



Runoff Processes in Alpine Catchments: Challenges and Opportunities

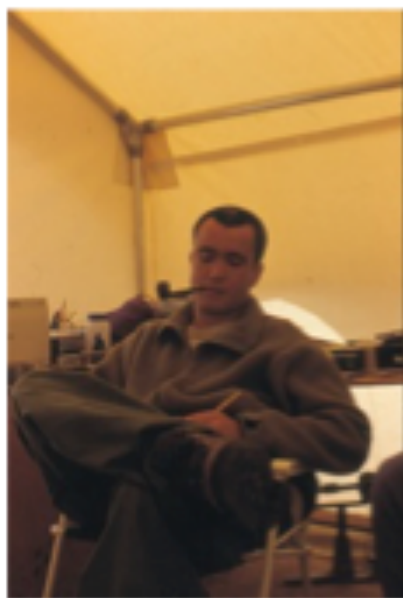
Sean K. Carey

School of Geography & Earth Sciences,
McMaster University,
Hamilton, Ontario, Canada





- Hydrologist and dabbler at McMaster University in Canada.
- Have been working in Wolf Creek, Yukon Territory Canada since 1995.
- Have also worked in British Columbia
- Interested in all things 'watery'. Interested in how 'cold' affects water and the environment. Also concerned with human impacts in 'remote' areas.



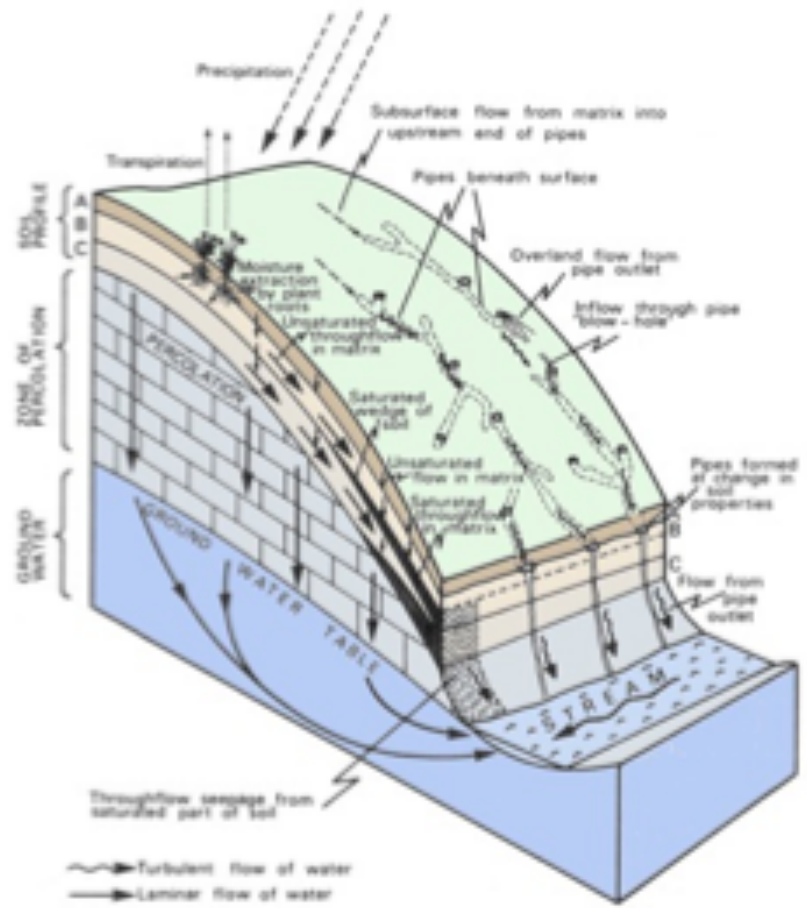
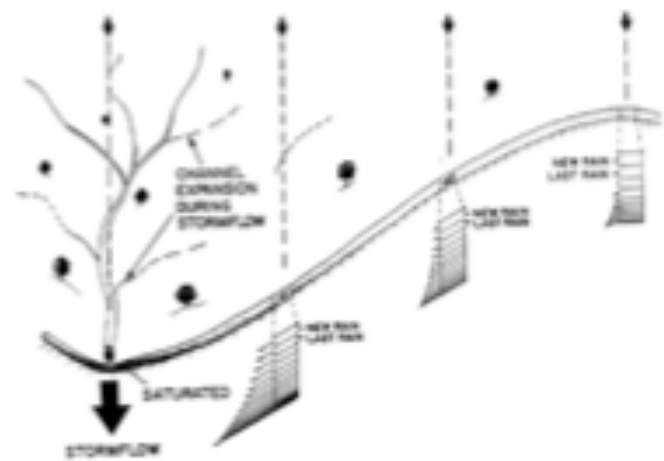


- From a streamflow generation perspective:
 - High energy
 - Complex geometry
 - Cryosphere influences (glaciers, snow, permafrost)

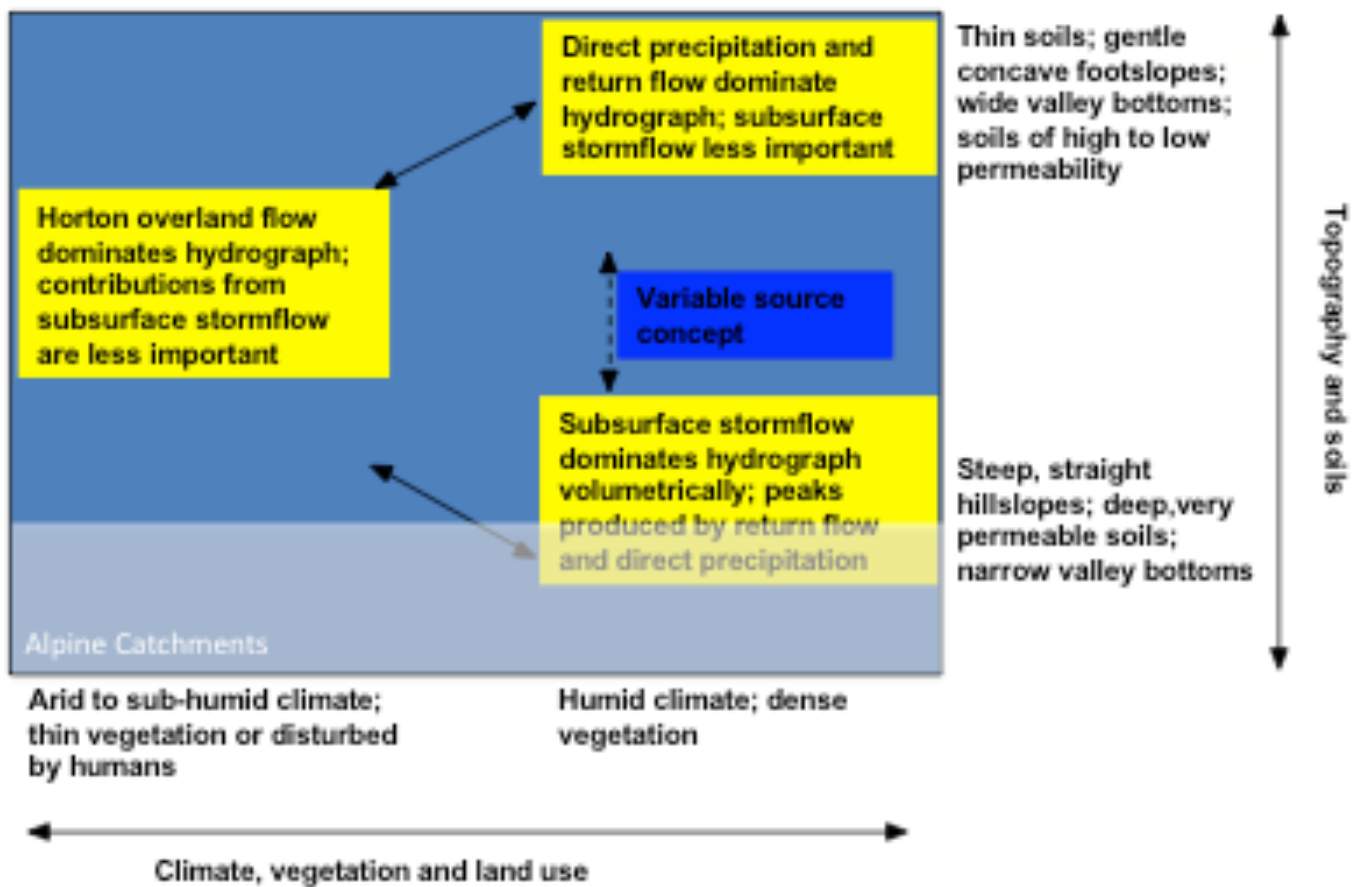




- Largely concerned with how water moved from hillslopes to streams
- Large literature beginning with Hortonian overland flow and moving to throughflow, saturated wedge, transmissivity feedback, etc.



From Kirkby (1978)



Runoff processes in relation to their major controls.

(From Dunne and Leopold, 1978)

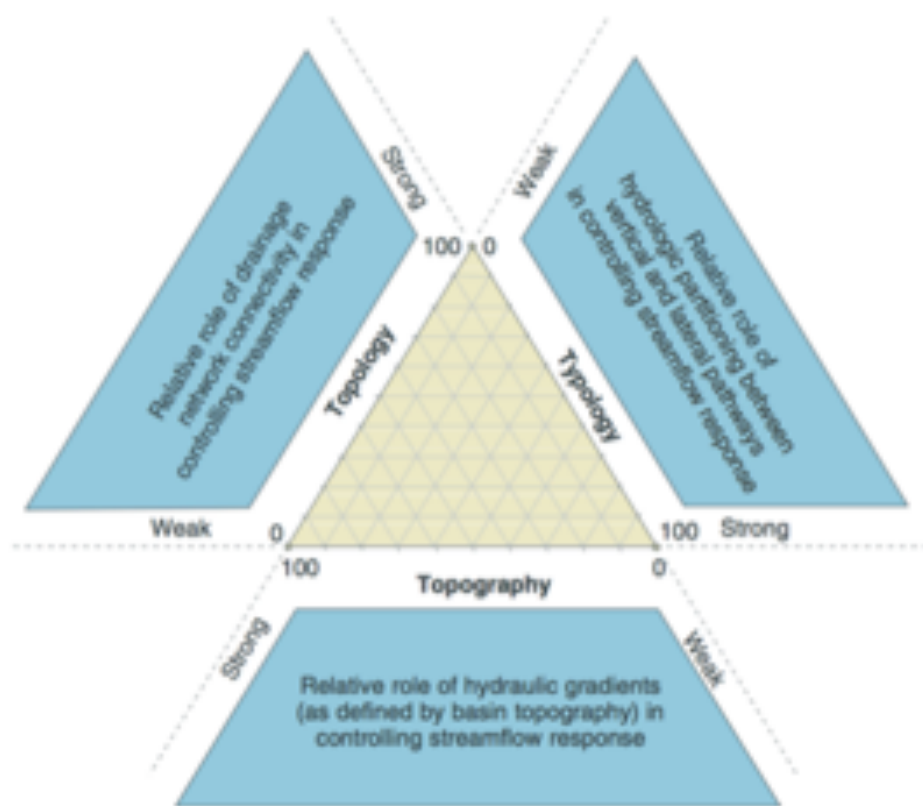
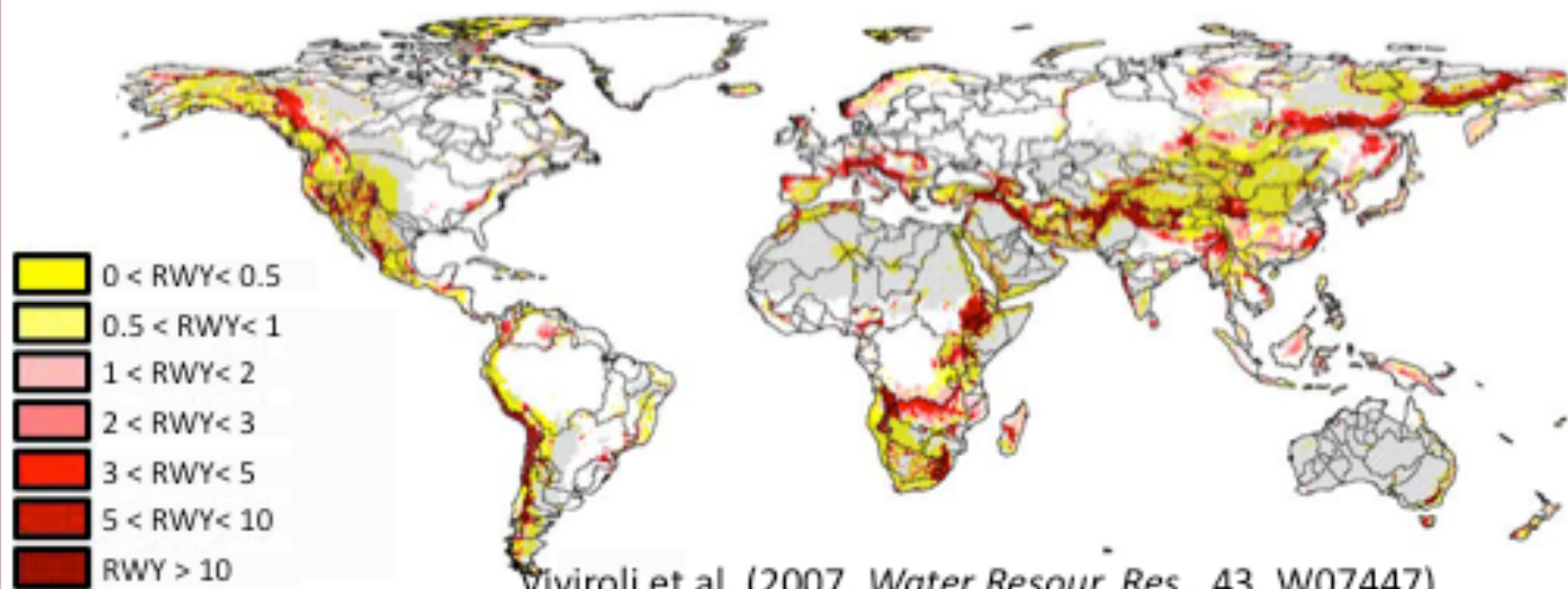


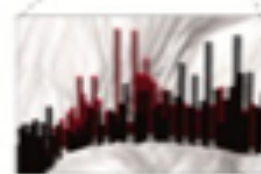
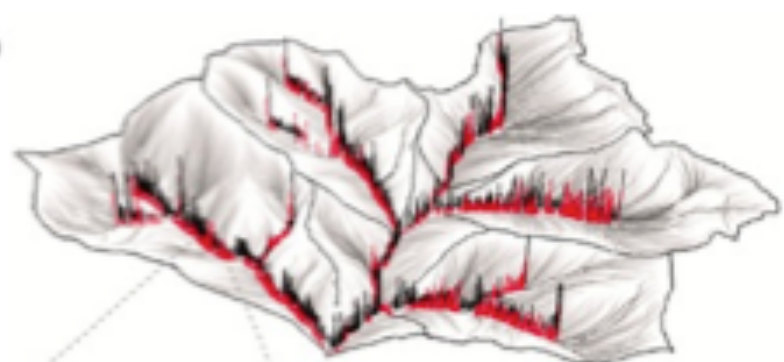
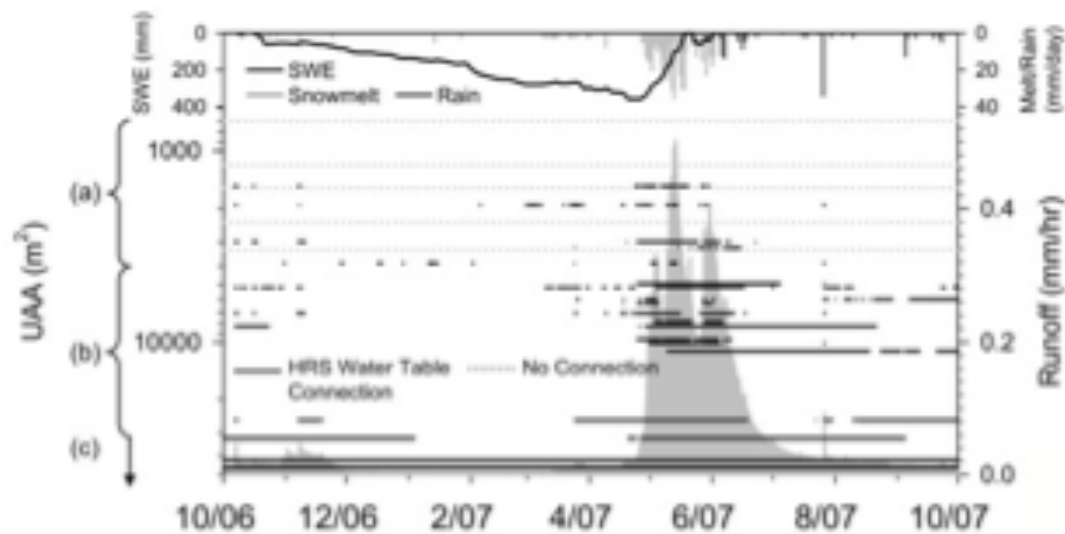
Figure 1. The T³ (typology—topography—topology) template and the relative role of each factor in the regulation of streamflow response



$$\text{Relative water yield (RWY)} = \frac{\text{Mountain runoff (mm/y)}}{\text{Lowland runoff (mm/y)}}$$



Viviroli et al. (2007, *Water Resour. Res.*, 43, W07447)



UAA (m²)
543,689
0
Logarithmically Scaled

HRS Connectivity
100%
0%
% of the water year

Jencso and McGlynn 2009 WRR

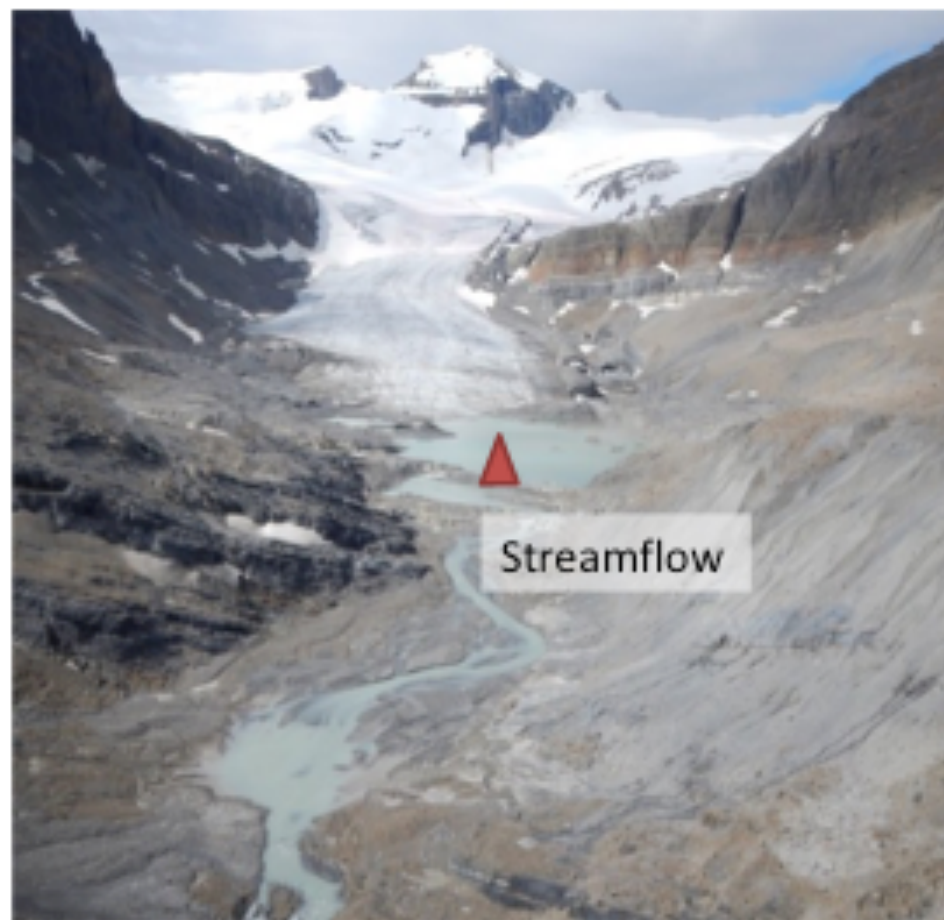
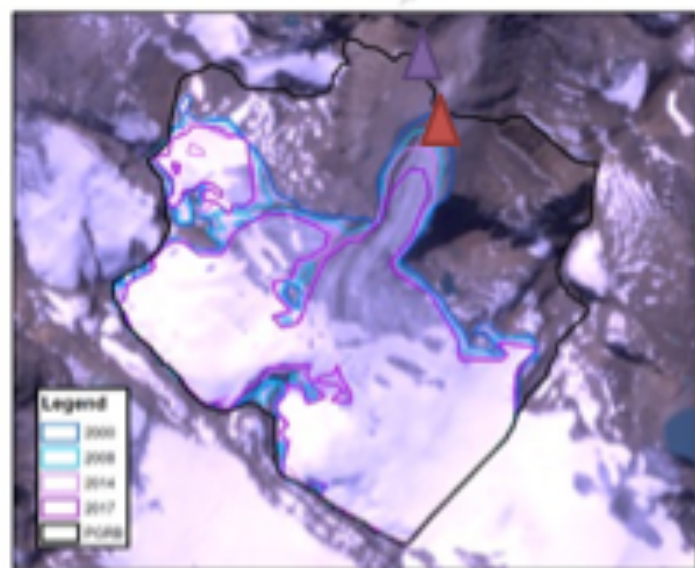


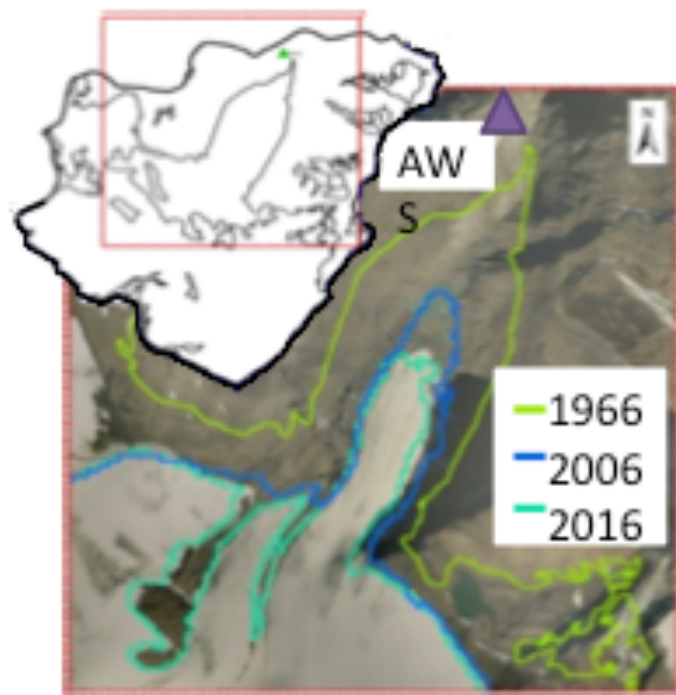
- Large thermal gradients
 - Elevation, aspect
- Poorly developed 'soils'
 - Large porosity, uncertain geological setting
- Frozen ground status
- Glaciers and perennial snowpacks

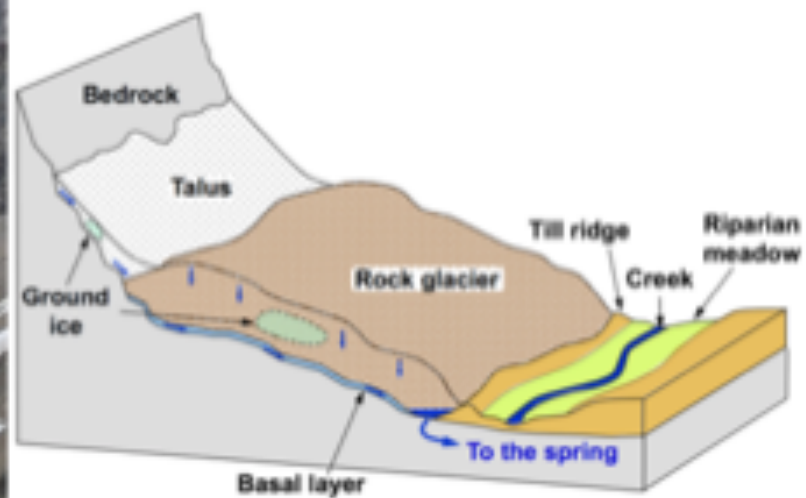


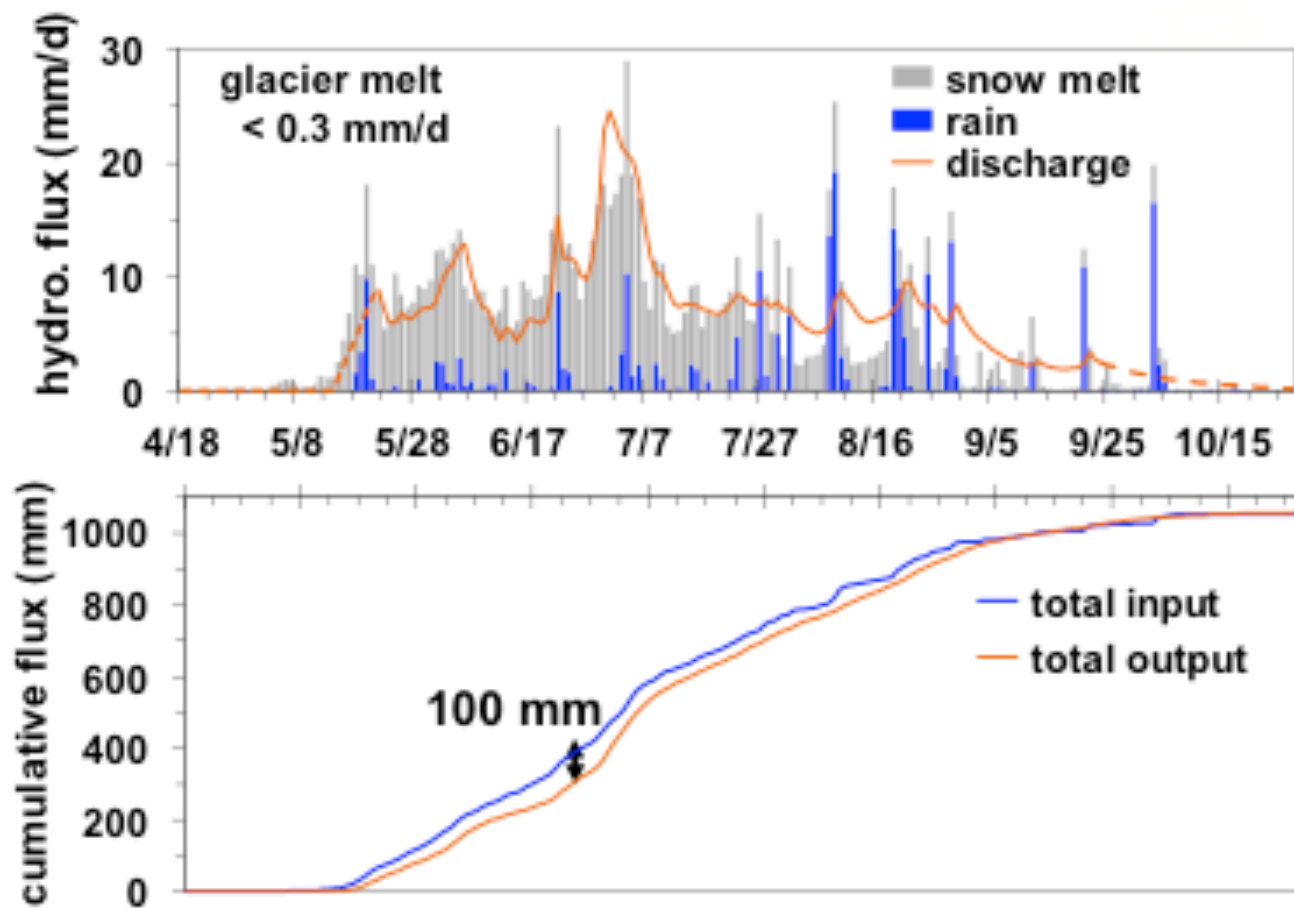


- Glacier and snow are key contributions to alpine streams
- Fairly well characterized, lots of healthy research
 - Focus is on global change

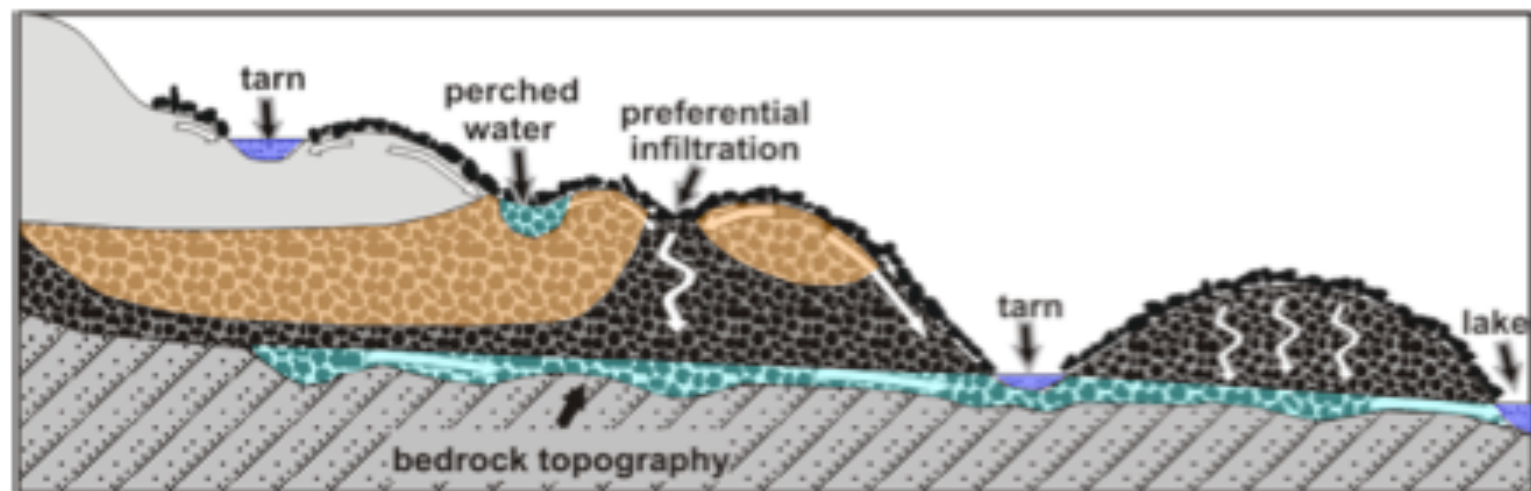


1966: 14.4 km²2016: 9.9 km²





Hood & Hayashi (2015, *J. Hydrol.* 521:482-497)



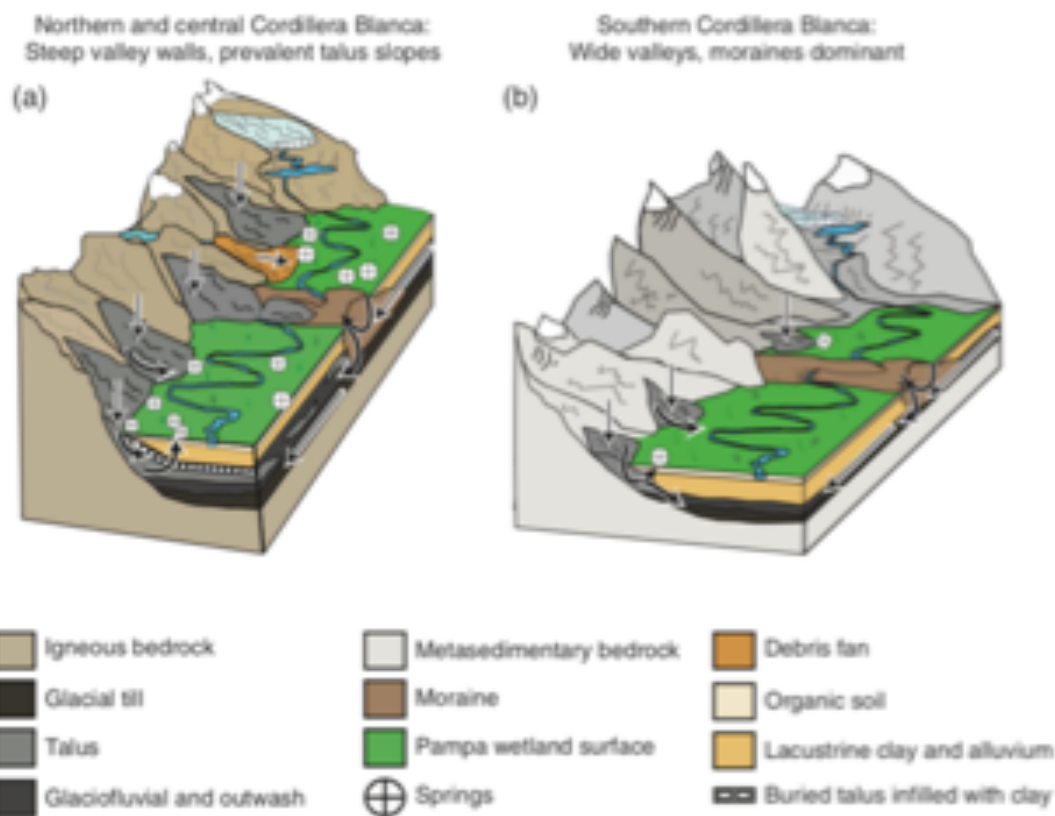
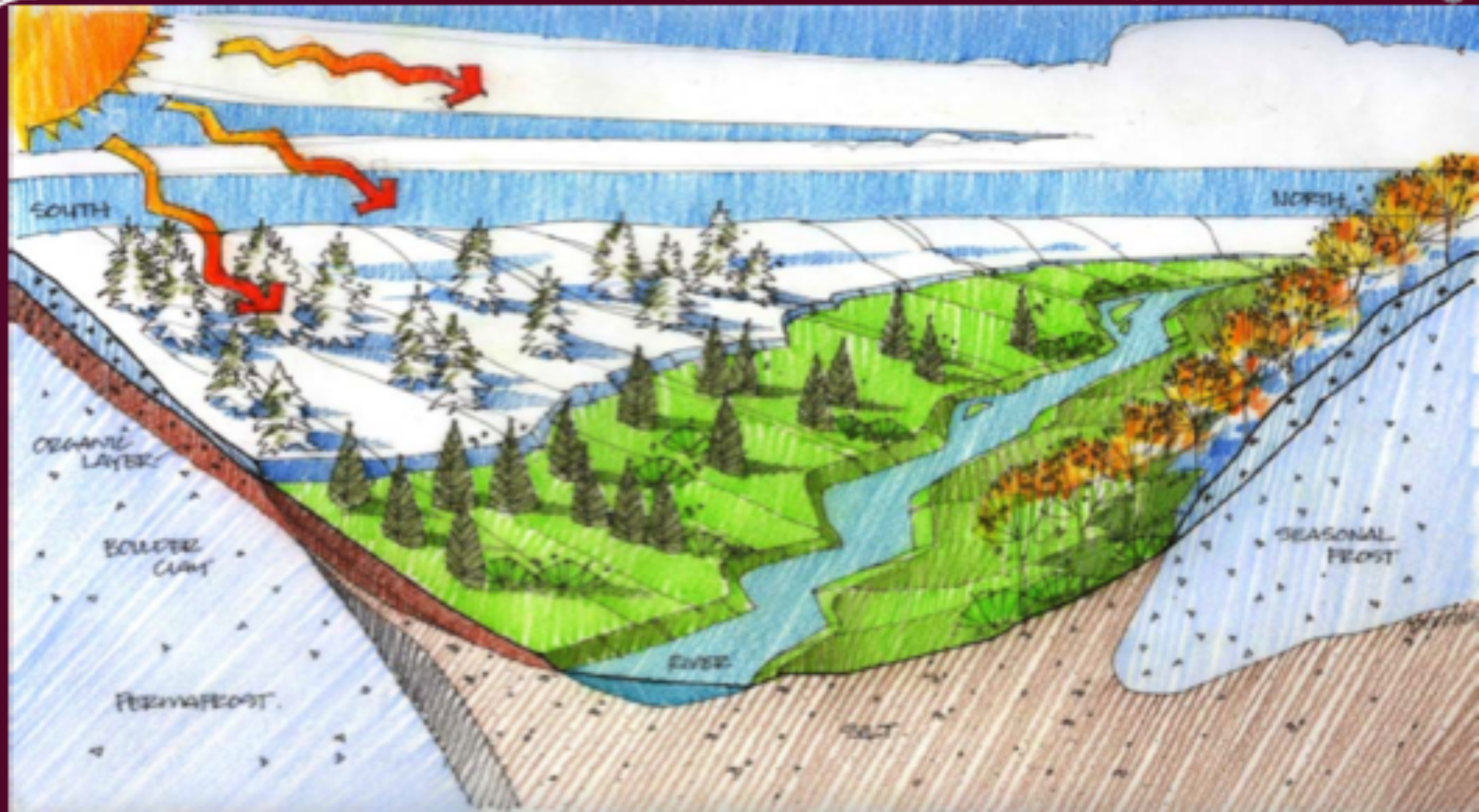


FIGURE 4 (a) Conceptual diagrams of groundwater flow for northern and central valleys (north of Huaraz) and (b) southern valleys (south of Huaraz) of the Cordillera Blanca. White arrows indicate direction of groundwater recharge, flow, and exchange



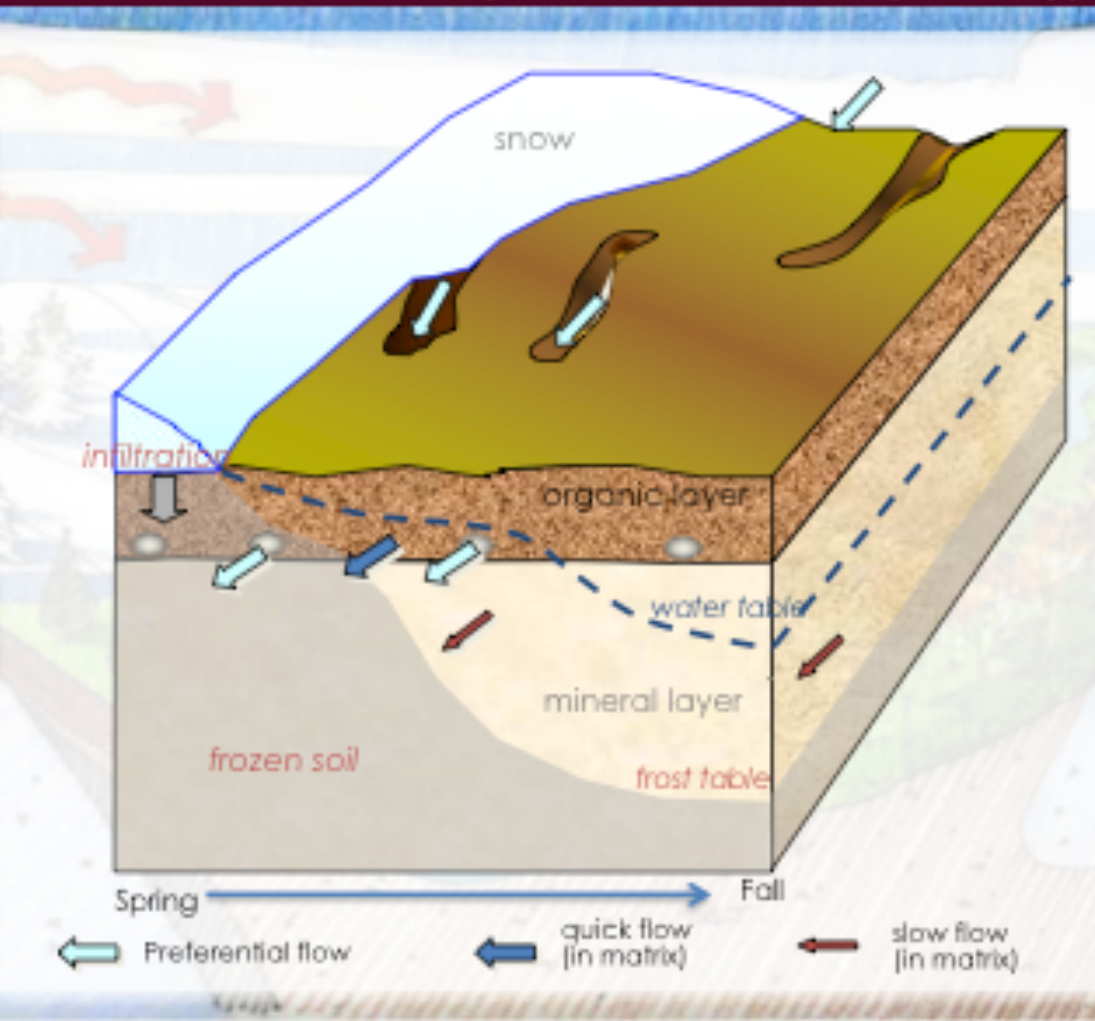


Colder alpine systems



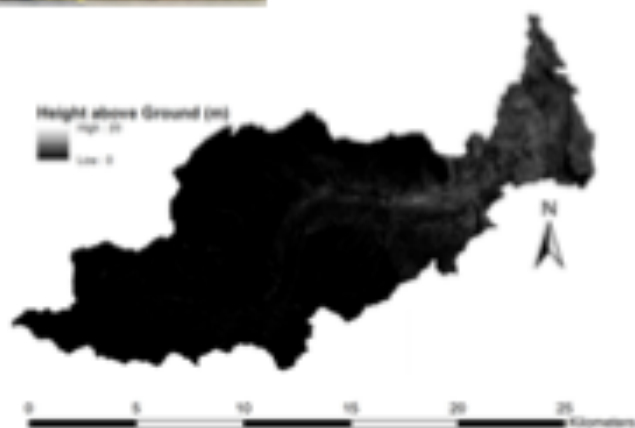
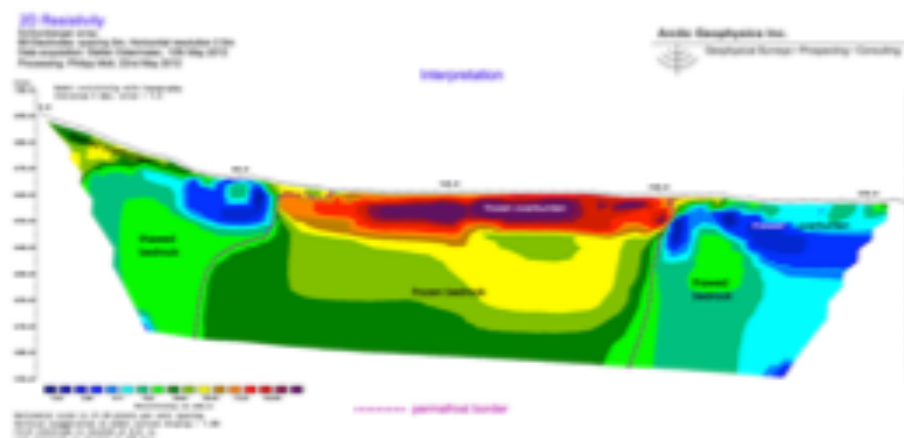


The importance of permafrost

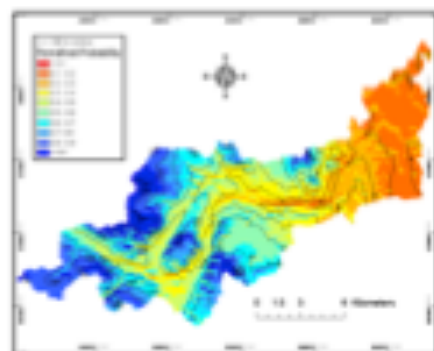
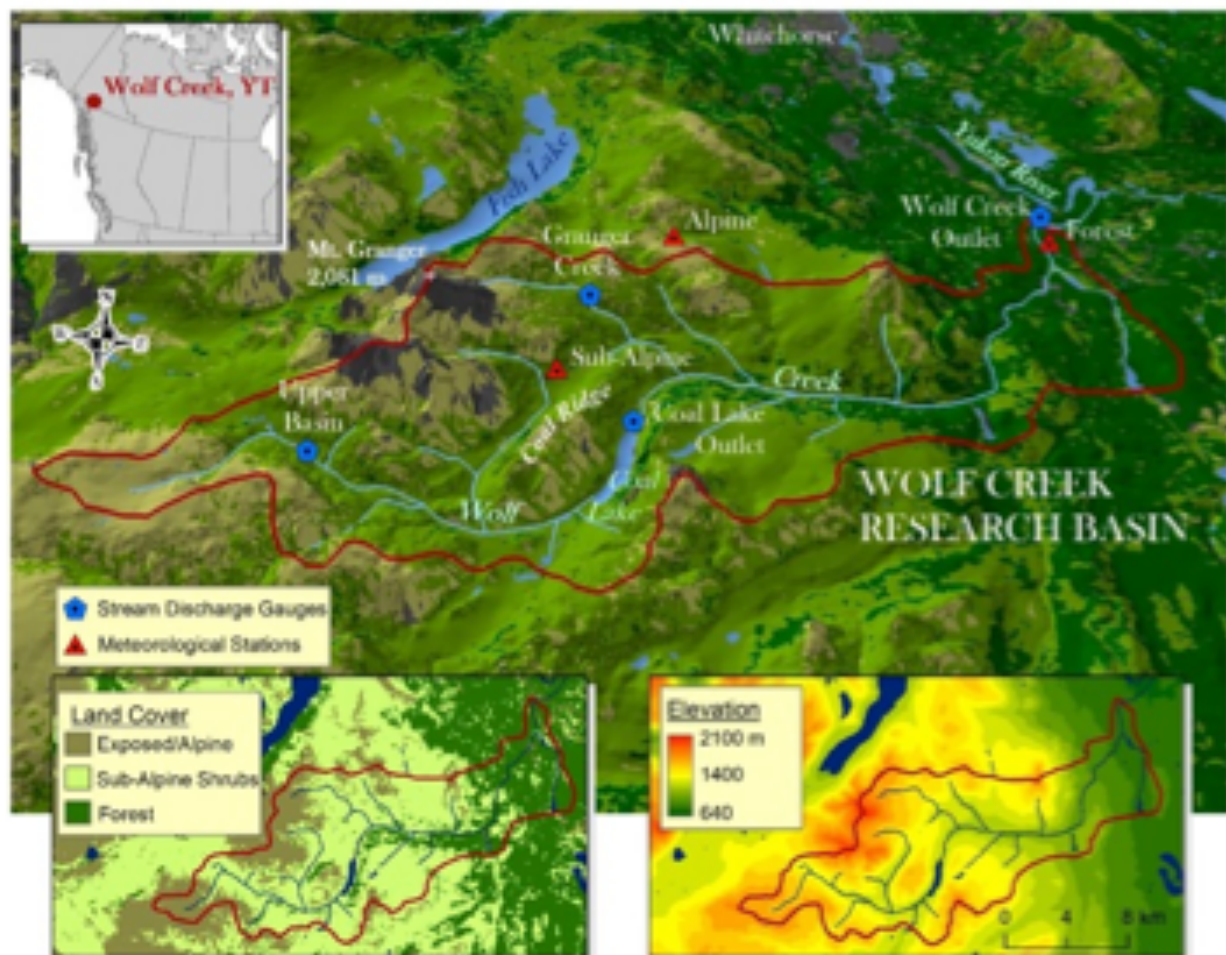




Largely an issue of characterization and representation

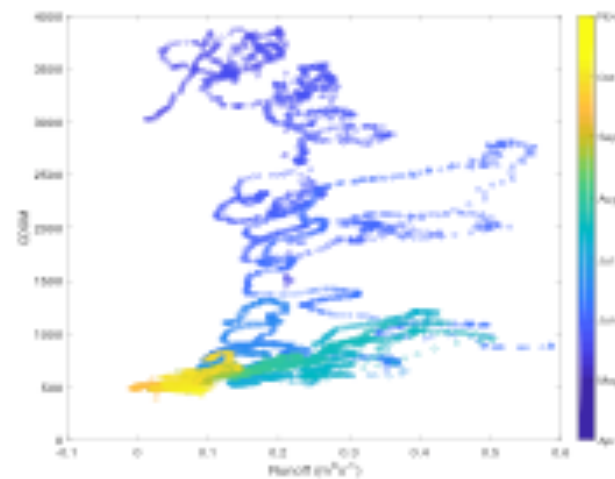
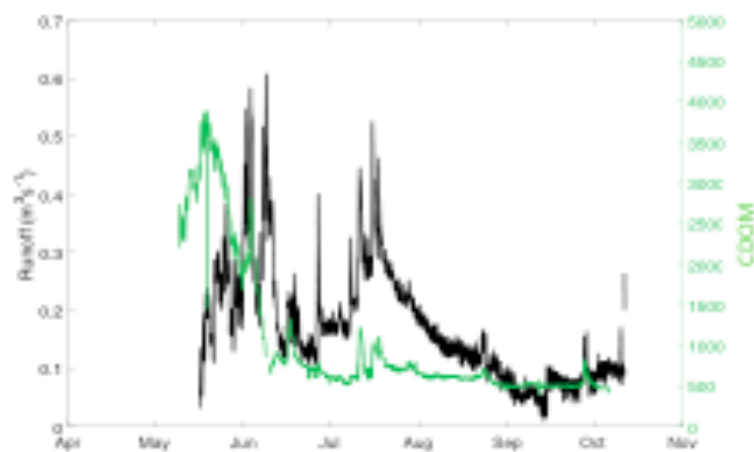
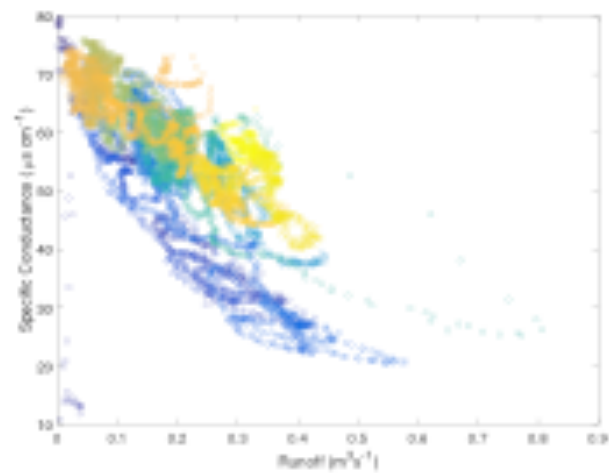
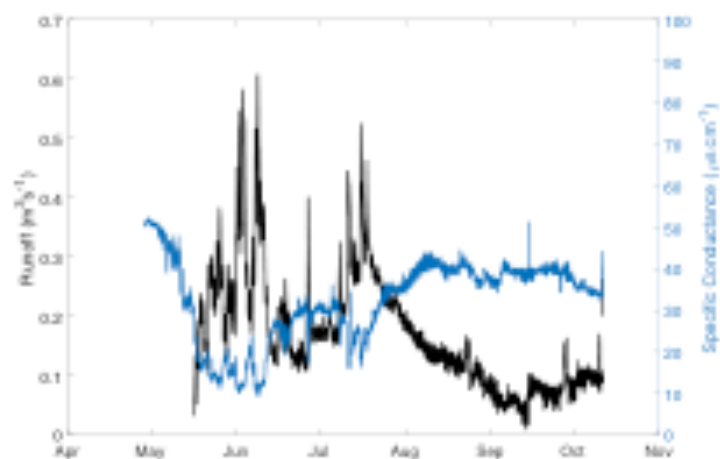


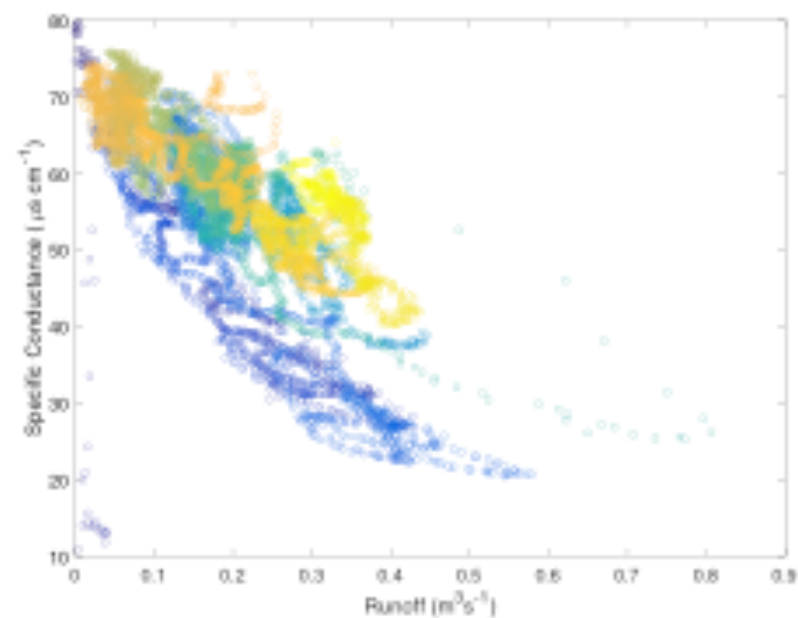




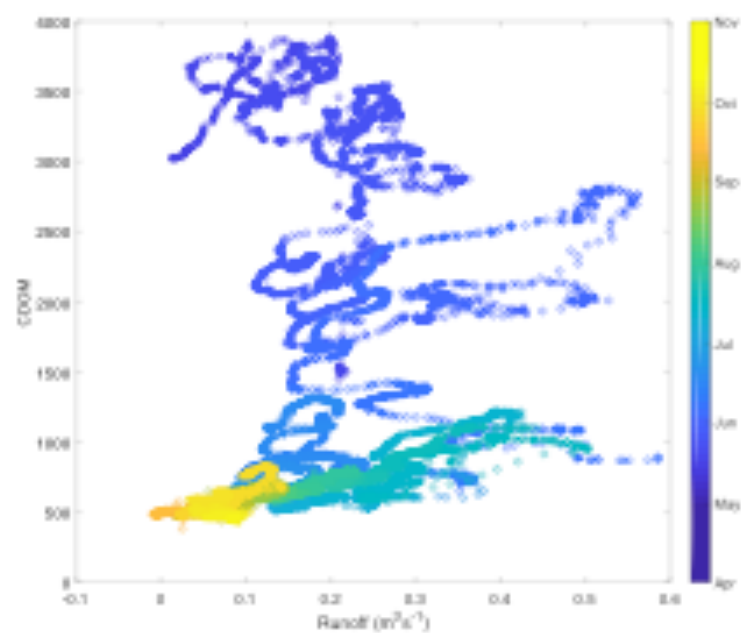
Permafrost probability map
Lewkowicz & Ednie 2004, PPP







- Seasonal trend show gradual shifts in Q-SpC as flow paths decline atop frozen ground and as catchment dries



- General clockwise 'flush'

- Analyze in event responses for Q, SpC and DOC
- How do patterns change, what does hysteresis tell us about runoff processes and overall watershed connections

Hydrological
Processes

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SPECIAL ISSUE CANADIAN GEOPHYSICAL UNION 2017

HydRun: A MATLAB toolbox for rainfall-runoff analysis

Weigang Tang , Sean K. Carey

Accepted manuscript online: 27 March 2017 Full publication history

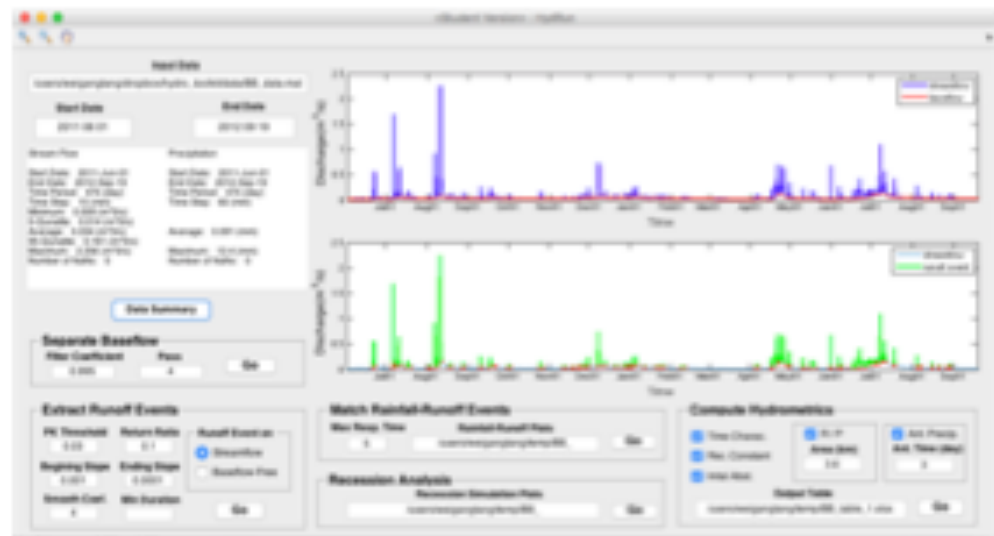
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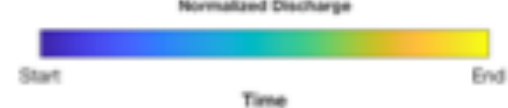
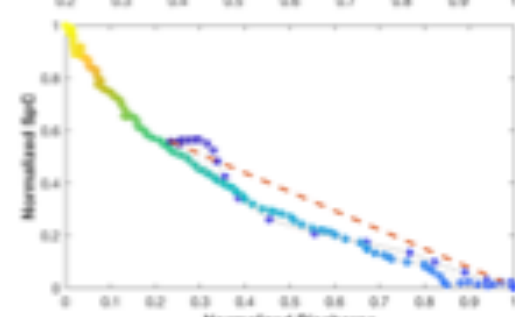
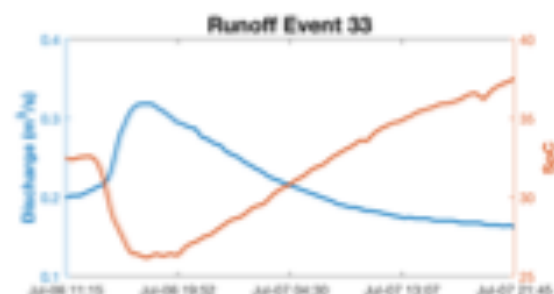
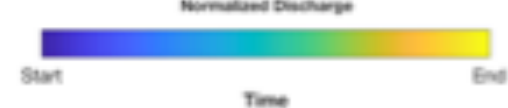
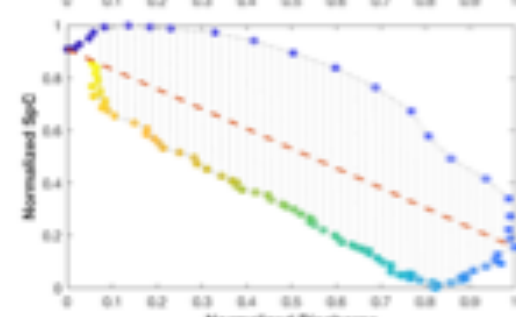
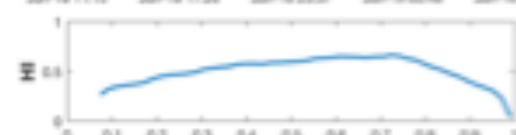
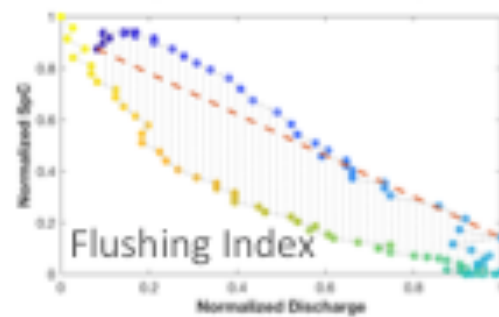
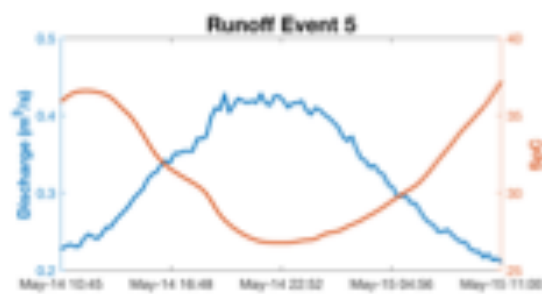
638 C. E. M. Lloyd et al. Testing an integral index for analysing storm discharge-concentration systems

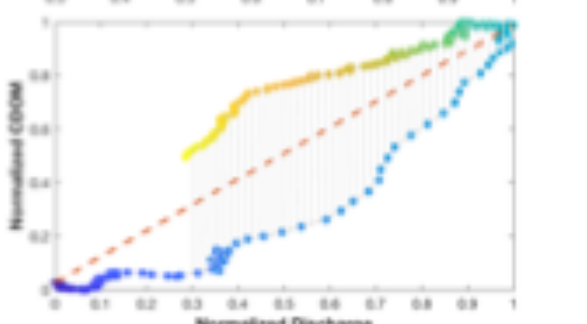
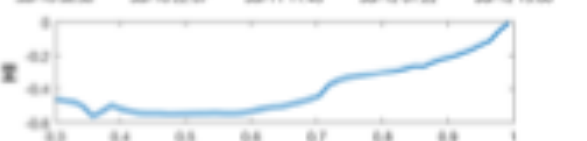
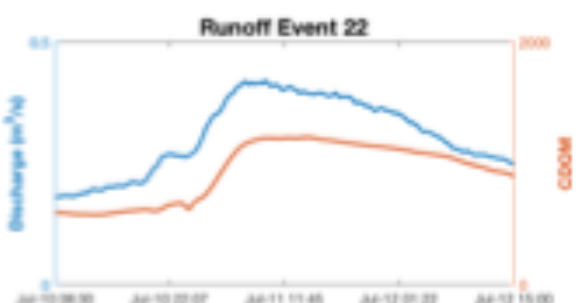
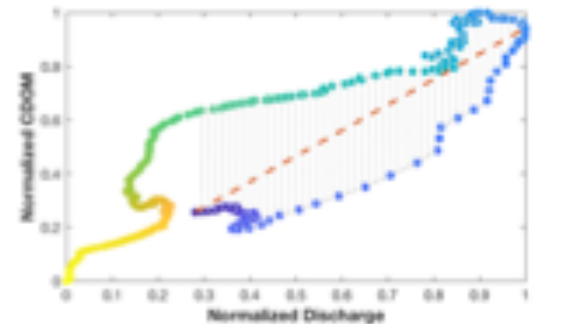
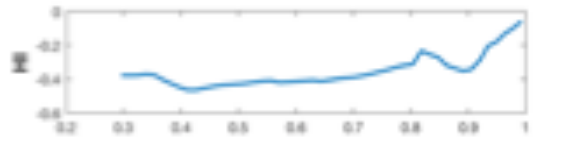
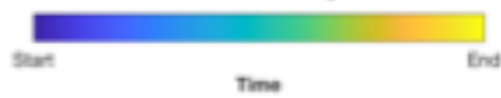
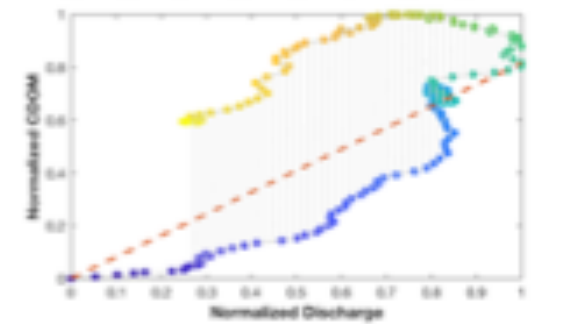


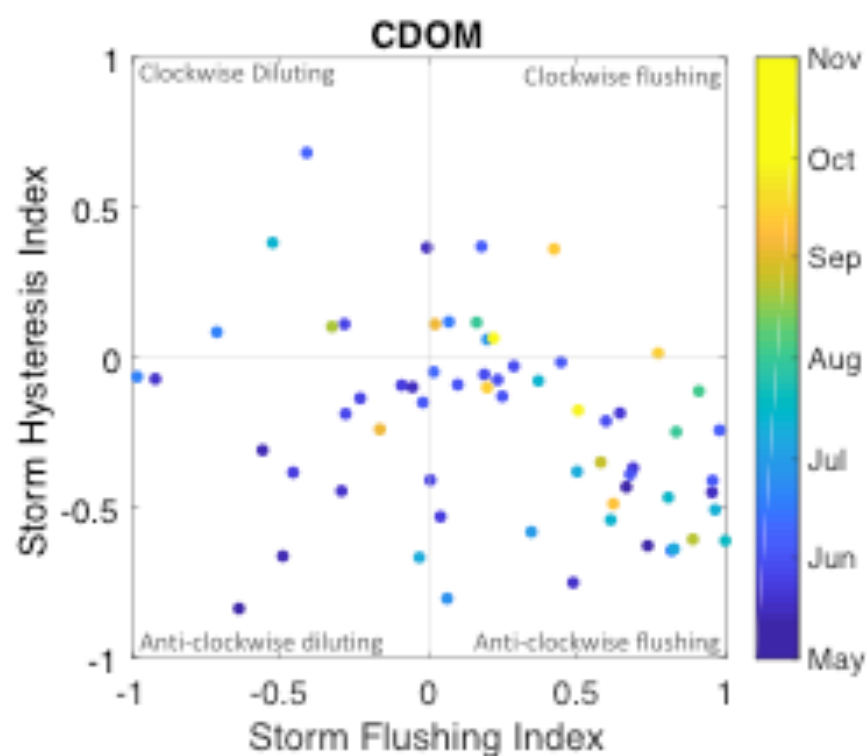
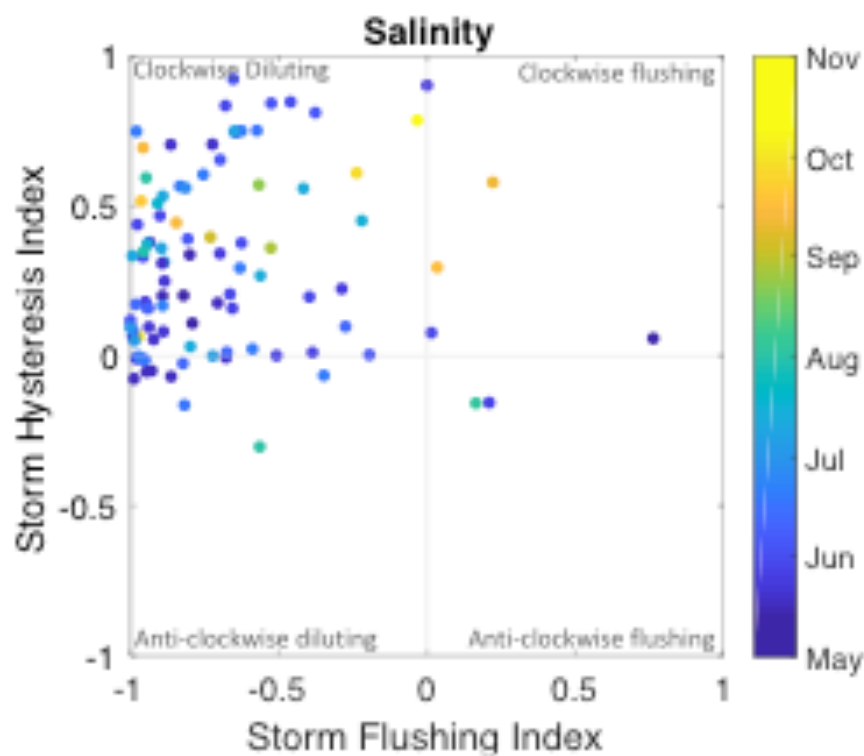
Figure 3. (a) Diagram of storm initial concentration. (b) Storm initial discharge on the basis of the calculated IC when the relationship between discharge and time data is used and (c) an idealized and conceptual storm. (Distinguishing the regions of increasing different quantities of time on the IC calculated, where IC_{10} and IC_{50} are the original and subsequent units in IC (DHR) methods, respectively and IC_{10} is the proposed method. (Colour-figure available online at [www.int-j-eas.com](#).)

Lloyd, C. et al. 2016. Hydrology and Earth System Sciences, 20, 625–632.



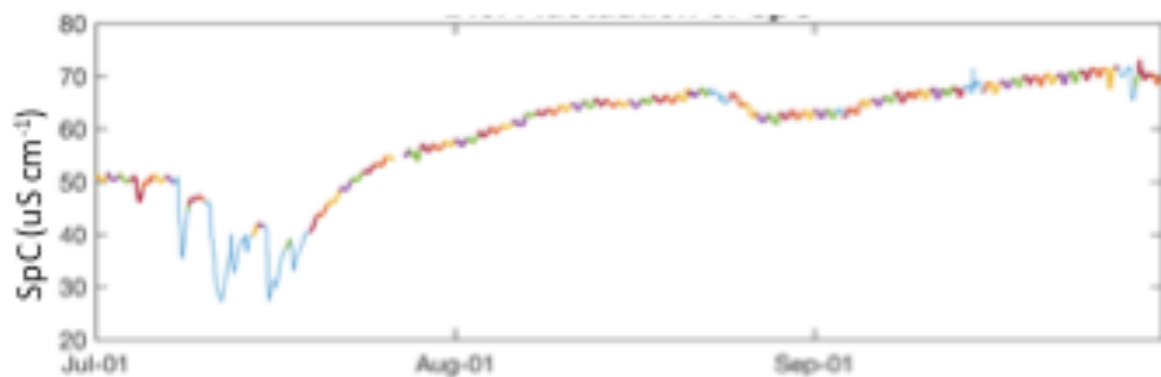
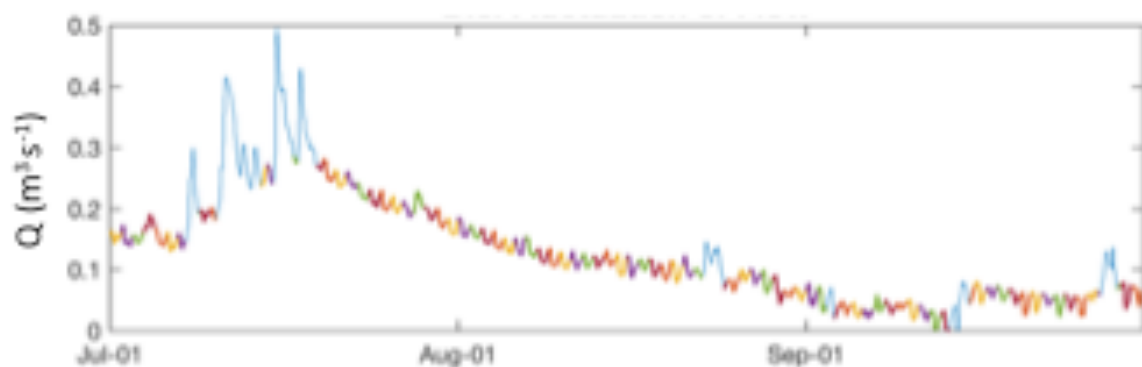






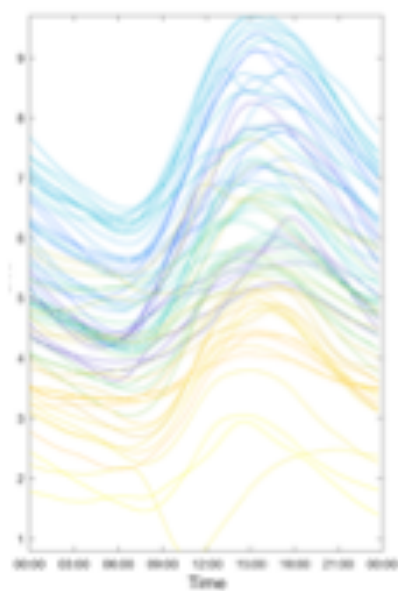


➤ Q and SpC from July 1 to Sept 30

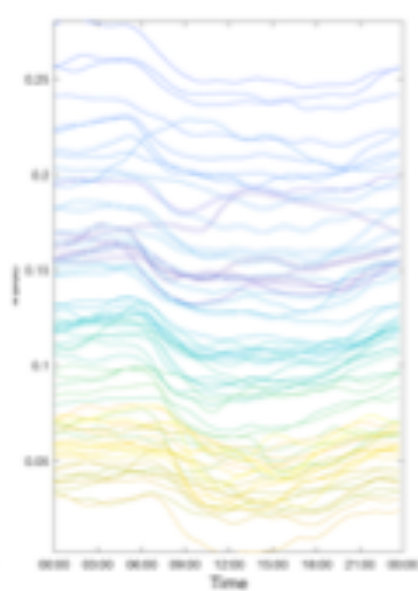




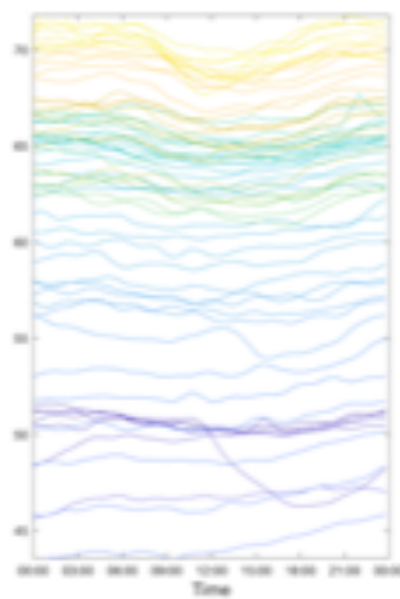
Water Temp



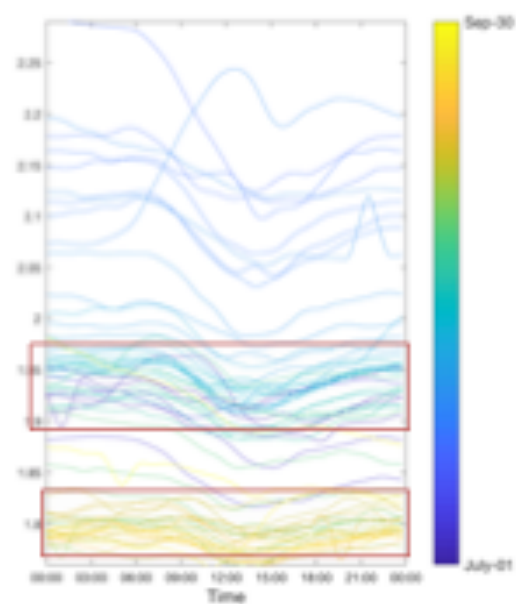
Discharge (Q)



SpC

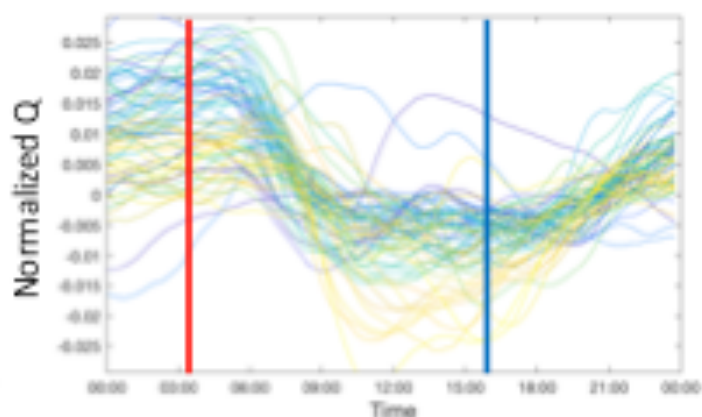
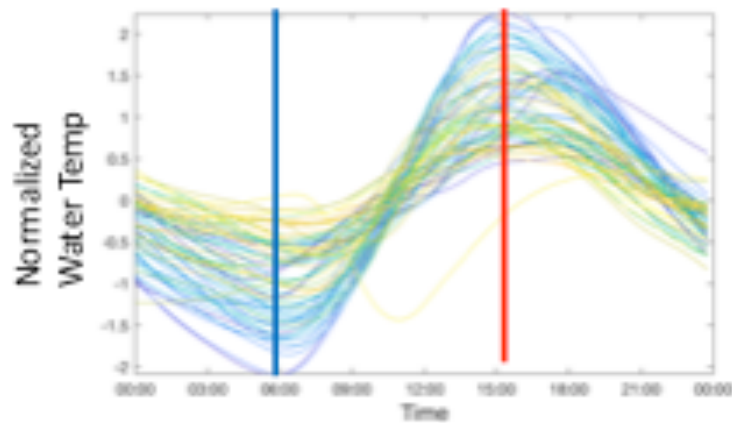


DOC

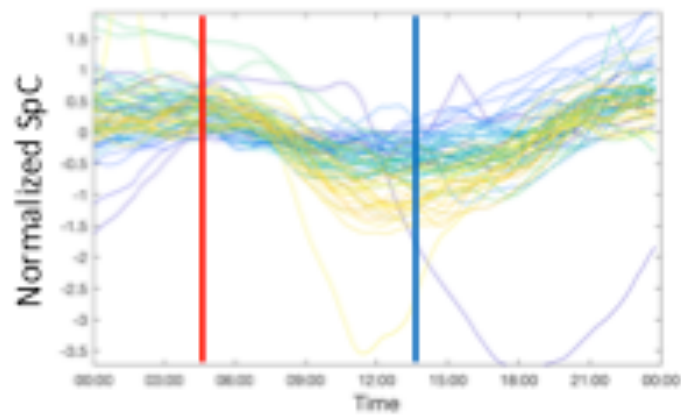
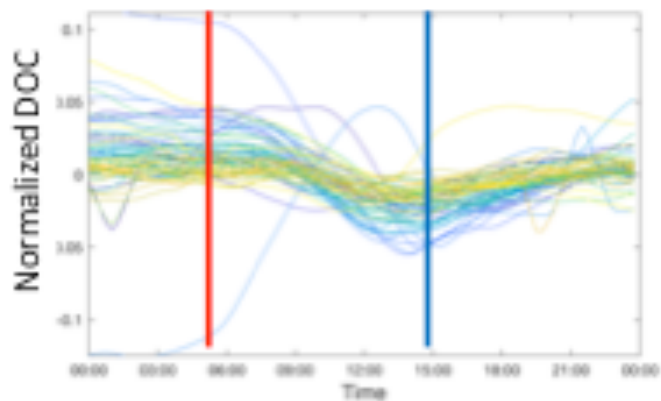
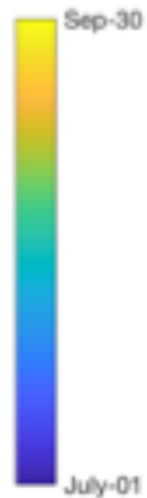


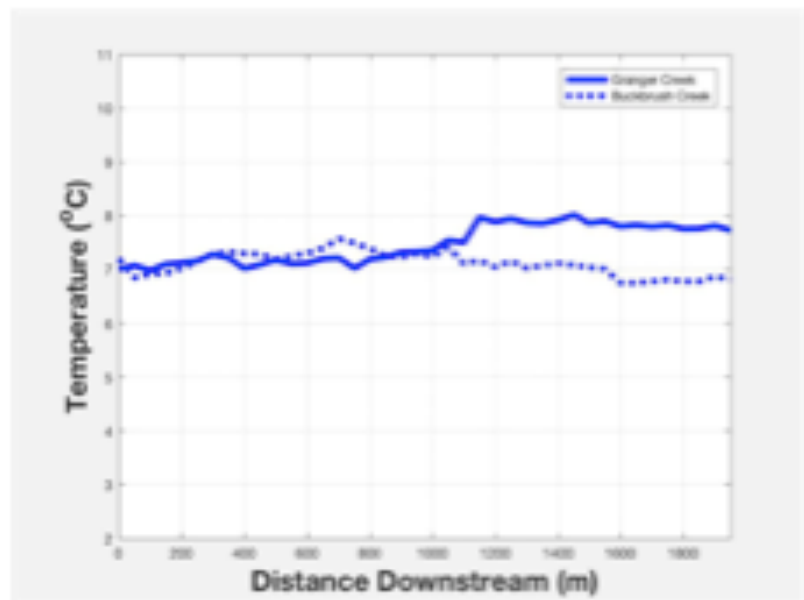
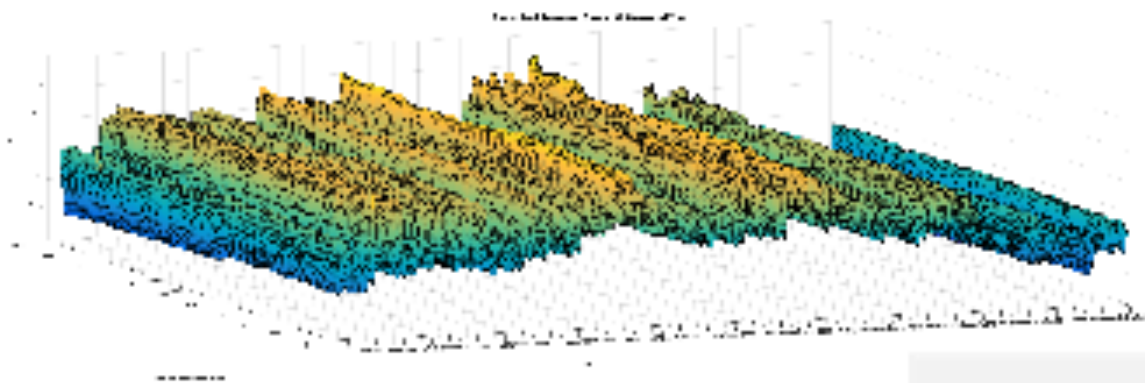
Sep-30

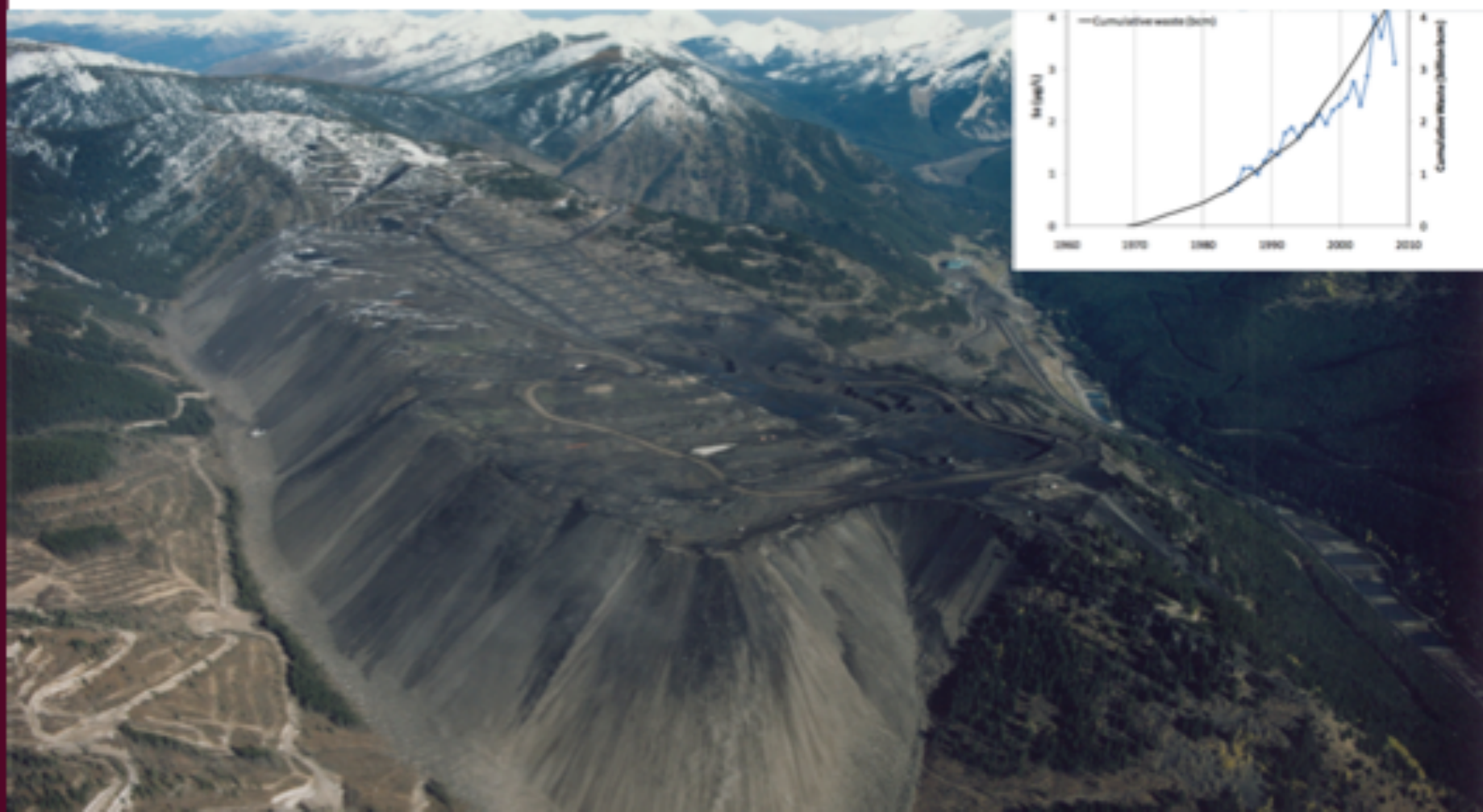
July-01

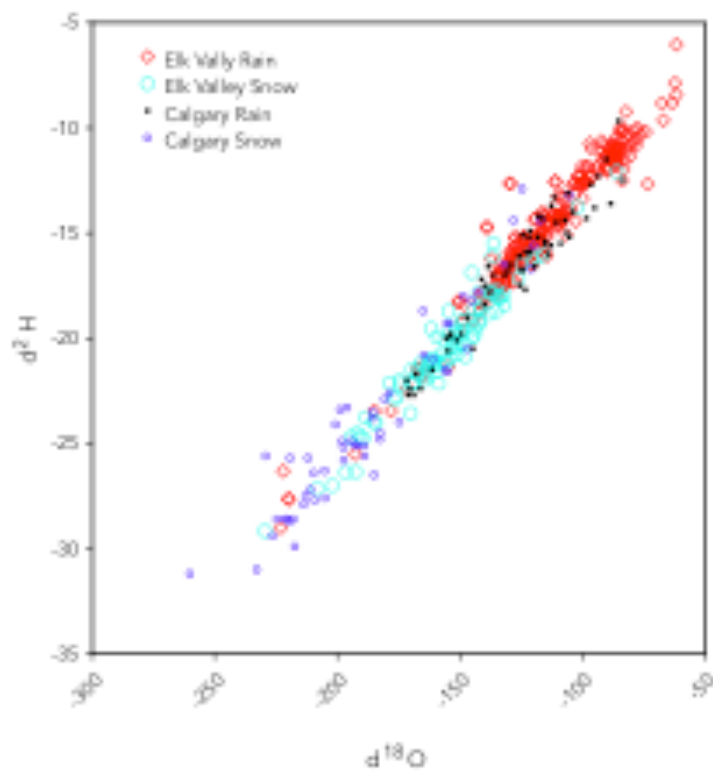


➤ Water temp out of phase means viscosity is not a driver

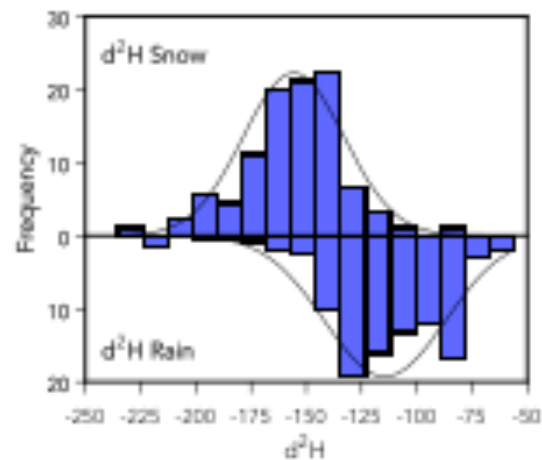
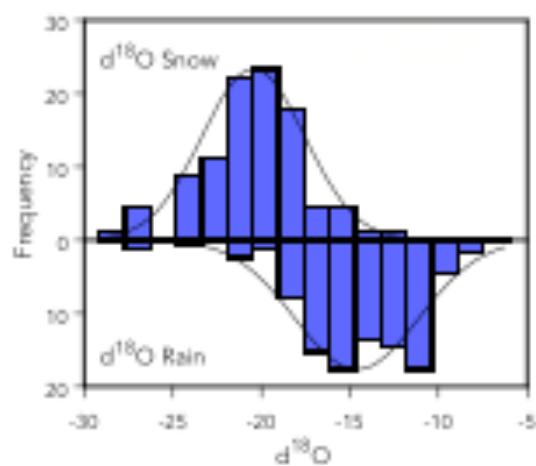


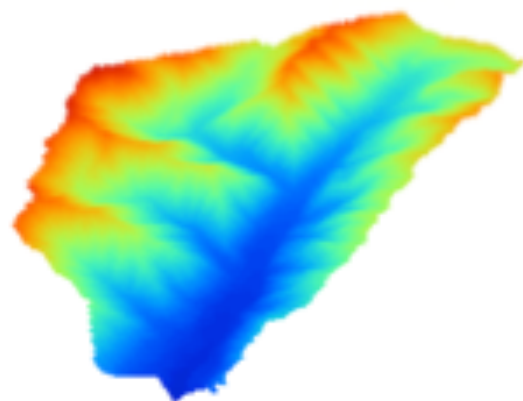
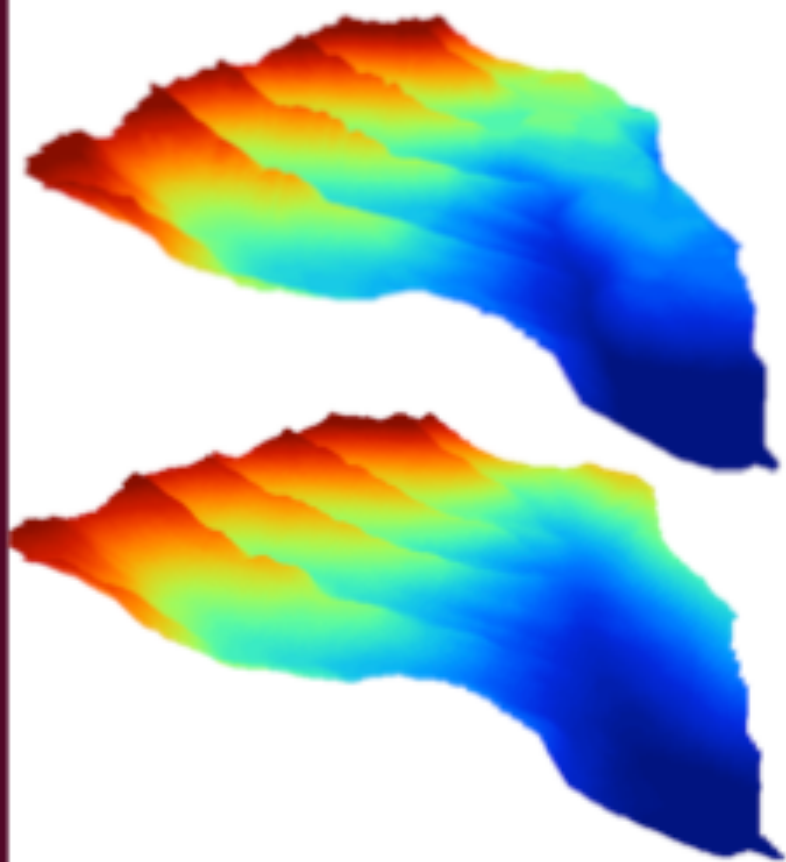




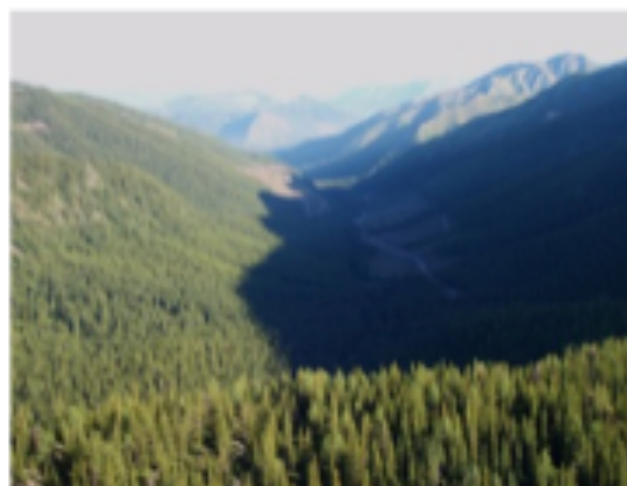


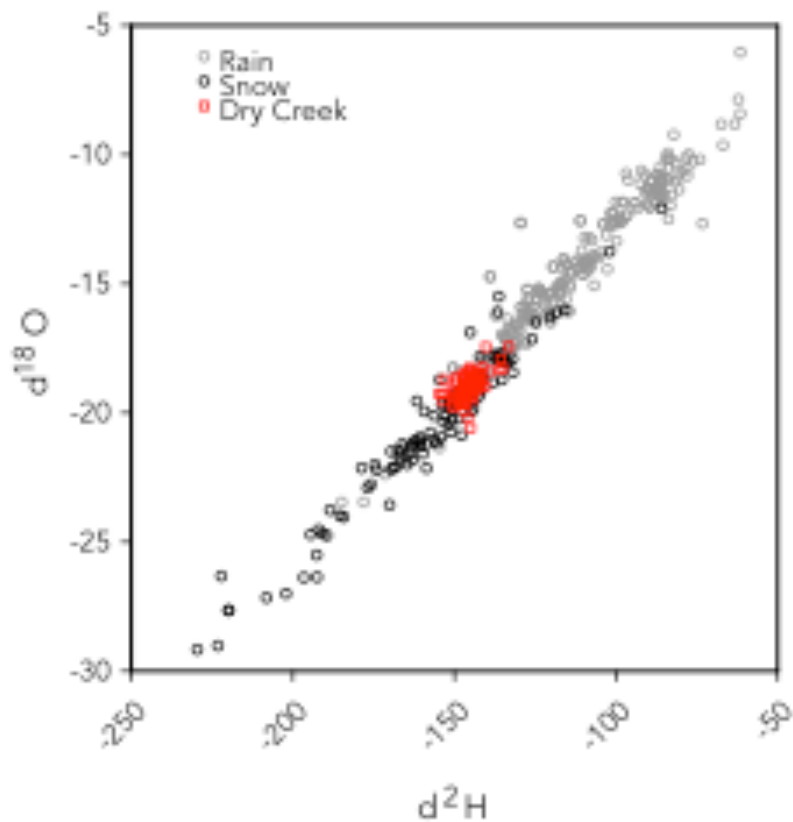
- ⇒ LMWL similar to Calgary
- ⇒ Distribution of snow and rain



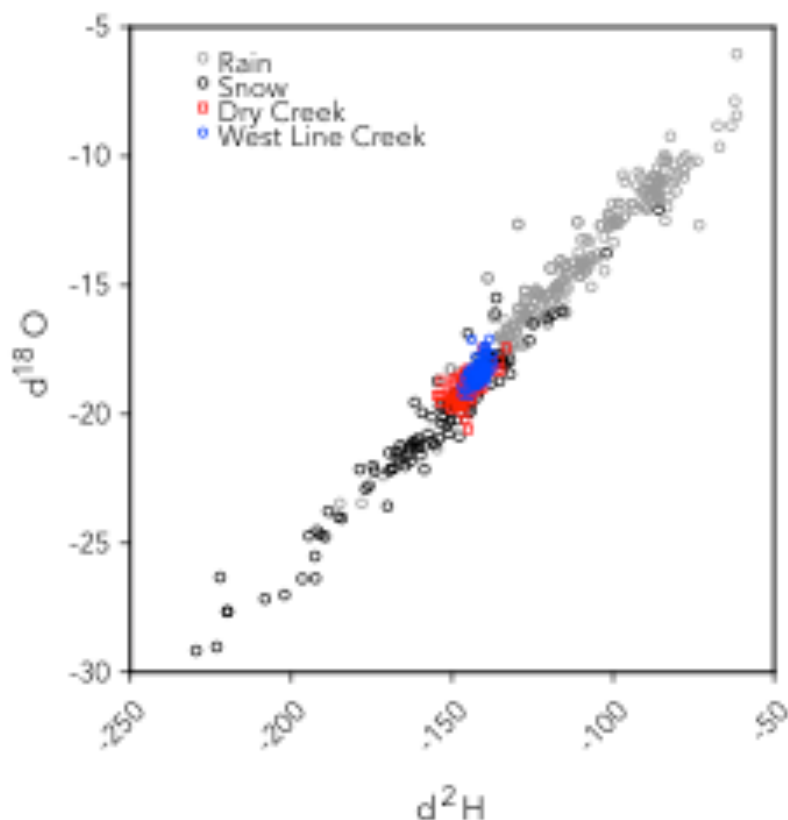


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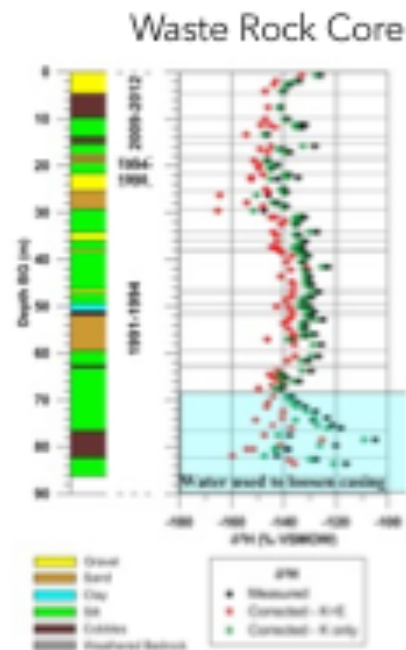
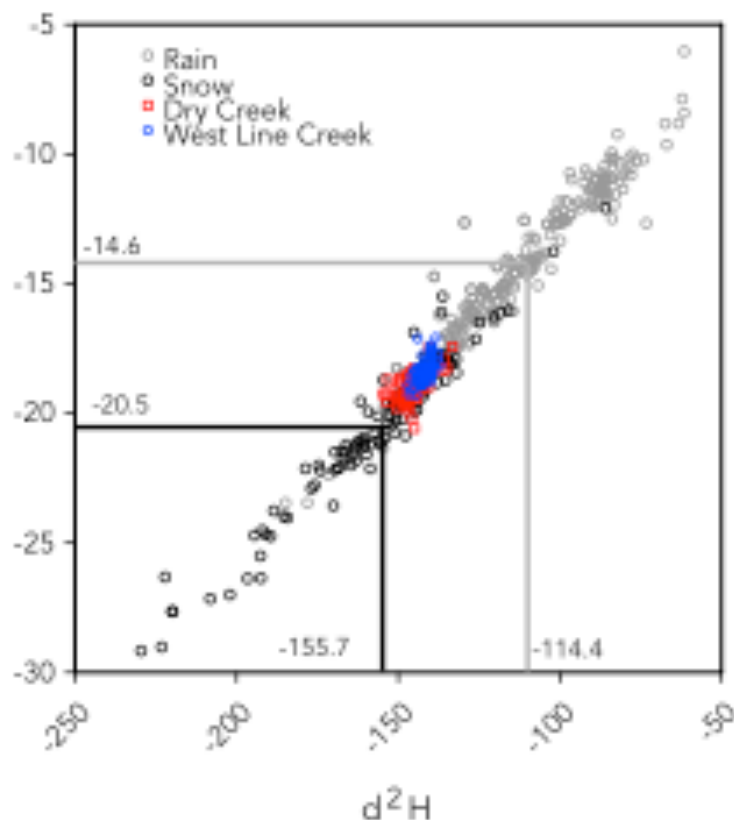
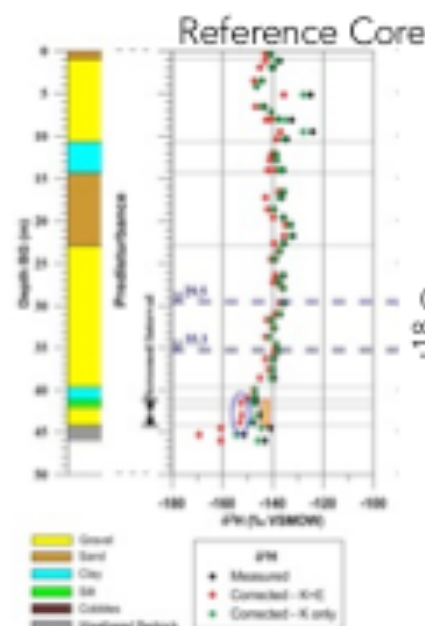




⇒ Dry Creek (reference) – little variation seasonally in streamflow signal

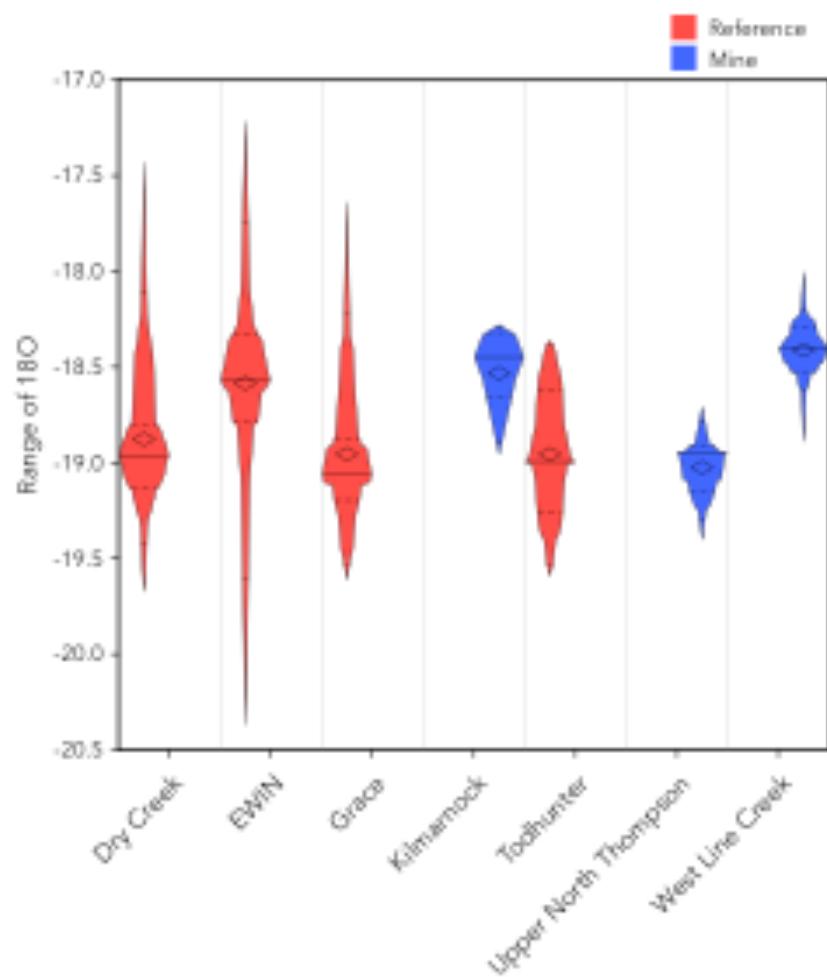


- ⇒ Dry Creek (reference) – little variation seasonally in streamflow signal
- ⇒ West Line Creek (mine) – signal even more dampened, heavier



Core figures provided by Lee Barbour,
U of Saskatchewan

- ⇒ Dry Creek (reference) – little variation seasonally in streamflow signal
- ⇒ West Line Creek (mine) – signal even more dampened, heavier
- ⇒ Signal more representative of snow, particularly at Dry Creek



⇒ Box percentiles of $\delta^{18}\text{O}$ from reference and mine-influenced watersheds

⇒ Greater distribution of isotopes commonly inferred as a proxy for decreased transit times and reduced storage



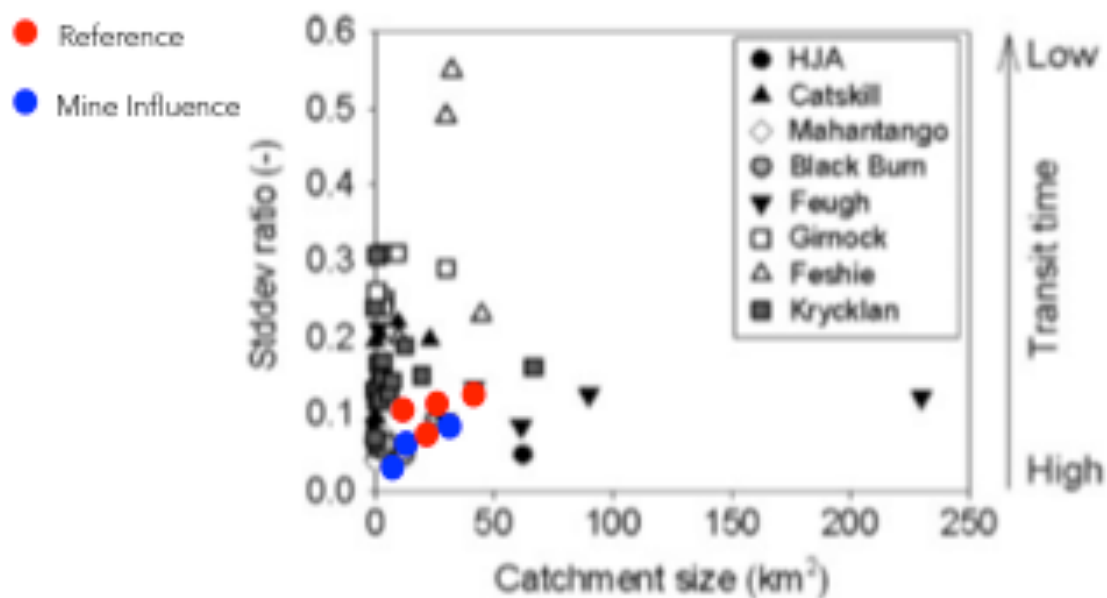


Figure 5. Relationship between catchment size and the inverse transit time proxy ITTP (ratio of standard deviations of $\delta^{18}\text{O}$ measurements of stream water to precipitation) for the eight catchments and their associated sub-catchments

Comparison of transit times with literature values from alpine watersheds
Tetzlaff et al., (2009). *Hydrological Processes*, 23, 945-953

Overview

Transit times—the link between hydrology and water quality at the catchment scale



Markus Hrachowitz,^{1*} Paolo Benettin,² Boris M. van Breukelen,¹ Ophelie Fovet,³ Nicholas J.K. Howden,⁴ Laurent Ruiz,³ Ype van der Velde³ and Andrew J. Wade⁶

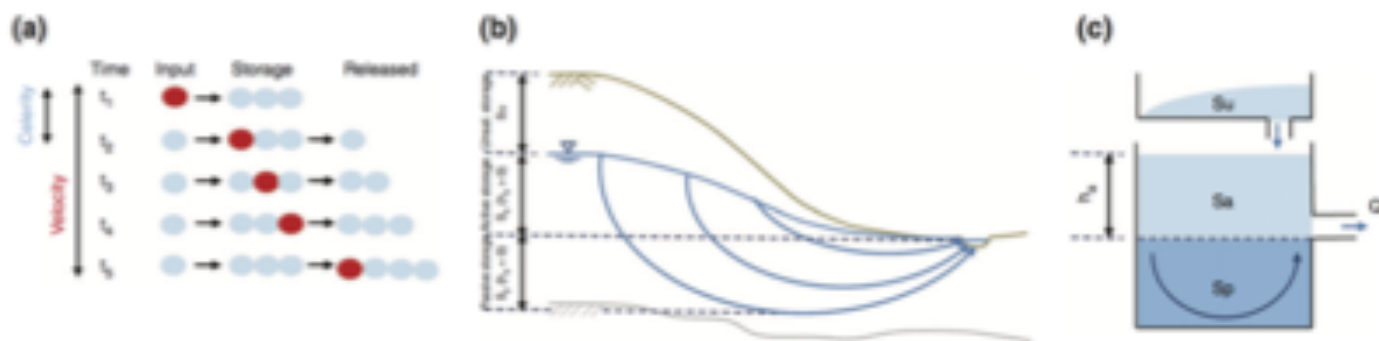
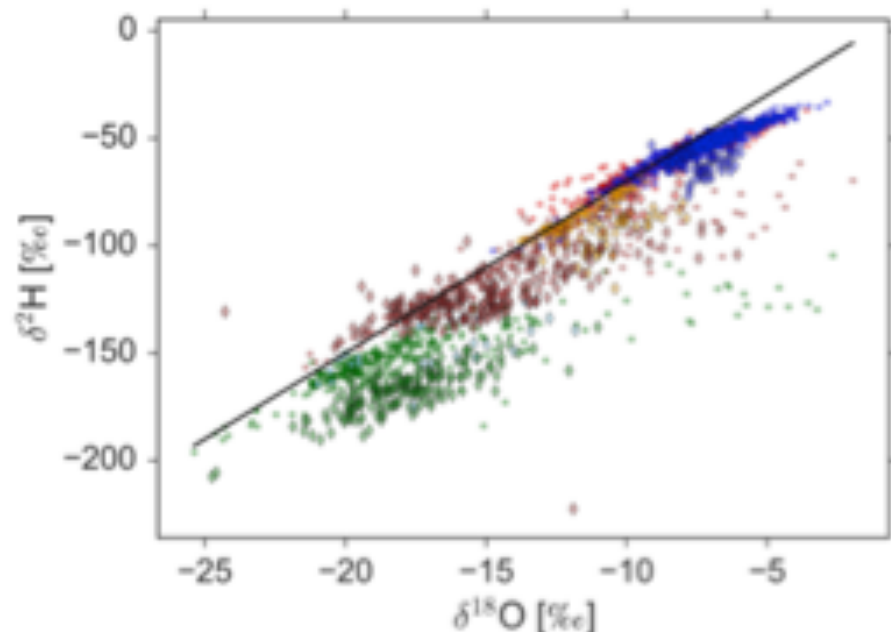
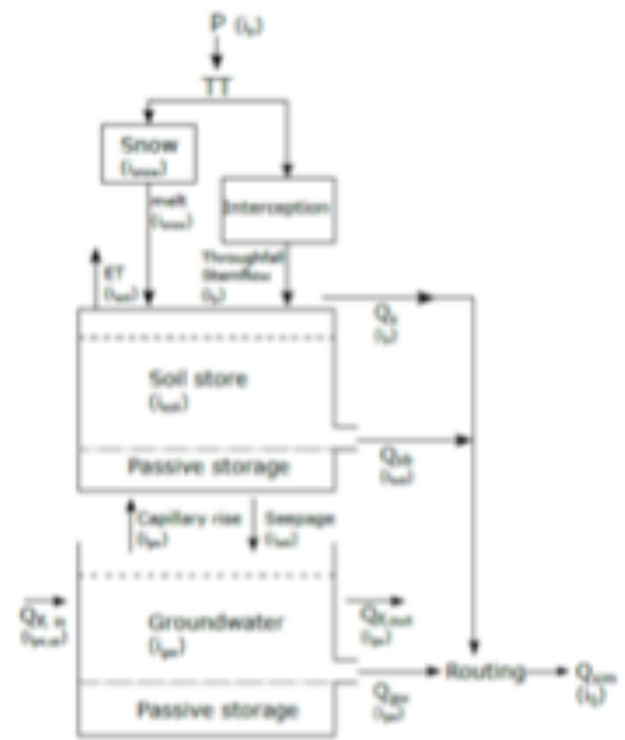


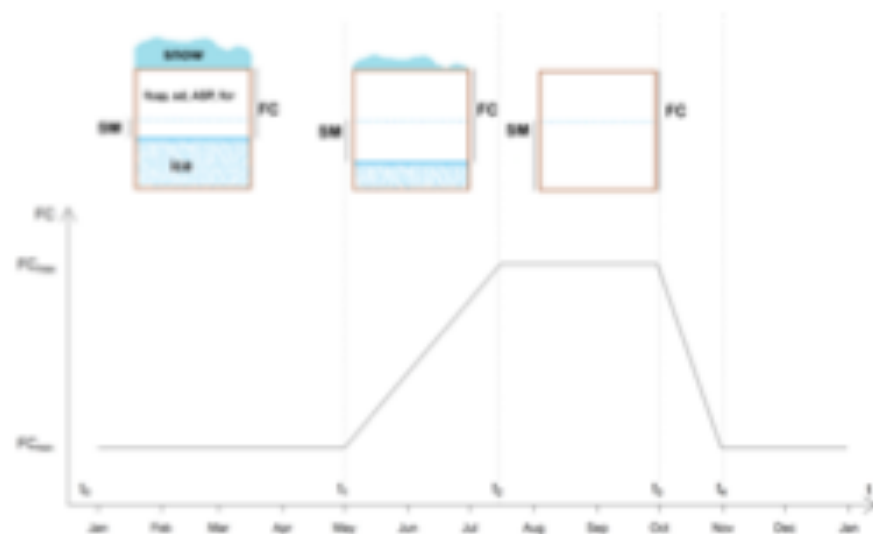
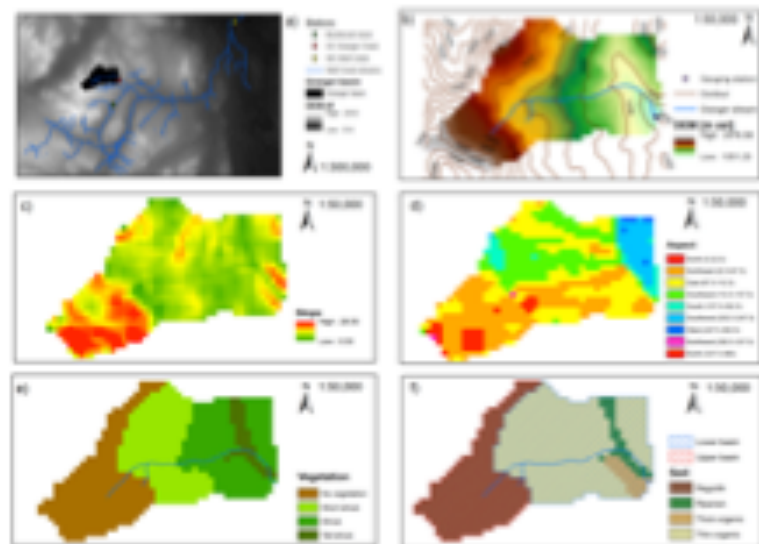
FIGURE 2 | (a) Conceptualization of the difference between celerity-driven hydrological response and velocity-driven transport processes using the analogy of a game of billiards. A new input at t_1 (red ball) causes a disturbance of the system that propagates with a celerity and that generates a response (blue ball) at t_2 . The red ball itself, however, is released from the system only at t_3 as it travels at a velocity that is much smaller than the celerity. (b) For a groundwater-dominated system, the propagation of the pressure wave to the stream is controlled by the wave celerity and the active storage S_a (i.e., the pressure head h_a) while the movement of the actual particles is controlled by the flow velocity and the length of the flow trajectory through a hydrologically passive storage volume S_p (after Ref 47), which (c) can be conceptualized in a model with a mixing volume below a given storage threshold. S_u represents the unsaturated zone whose nonlinear behavior is indicated by the curved line.



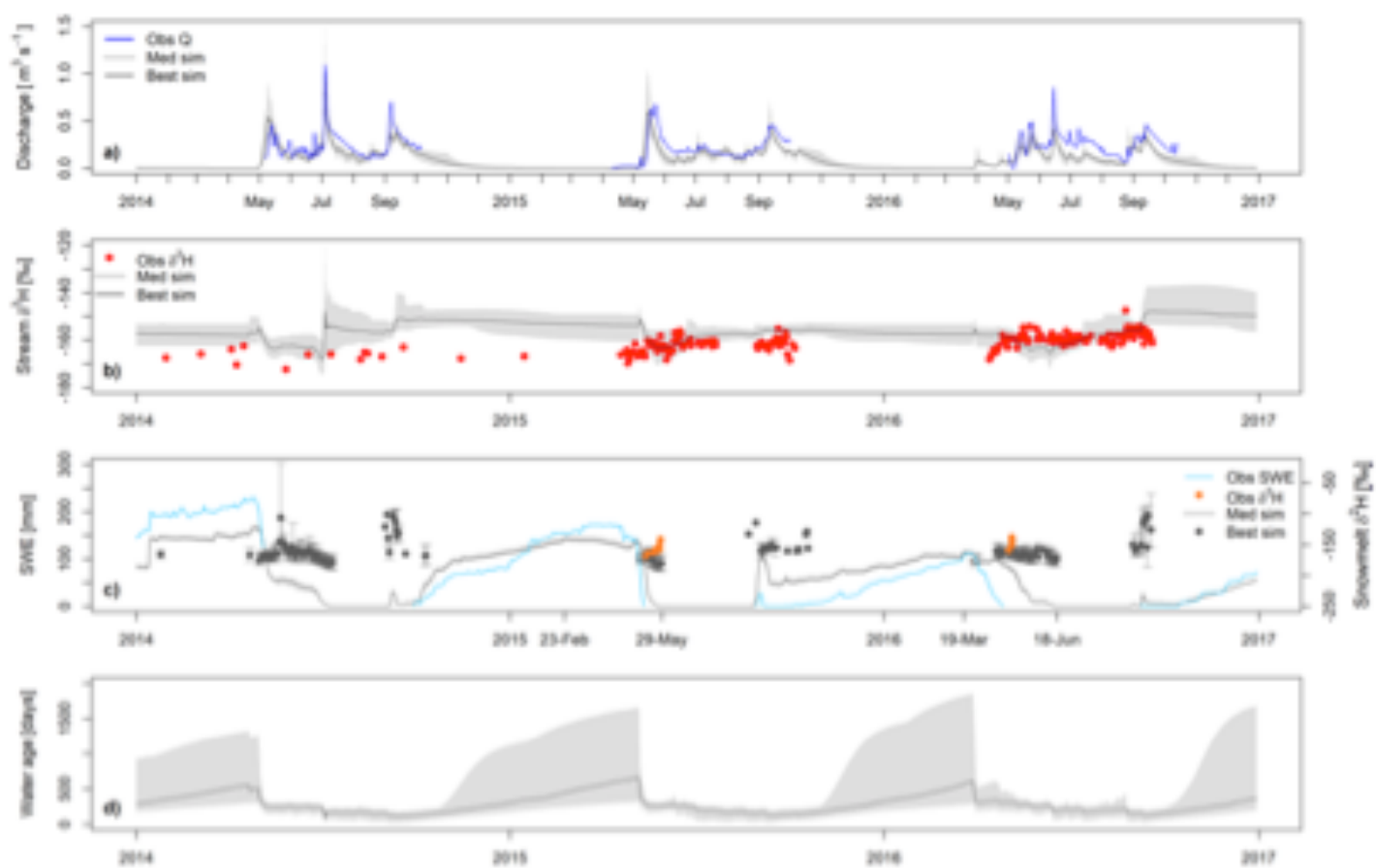
- GMWL
- Vegetation at Brunftland Burn
- Soil at Wolf Creek
- Vegetation at Dorset
- Vegetation at Dorset
- Soil at Brunftland Burn
- Soil at Krycklan
- Vegetation at Krycklan
- Soil at Moss Creek
- Vegetation at Moss Creek
- Vegetation at Dry Creek
- Soil at Dry Creek
- Vegetation at Dry Creek

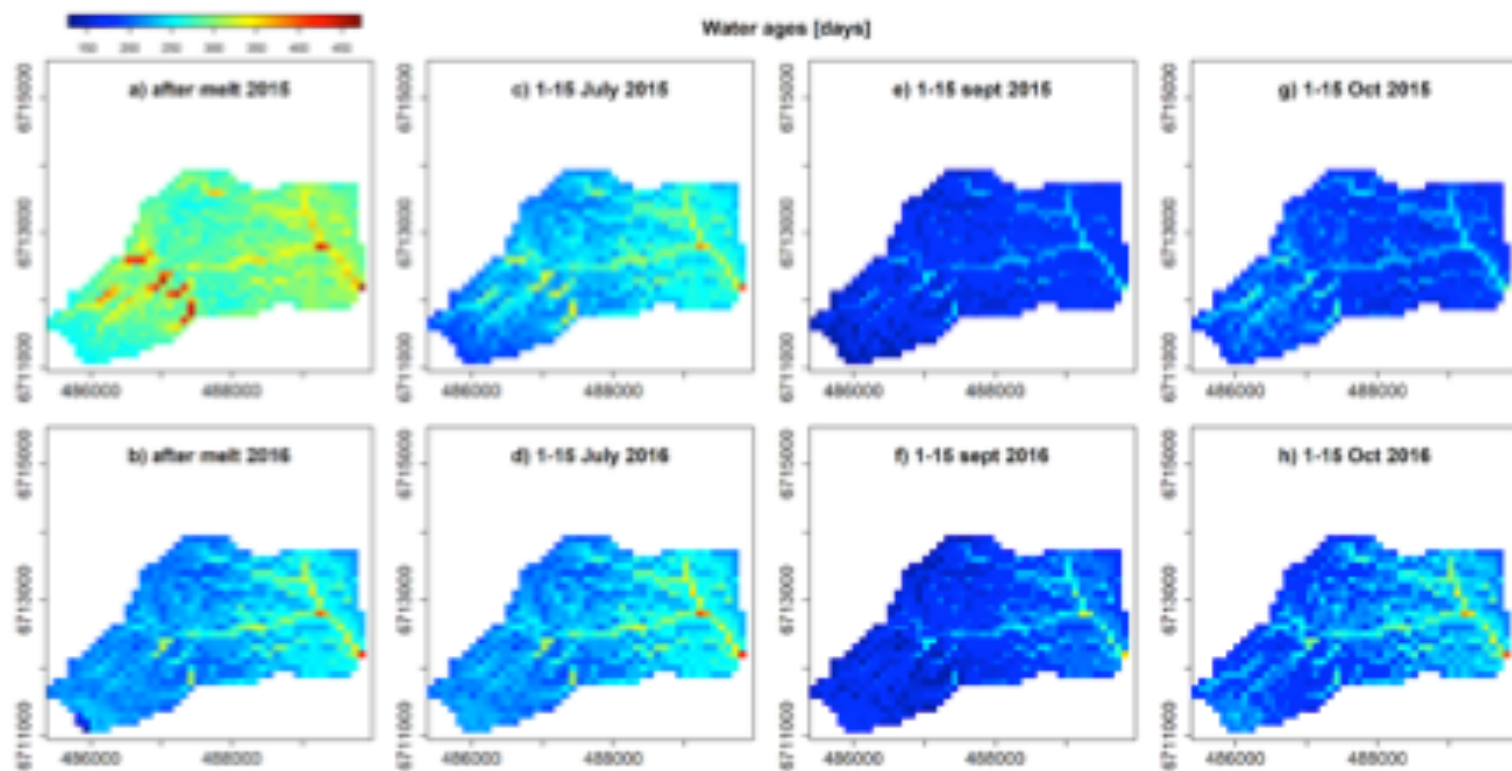


➤ VEWA – International inter-comparison project with common methodologies to trace water and isotopes through the soil-plant-stream continuum.



STARR - Ala-aho et al. (2017)







Thank you!