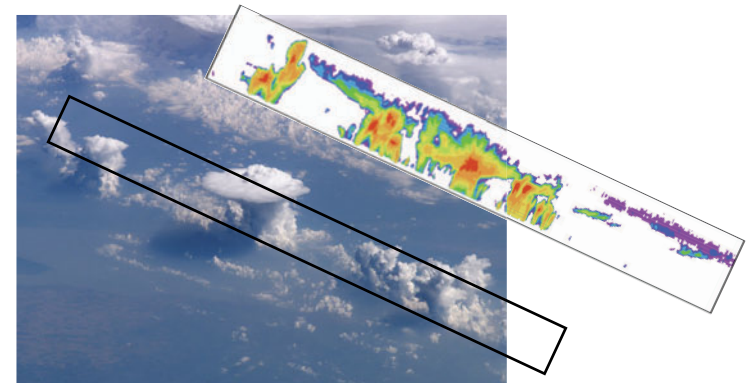


## Implication of the study

Formulation of convective parameterization is based on convective dynamics via convective mass flux



Global observations of convection mostly focus on observing hydrometeors



$G(t)$  contains rich information about convective dynamics and has the potential of becoming an effective diagnostic of the representation of convection and convective transport in global models.