Anthropogenic aerosol perturbations in global storm-resolving simulations

Aerosols and clouds

Aerosols originate from natural processes, like dust storms or sea spray, but also from human activities, like biomass burning or fuel combustion. Despite their small size, aerosols strongly influence Earth's radiation balancex.[2] Aerosols scatter and absorb radiation referred to as aerosolradiation interactions but also modify the properties of clouds, as cloud droplets form on aerosol particles, referred to as aerosol-cloud interactions.[1]



(adapted from Boucher *et al.*¹)

Storm-resolving simulations

Storm-resolving simulations resolve atmospheric motions on scales of about $5 \,\mathrm{km}$ and consequently represent important atmospheric processes like convective updrafts that were parameterized previously.[6] Regional storm-resolving simulations revealed significant effects of aerosols on clouds and provided insights into the underlying processes and drivers.[4, 7, 5]

Our poster presents first results from global storm-resolving simulations. In contrast to regional simulations, global simulations include the coupling to large-scale circulation and in particular the budgetary constraints on precipitation due to the conservation of energy and water.[3]

Our simulations are performed with the atmospheric model ICON[10] in which aerosol perturbations are represented with the plume model MACv2-SP.[8] The sea surface temperature and sea ice are prescribed. Our analysis includes 40 days in the biomass burning season from 22th August to 30st September 2020.[9]





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<u>~</u> 0.75

between the perturbation (aero33) and reference (aero32).

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Cloud water and ice

<u>\$</u> 0.50 <u>وَ</u> 0.25 ප ත 10 $\widehat{}$ and my mm Simulation time (days)

Figure 3. Maps of water vapor, cloud water, and cloud ice; vertically integrated and averaged from 20th to 30th September (from day 30 to 40).









aero32





Figure 2. Global means of water vapor, cloud water, and cloud ice; vertically integrated. Shown is the difference



Figure 4. Global means of outgoing longwave and shortwave radiation at the top of the atmosphere. Shown is the difference between the perturbation (aero33) and reference (aero32).



Figure 5. Maps of the outgoing longwave and shortwave radiation at the top of the atmosphere over cloudy sky; vertically integrated and averaged from 20th to 30th September (from day 30 to 40).



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Radiation fluxes

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