

**Bringing together observations and
models to understand the world
and improve predictions of it**

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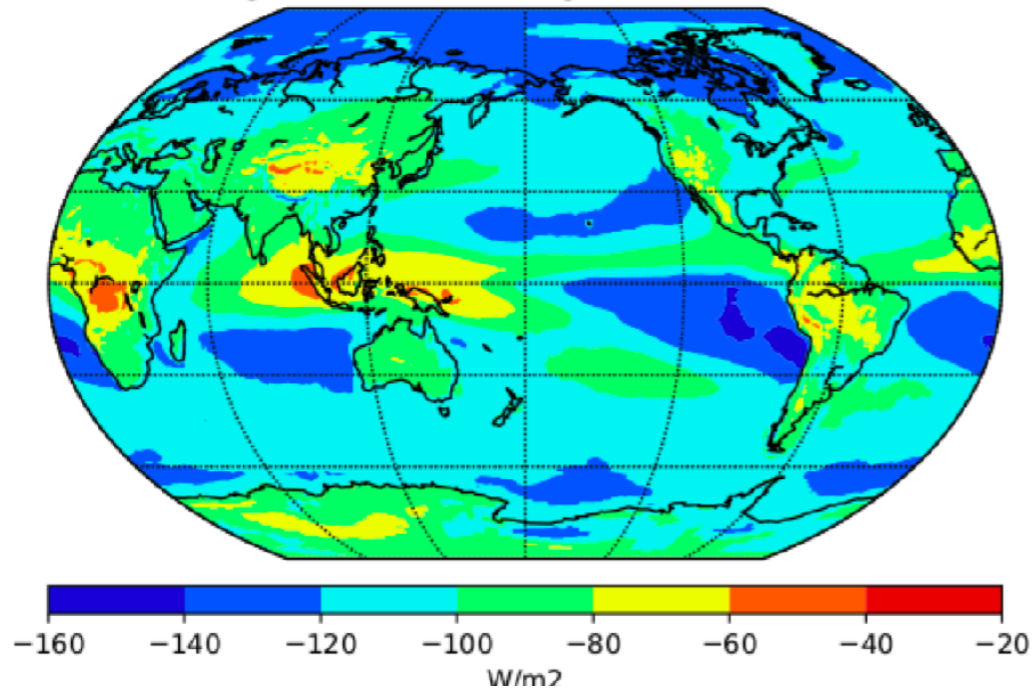
Thank you to lots of people

**Global observations set
the scene**

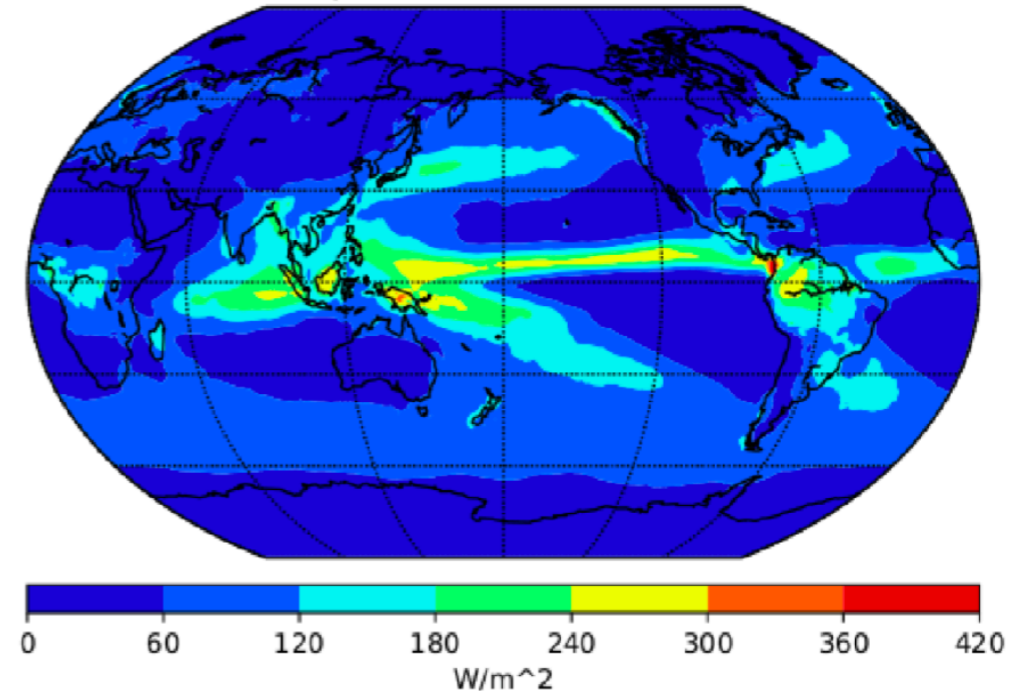
What is happening?

Climatologically, the “tropics” (30 N to 30 S) are close to Radiative Convective Equilibrium (RCE).

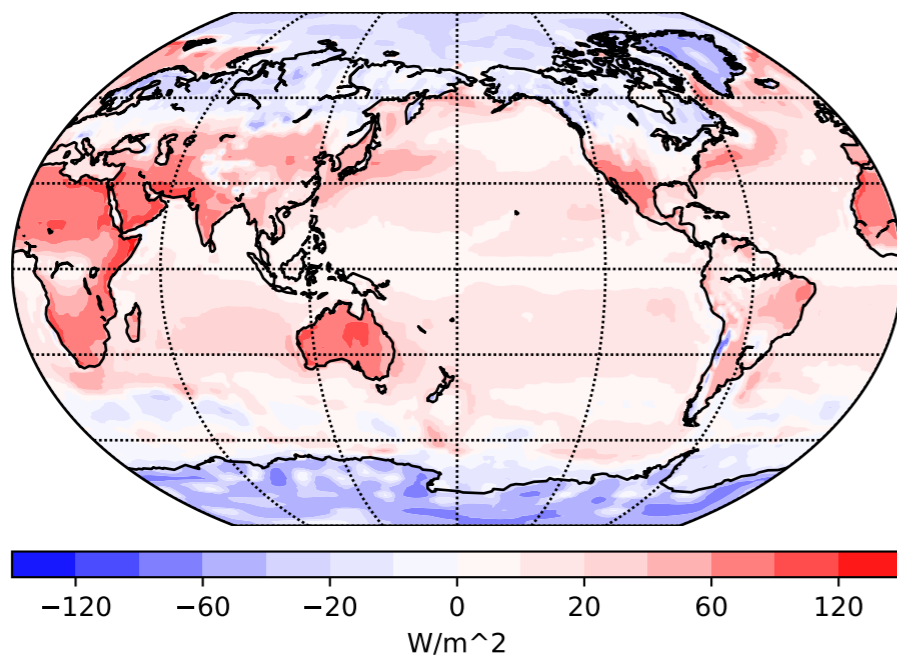
Averaged radiative cooling - CERES 2001/2009



Precipitation - GPCP 2001/2009



Sensible Heat Flux - NCEP 2001-2009



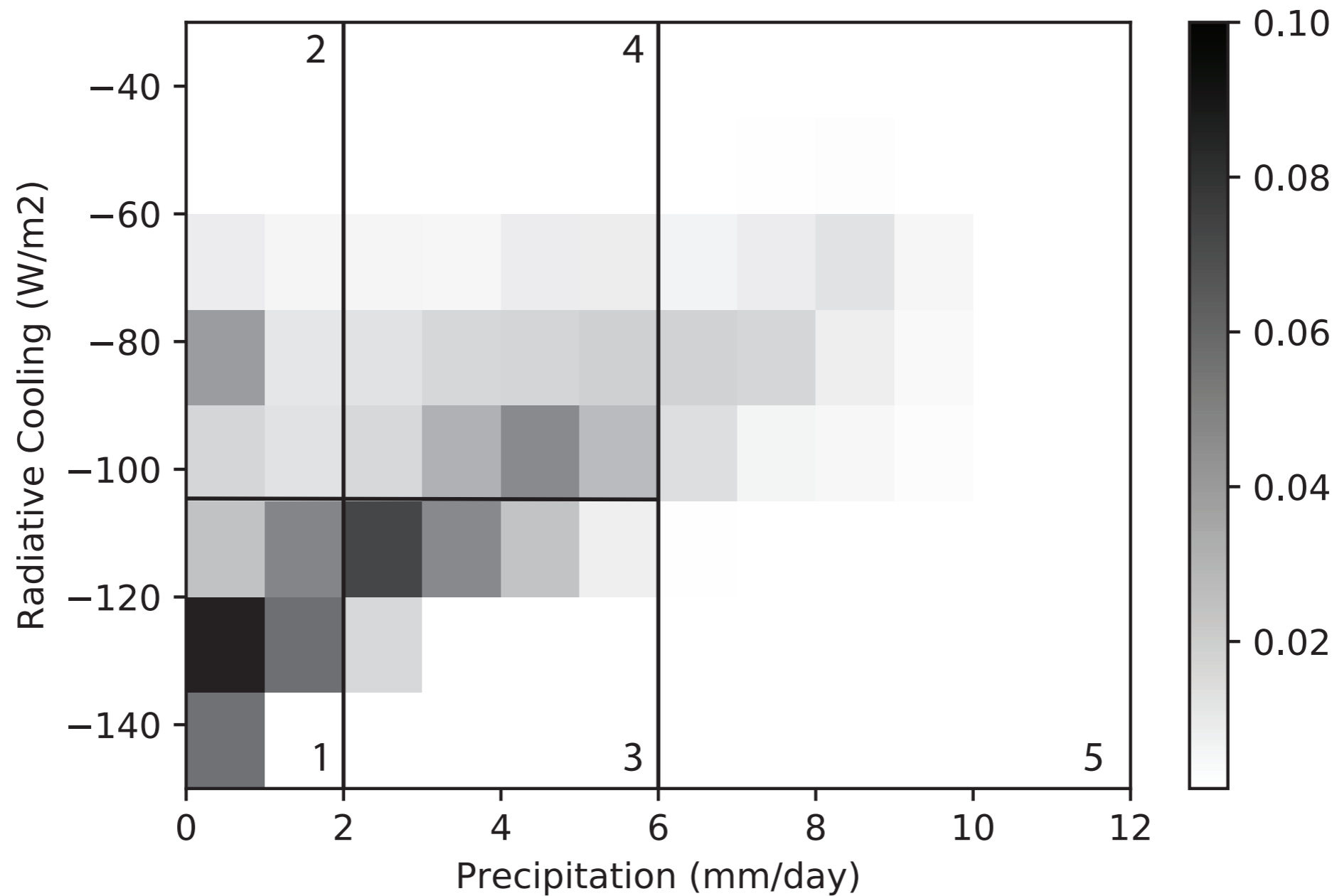
$$\langle Q_{Rad} \rangle = -104.5 W/m^2$$

$$\langle P \rangle = 85.2 W/m^2$$

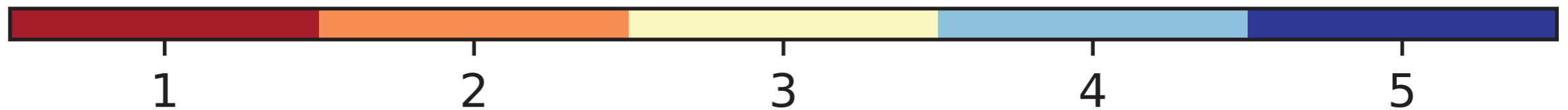
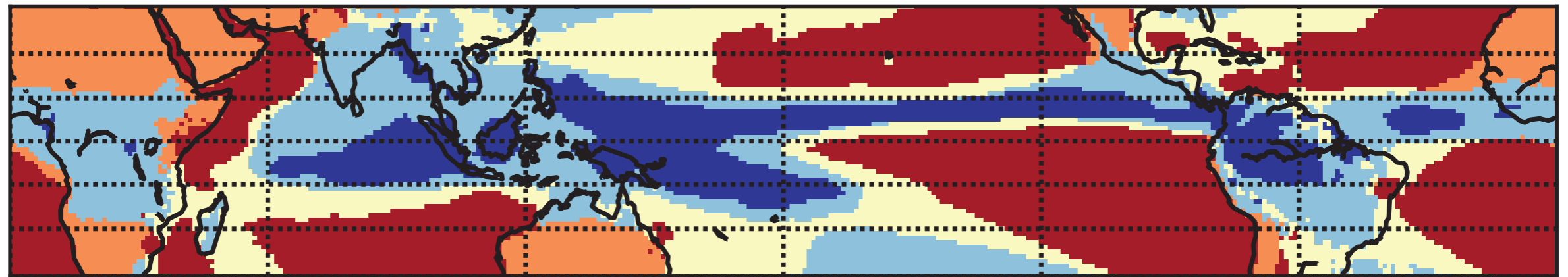
$$\langle SSHF \rangle = 22.6 W/m^2$$

$$Residual = 3.3 W/m^2$$

A histogram of the climatological radiative cooling and precipitation in the tropics shows that this is not true locally!

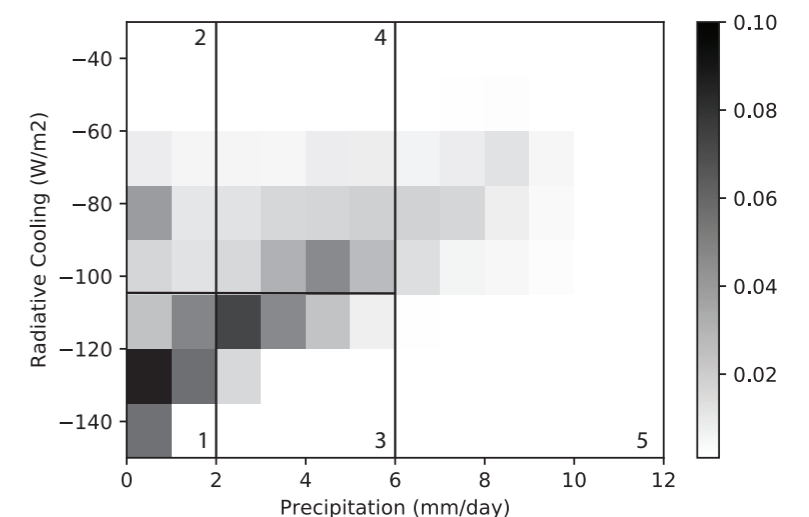


It is clear that the tropical RCE is only established by the “cooperation” of very different regions in the tropics.

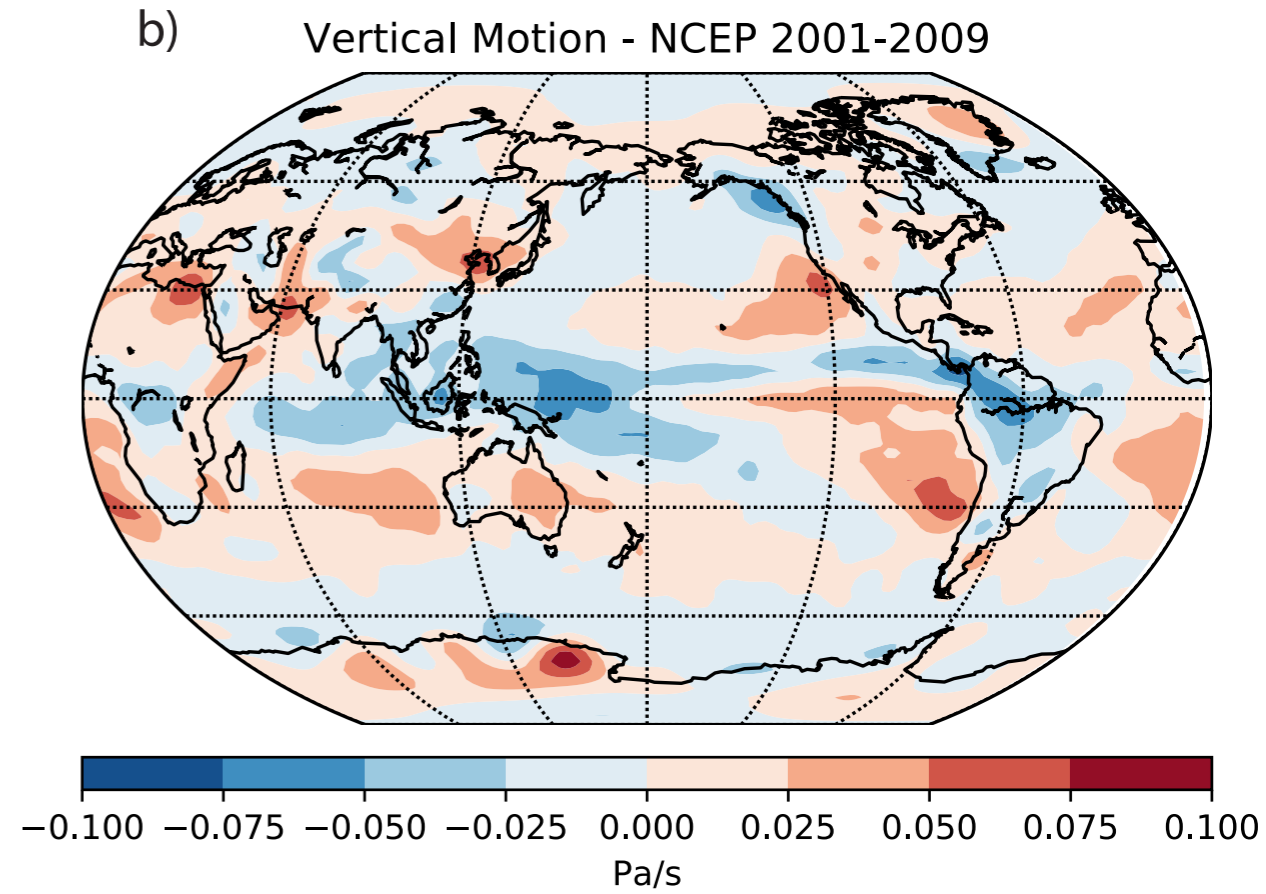
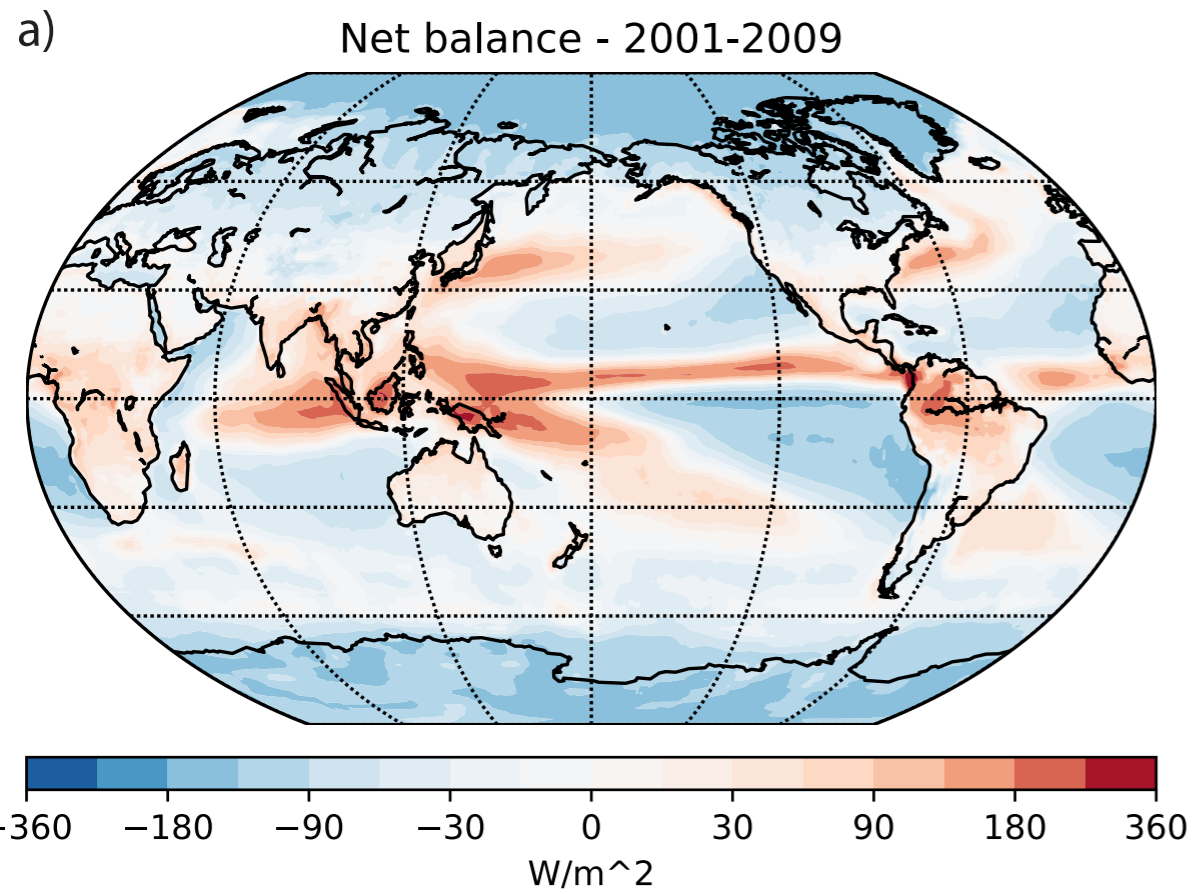


Regions of the Mean histogram

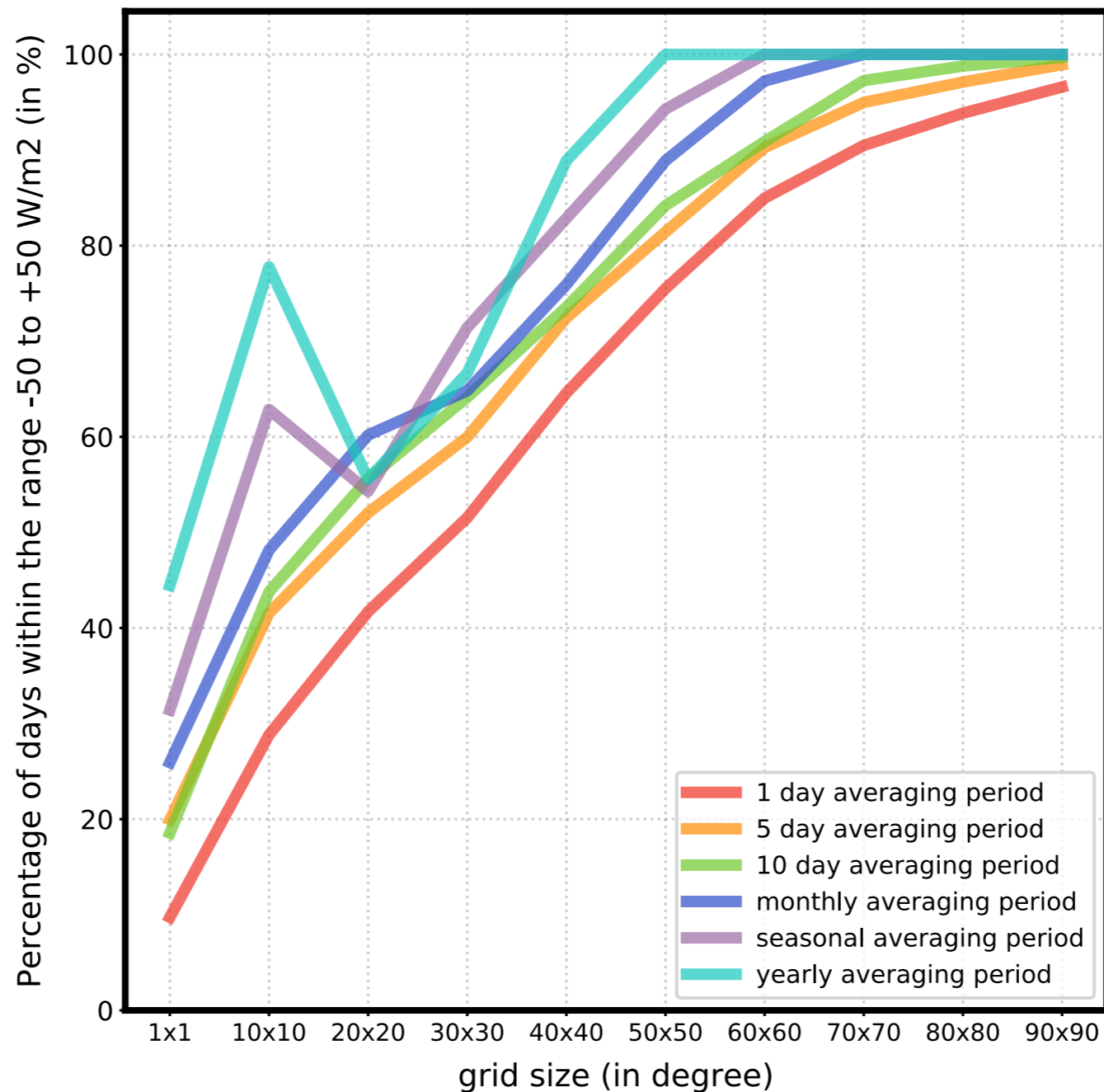
- 1 = Strong cooling, no rain**
- 2 = Weak cooling, no rain**
- 3 = Strong cooling, medium rain**
- 4 = Weak cooling, medium rain**
- 5 = Strong rain, never cools strongly**



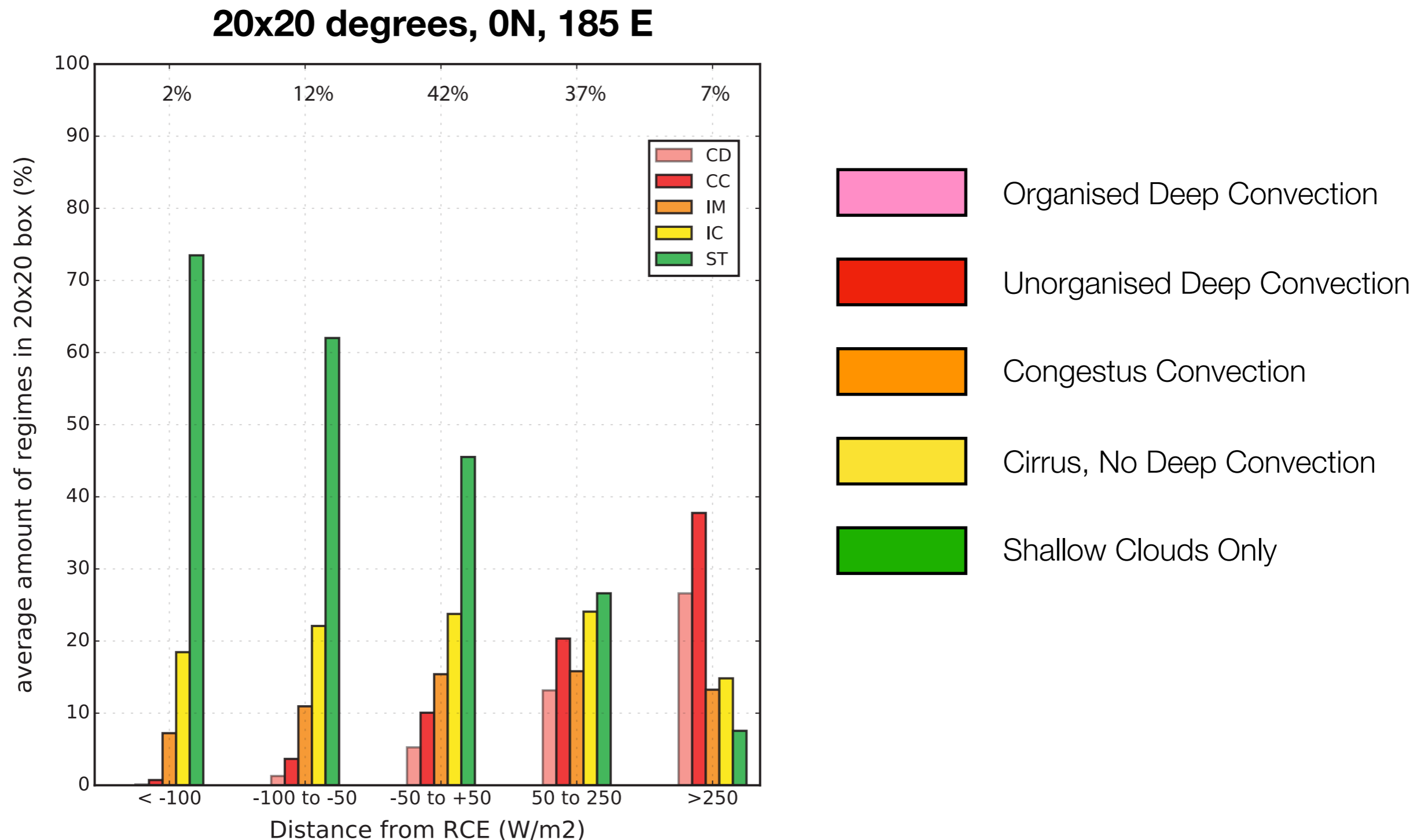
It is the circulation that facilitates the large-scale equilibrium



We define a “distance from RCE” as the difference between radiative cooling, sensible heat flux, and precipitation and ask what fractions of days are in RCE (-50 to + 50 W/m² difference) as a function of spatial and temporal averaging scale.

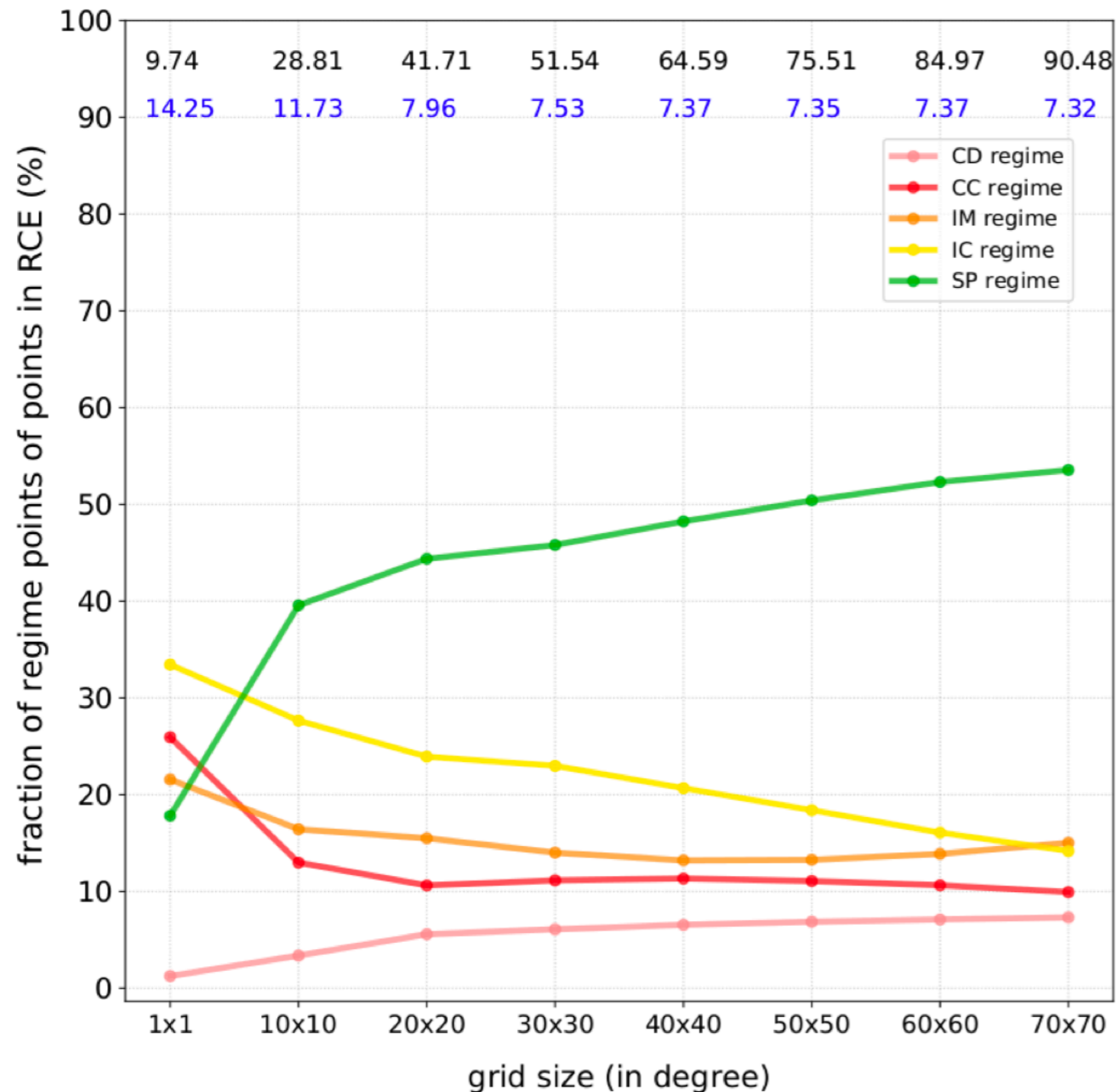


In relating RCE to cloud states, we find that achieving RCE requires the “*right*” mix of deep and organised convection and suppressed conditions.



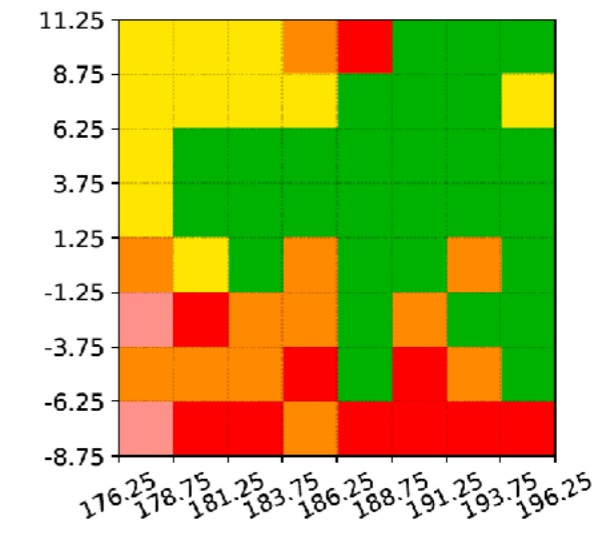
The ratio of active to suppressed conditions appears to be quite scale-invariant.

West Pacific (0 S, 185 E)

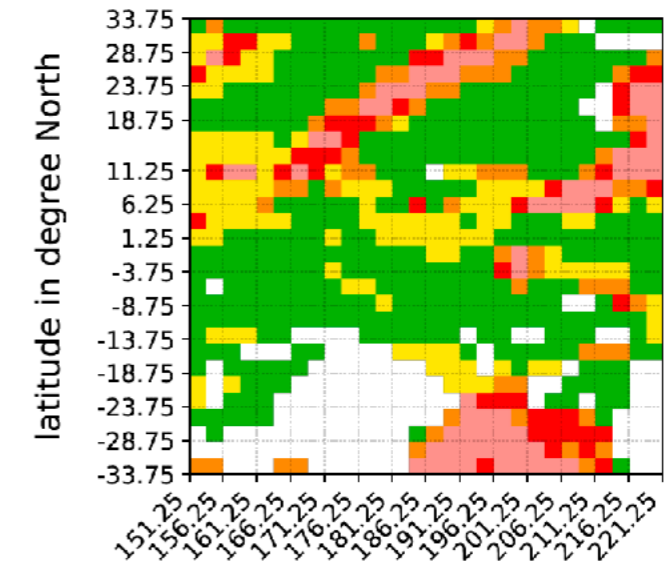


Snapshots

Distance from RCE: 8.32373835633

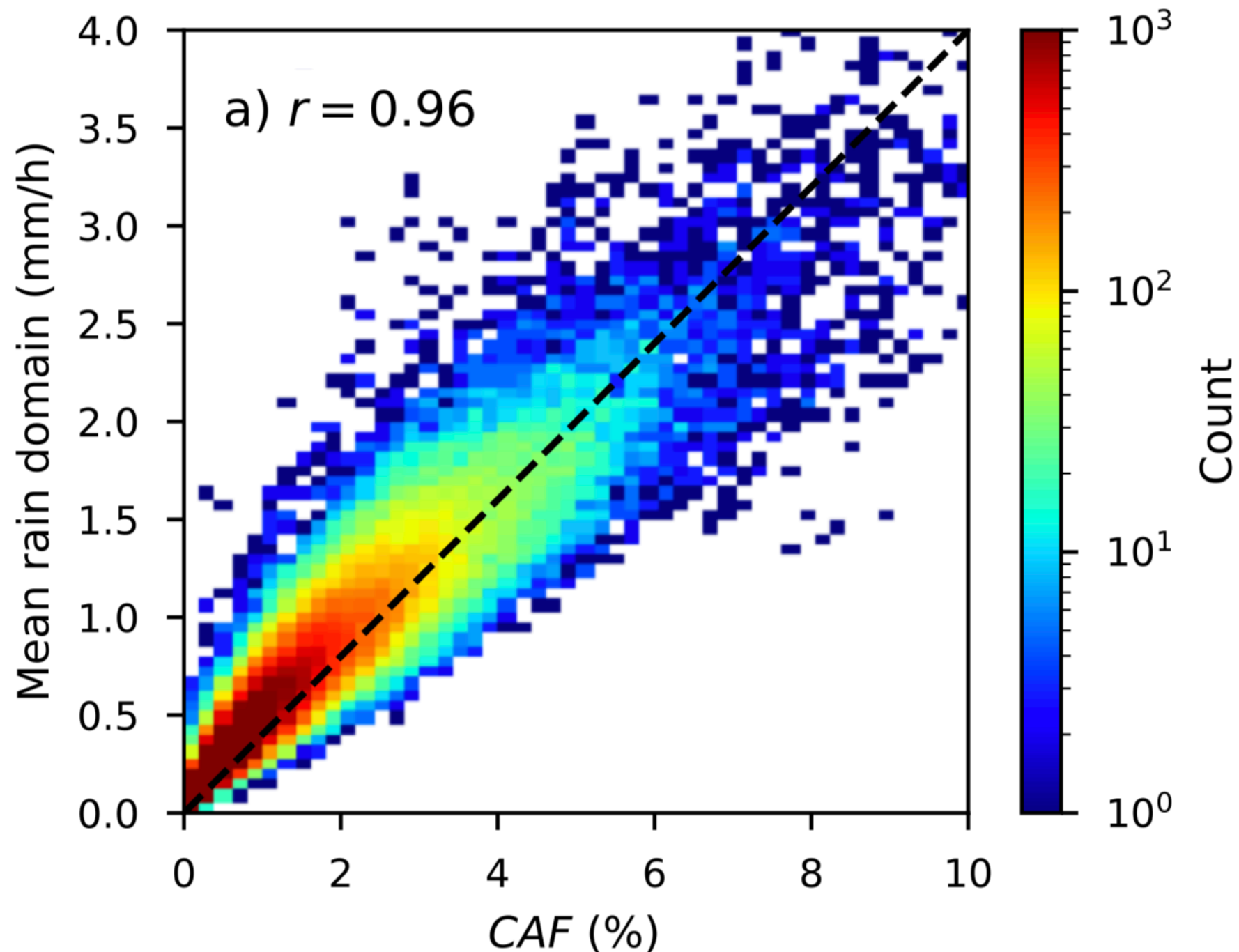


Distance from RCE: -5.54540884029



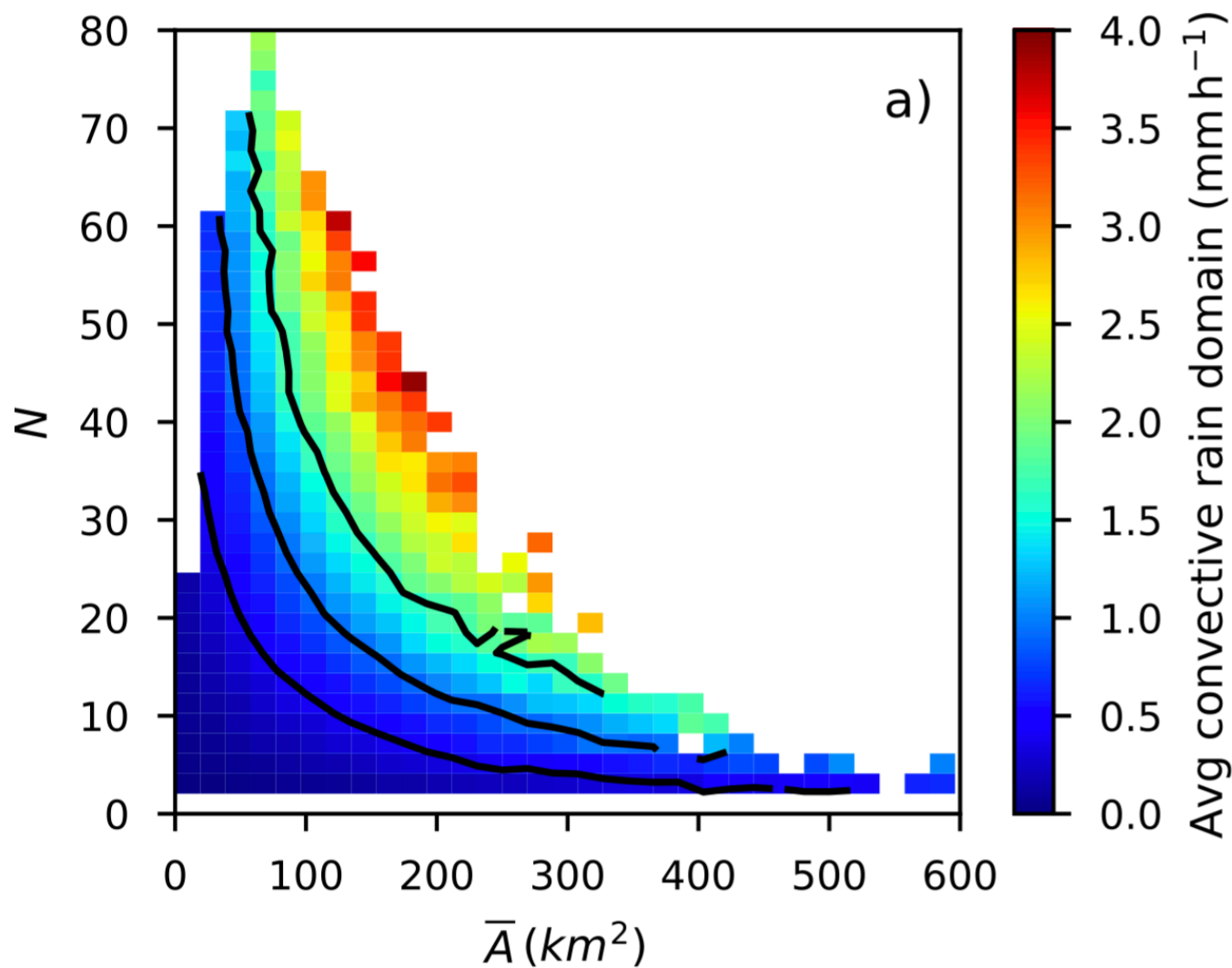
**Local observations
provide process details
How might it work?**

Radar observations from Darwin show that area-average rainfall in a convecting environment relates strongly to the area that is raining

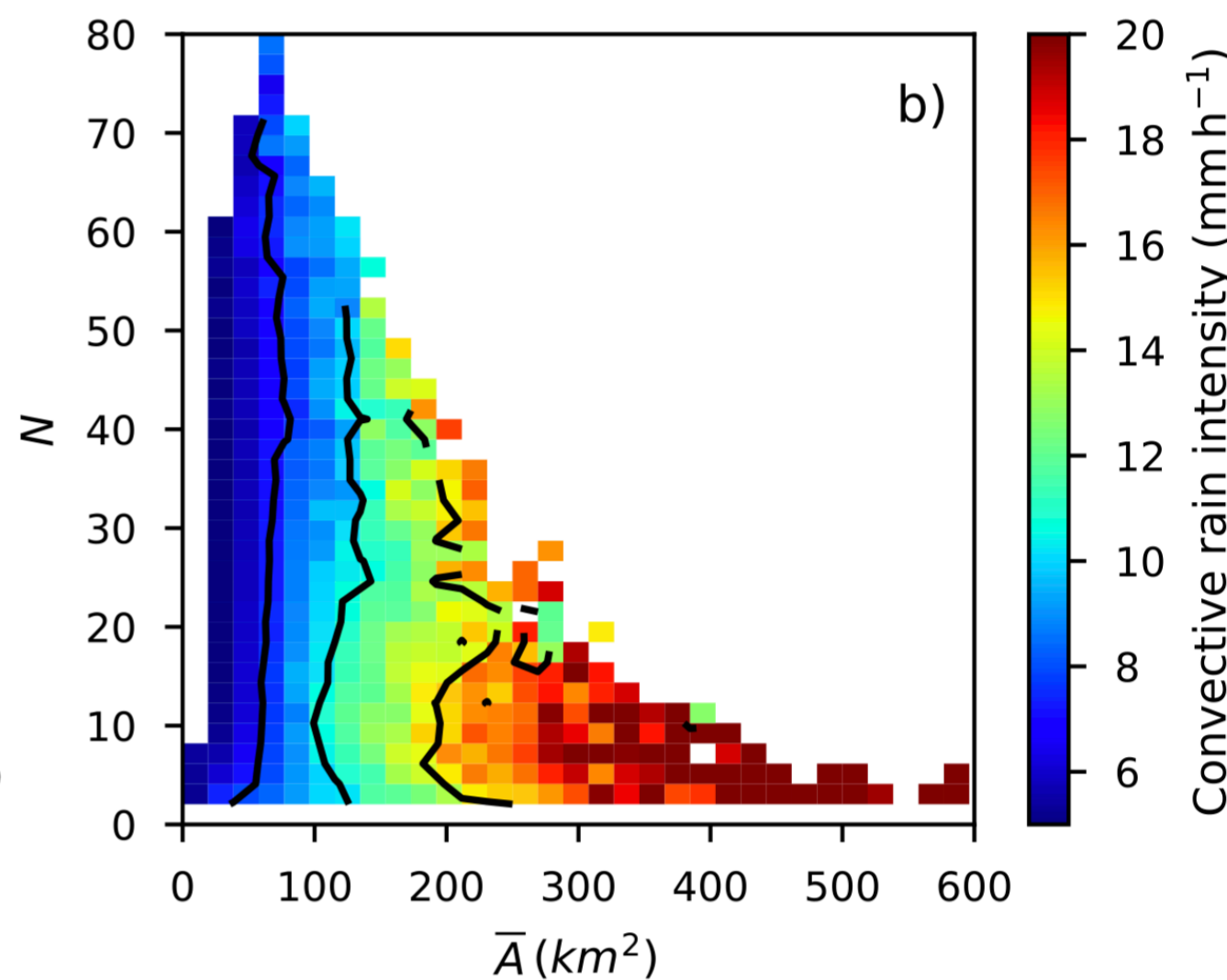


Splitting the area fraction into number of clouds and their size reveals interesting behaviour

Area mean rainfall

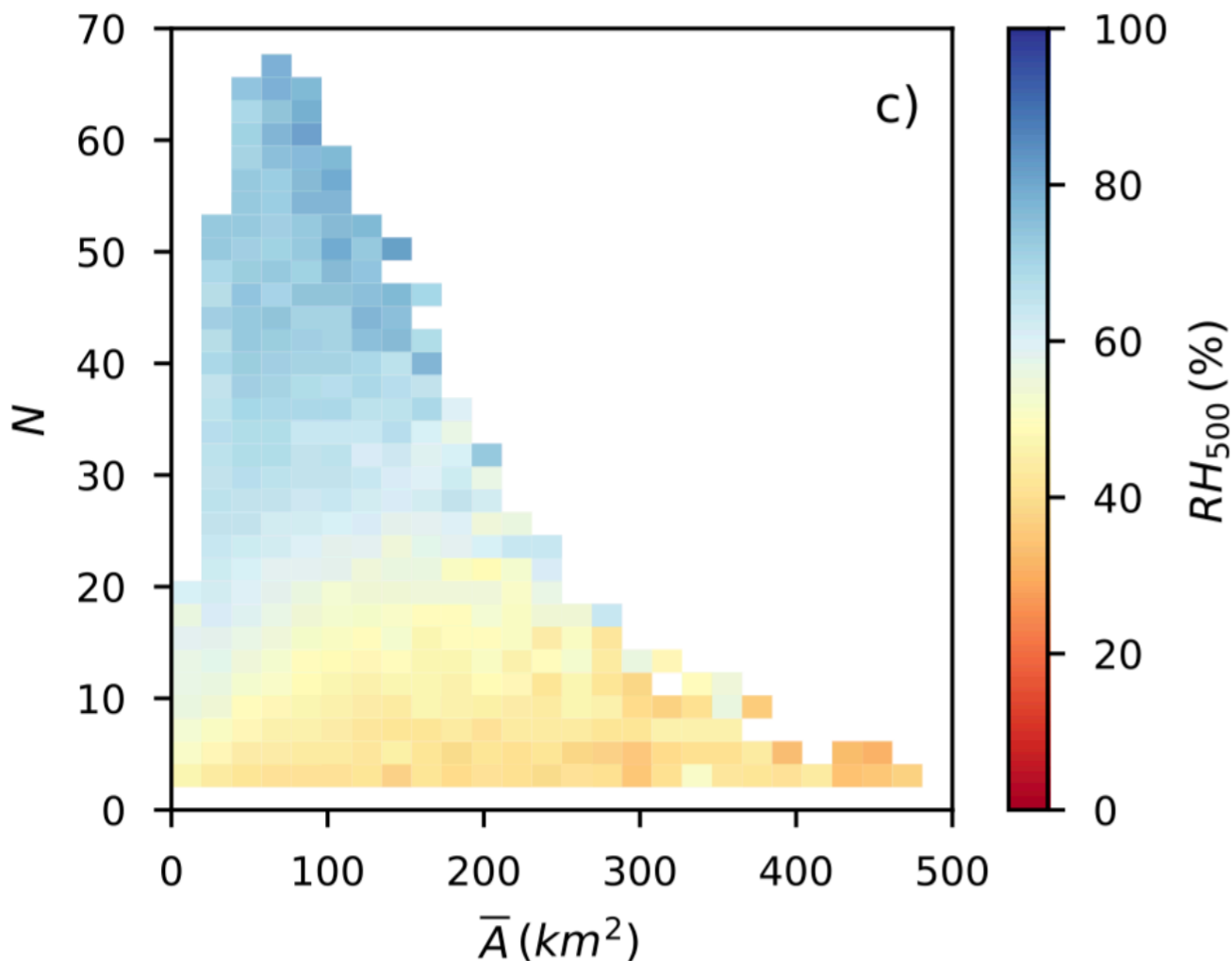


Local rain intensity

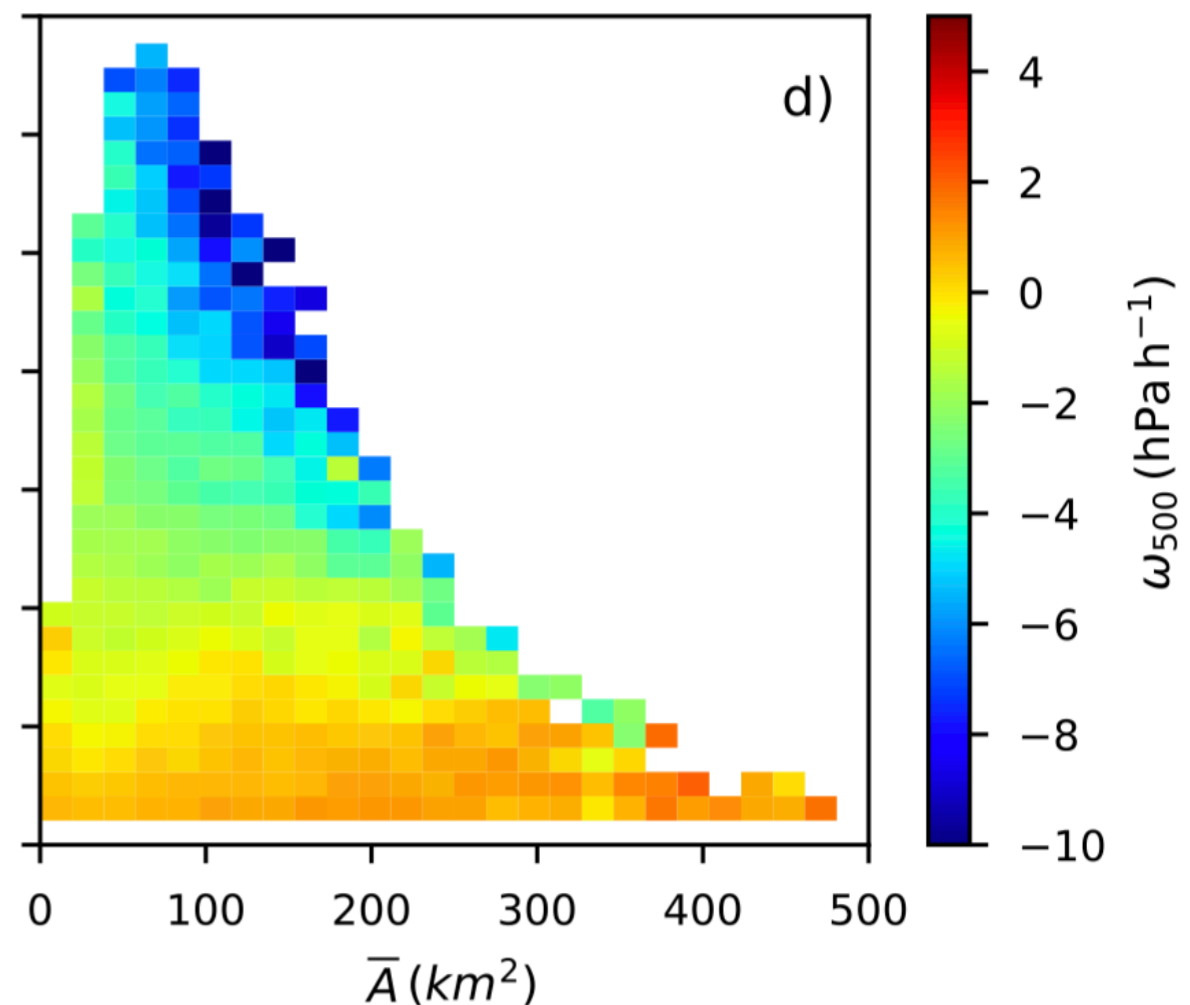


Investigating the relationship of the convective cloud ensembles characteristics to the large scales sets the scene for parametrisation

Mid-tropospheric RH



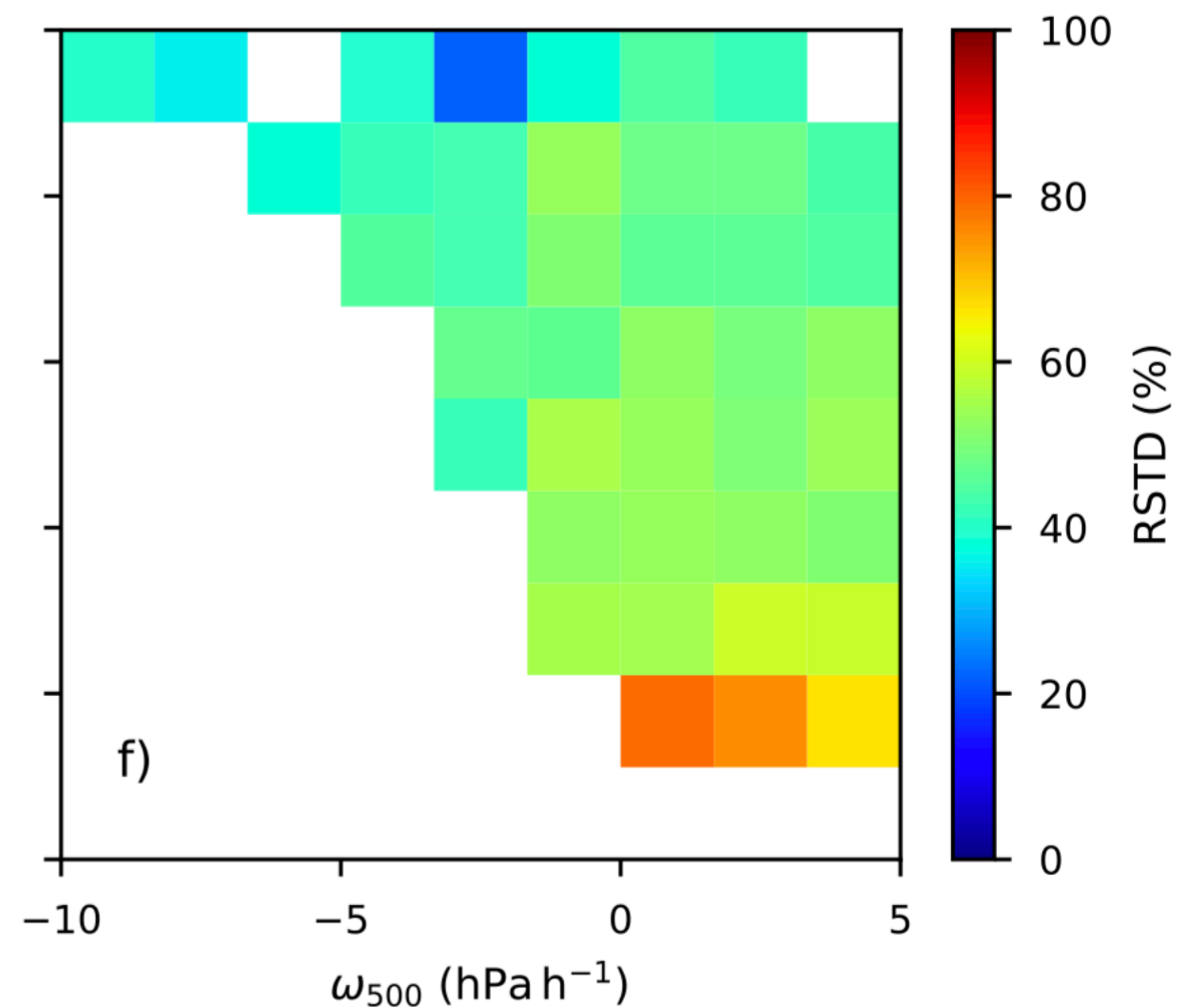
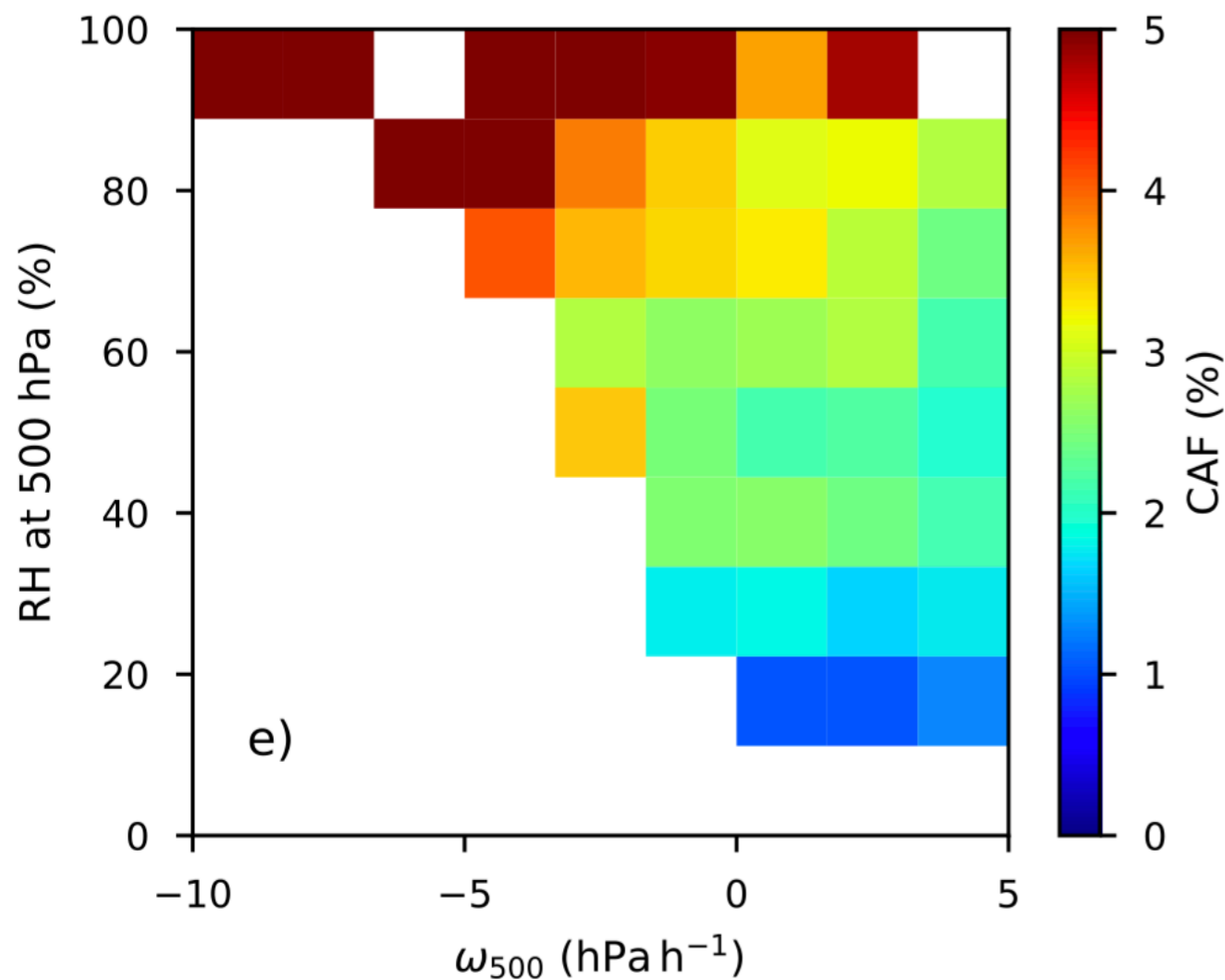
Mid-tropospheric vertical motion



Investigating the relationship of the convective cloud ensembles characteristics to the large scales sets the scene for parametrisation

Cloud area fraction as $f(\text{RH}, w)$

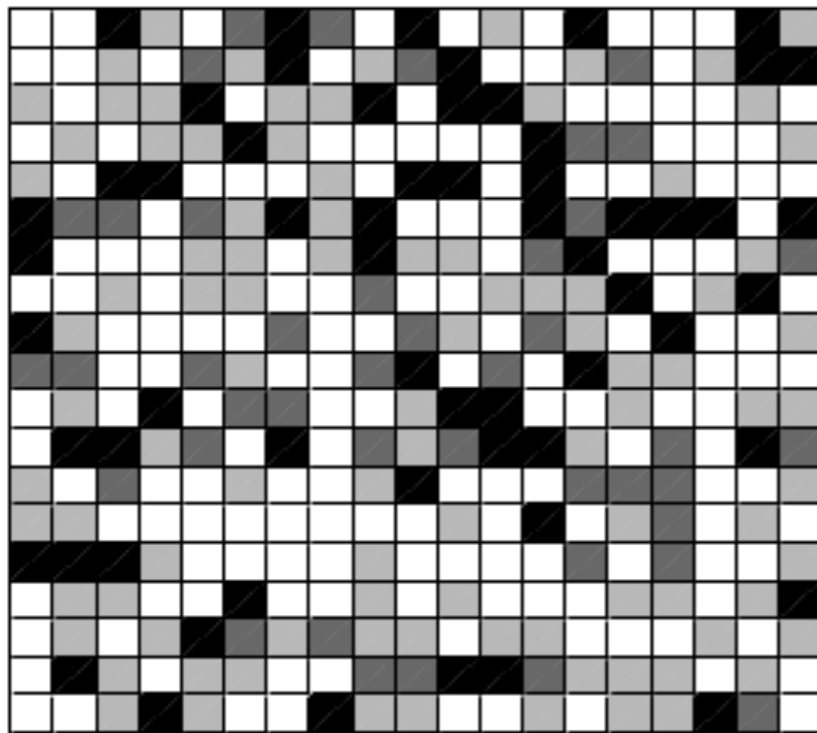
Within-bin standard deviation



**Translating what we've
learned into models**

Based on our findings we divide the mass-flux that models use to describe convection into area and velocity and predict area with a statistical model

$$\overline{\rho w' s'} \sim M_c (s_c - \bar{s}) \quad M_c \approx \rho \sigma_c w_c$$



States considered:

- Clear
- Congestus
- Deep Convection
- Stratiform Clouds

$$R_{01} = \frac{1}{\tau_{01}} \Gamma(C) \Gamma(D),$$

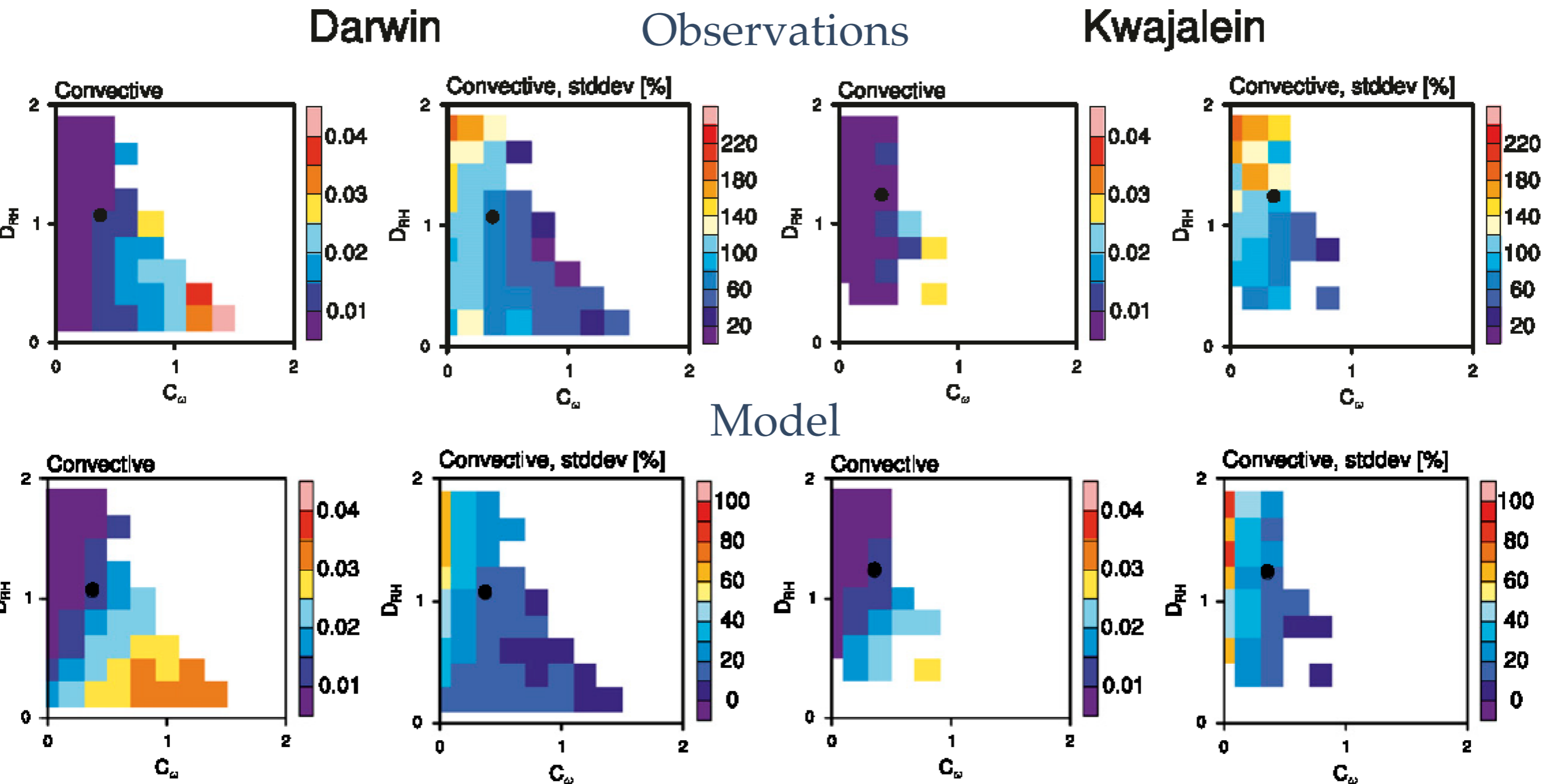
Clear to congestus

$$R_{02} = \frac{1}{\tau_{02}} \Gamma(C) (1 - \Gamma(D)),$$

Clear to deep

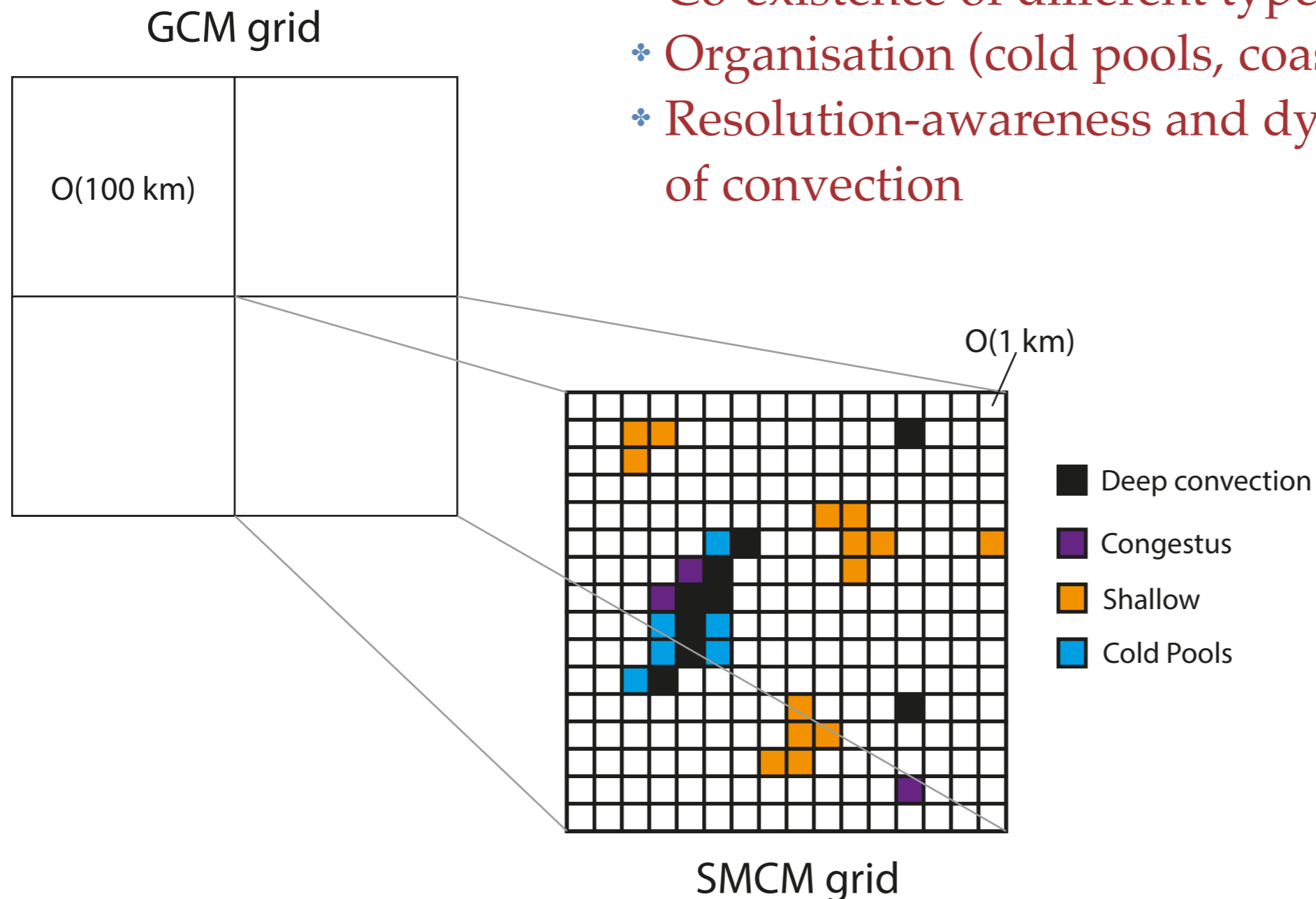
So large C (ascent) and large D (low RH) favour congestus, while large C and small D (high RH) favour deep convection.

The model captures the the observed behaviour well



The new framework can in principle deal with all issues currently being discussed in the parametrisation community

- ❖ Stochastic behaviour of convection
- ❖ Co-existence of different types of convection
- ❖ Organisation (cold pools, coast lines, other)
- ❖ Resolution-awareness and dynamic switching off of convection



The vertical cloud structure will be calculated separately for each cloud type using an entraining plume or any other cloud model

