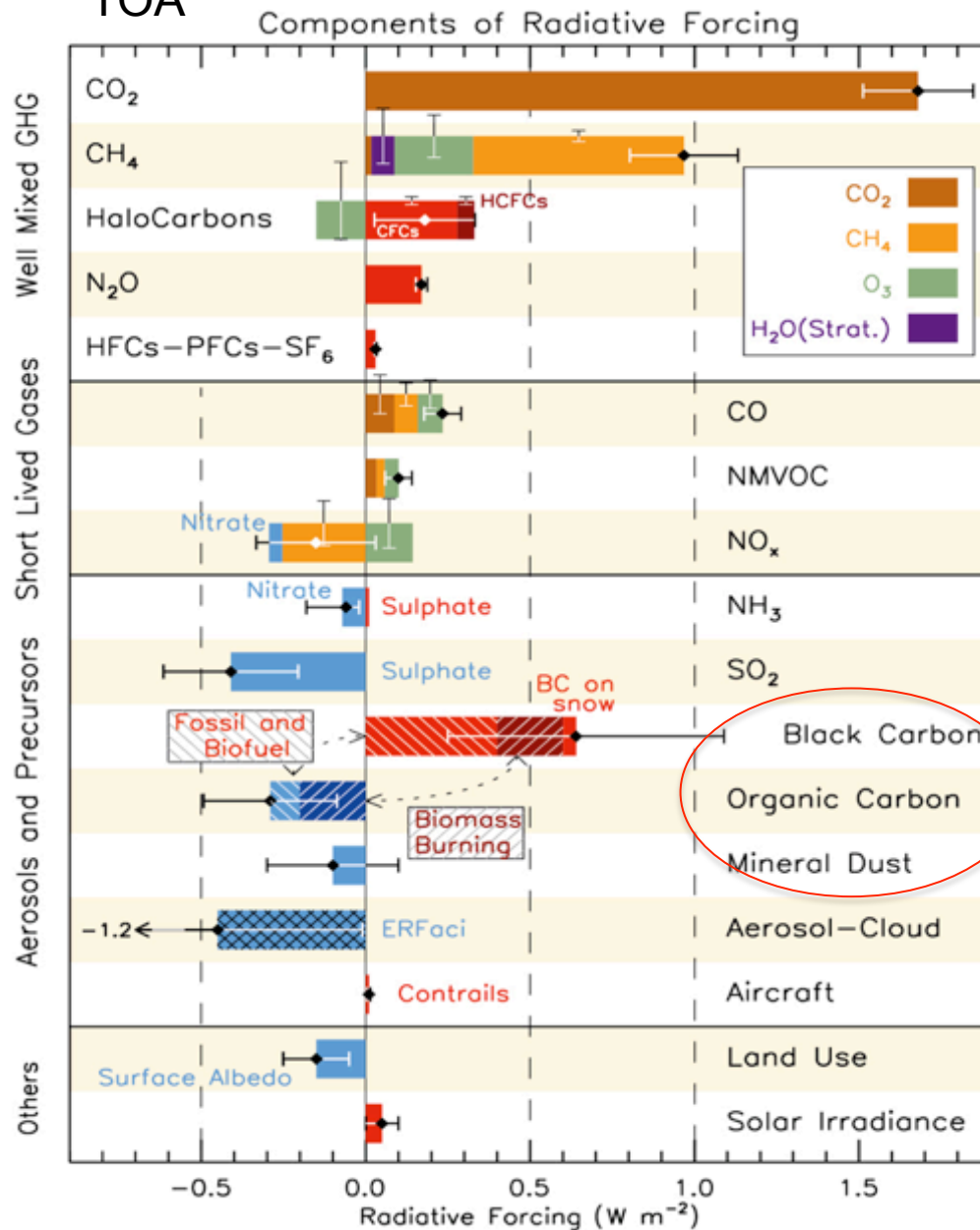


Impacts of aerosol snow-darkening effects on Eurasian hydroclimate and heat waves during boreal spring and summer

William K. M. Lau
ESSIC, University of Maryland

Coauthors: Jeong Sang/Gonju U, M. K. Kim/Gonju U. ,
K. M. Kim/GSFC, R. Koster/GSFC,
T. J. Yasunari/Hokkaido U.

TOA



- Aerosol-Cloud-Climate (ACC) feedback is important for global and regional climate change

- ACC feedback processes involve both anthropogenic and natural aerosols

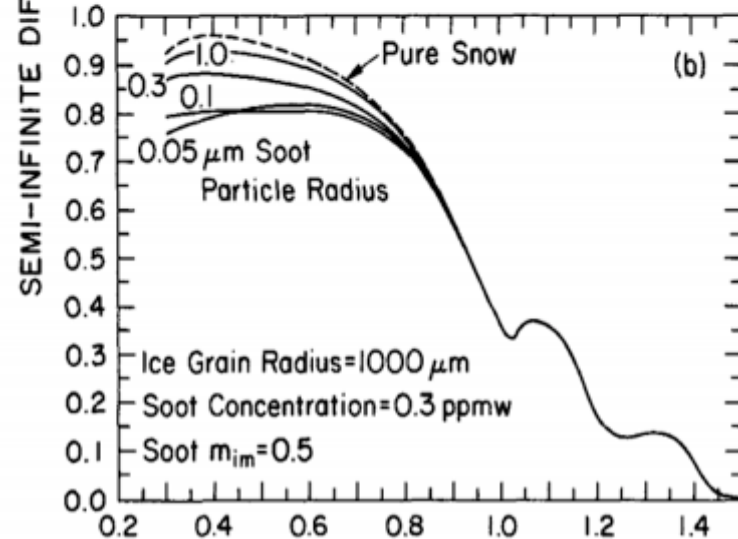
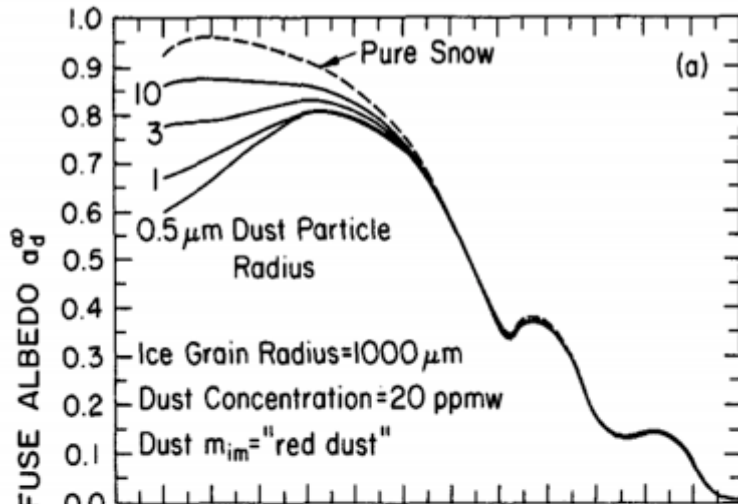
- Globally, natural aerosols are 6- 10 times more abundant than anthropogenic aerosols

- Absorbing aerosols (BC, OC and mineral dusts) amplify atmospheric heating

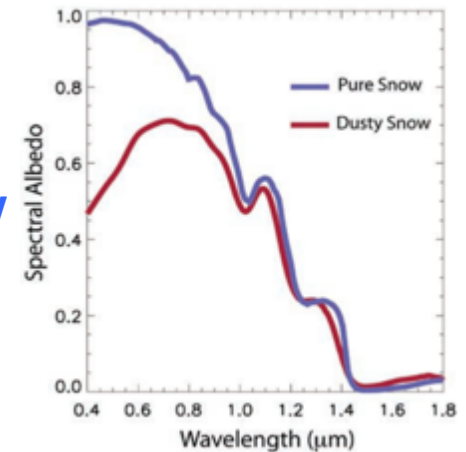
- Snow-darkening by absorbing accelerate snowmelt, amplify global warming by snow-surface albedo feedback

Snow Darkening Effect (SDE)

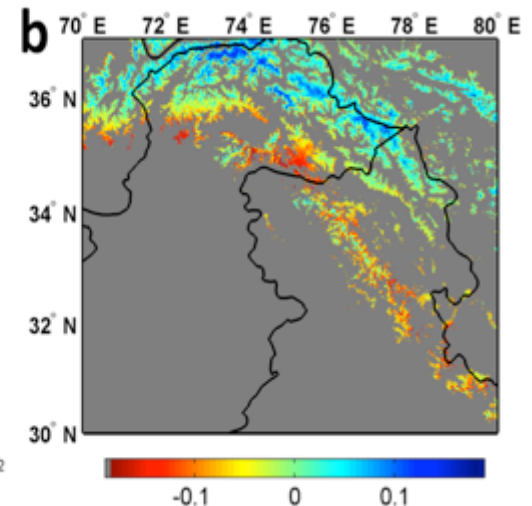
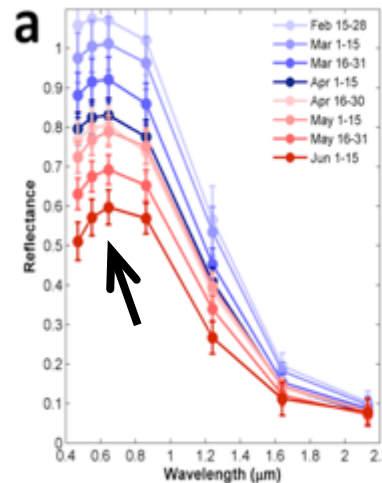
Dust and black carbon reduce snow albedo in the visible wavelengths



Warren and Wiscombe, 1980



Painter et al., 2009



Gautam et al., 2012

GEOS5 model with a new snow-darkening GOddard SnoW IMpurity (GOSWIM) module

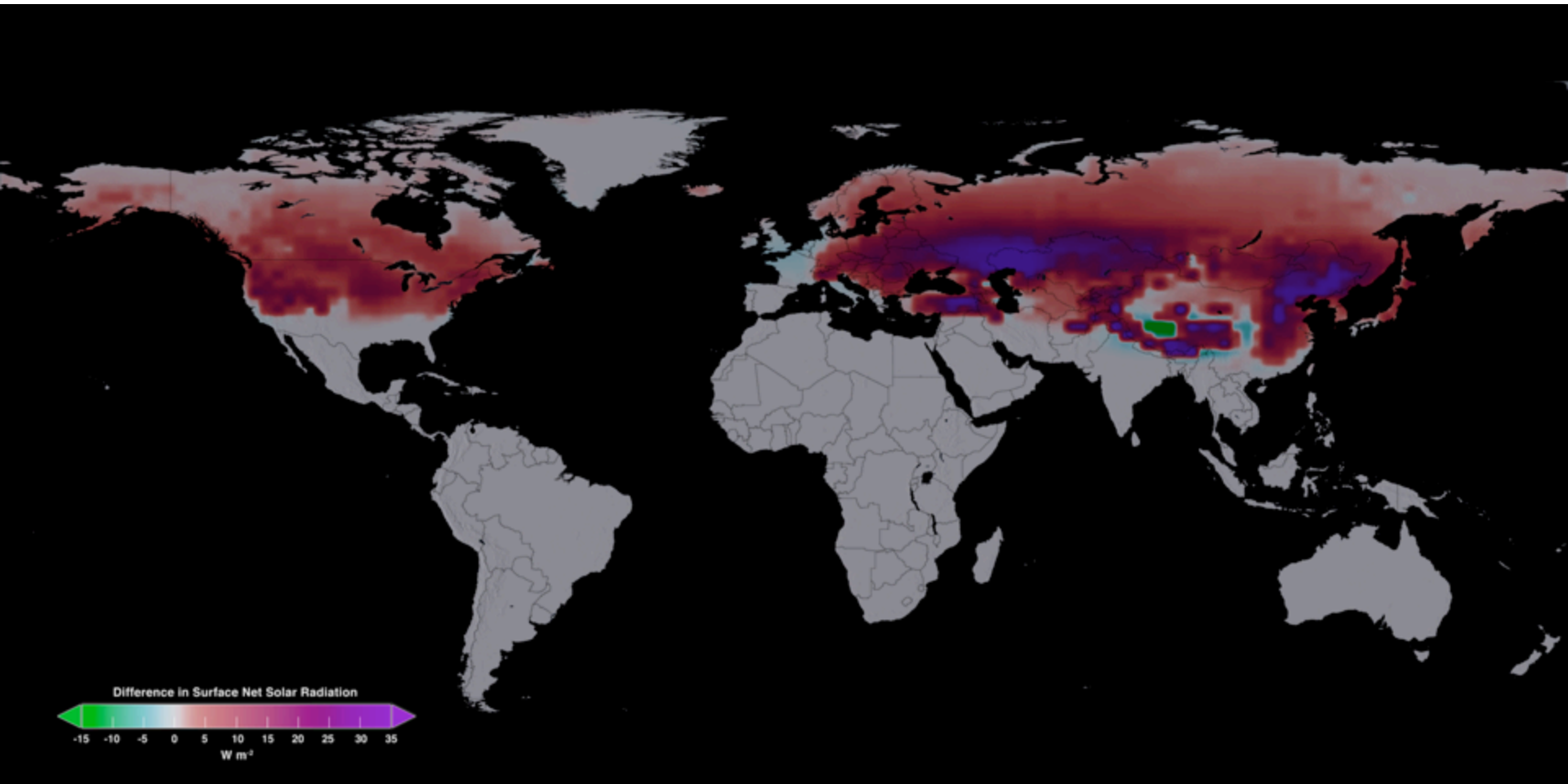
| Model | Period | ENS | Experiment | |
|-----------------|---|--------------------------|------------|-------------------------------|
| NASA- GEOS_5 | 2002.01.01 ~2011.12.31 (10 years) | ENS_MEAN (10 members) | SDE | with snow darkening effect |
| | | | NSDE | without snow darkening effect |

Snow impurities include dust, BC, and OC (natural + anthropogenic)

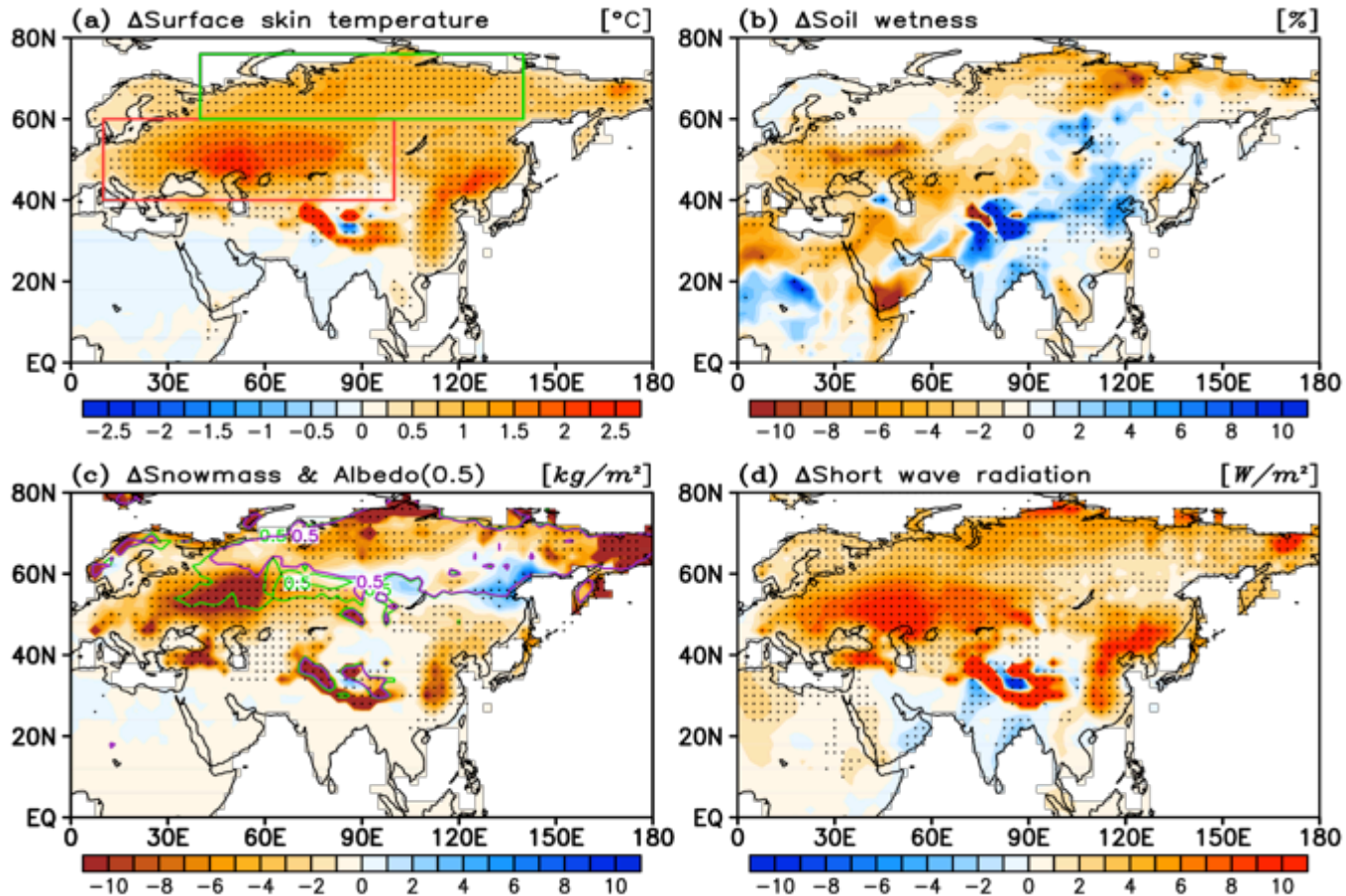
Snowdarkening effects by absorbing aerosols on:

1. Global climate, regional land surface energy and water balances (Eurasia/Central Asia, East Asia, North America and Tibetan Plateau (Yasunari et al. 2015, JGR)
2. [Hydroclimate feedback and heatwaves over Eurasia](#) (Lau et al., 2018, JGR, in press)
3. Effects on Asian monsoon precipitation : A revisit of the Blandford hypothesis (Lau and Kim, 2018, in preparation)

Snow-Darkening Effect (SDE): reduced snow albedo by deposition of absorbing aerosols (dust, BC, OC) leads to increased in net surface solar radiation, during the boreal summer melting season (MAM)

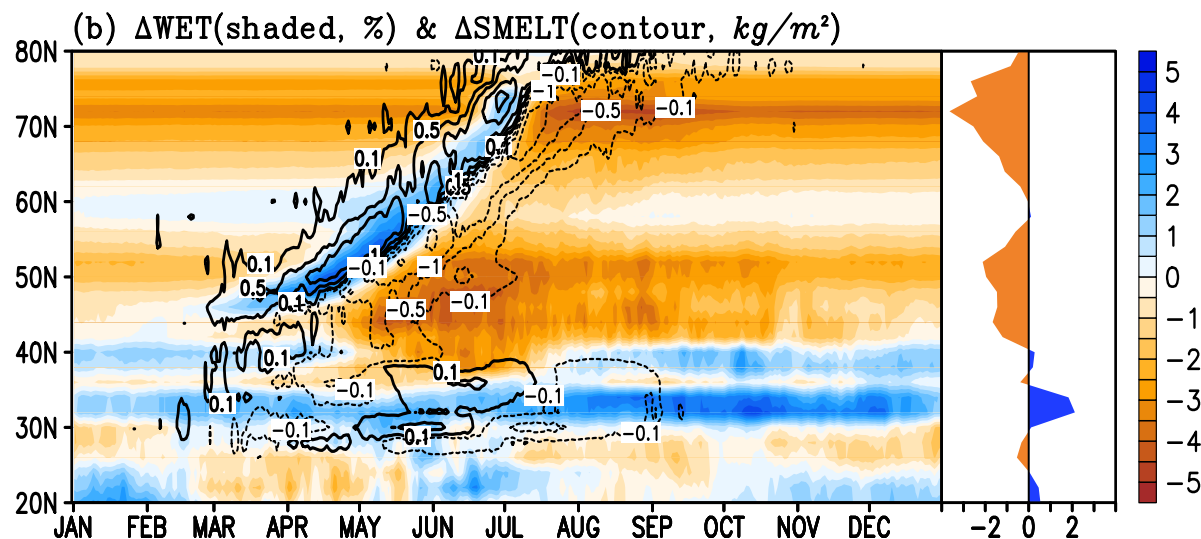
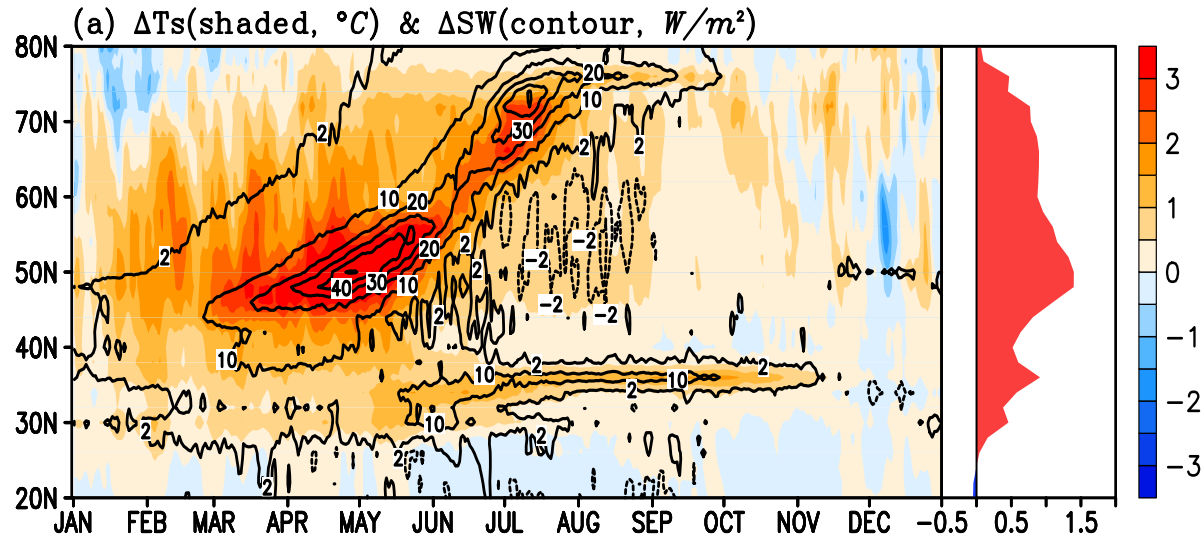


SDE induced annual mean climate change



- warmer land $> 40^{\circ}\text{N}$, entire Eurasia, TP, and East Asian regions
SDE increases surface absorption of SW radiation, by albedo reduction of snowcover regions, leading increased snowmelt, but
- changes in soil wetness and relationship with temperature is not obvious

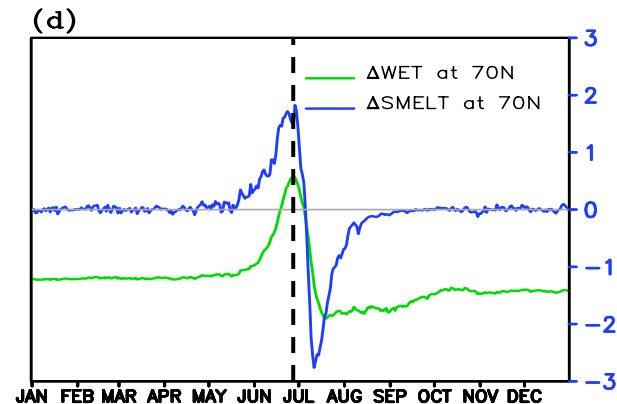
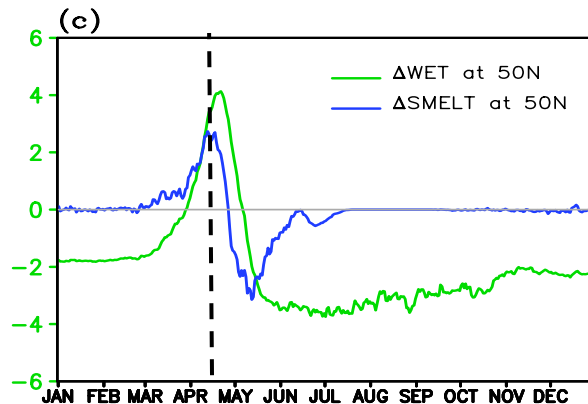
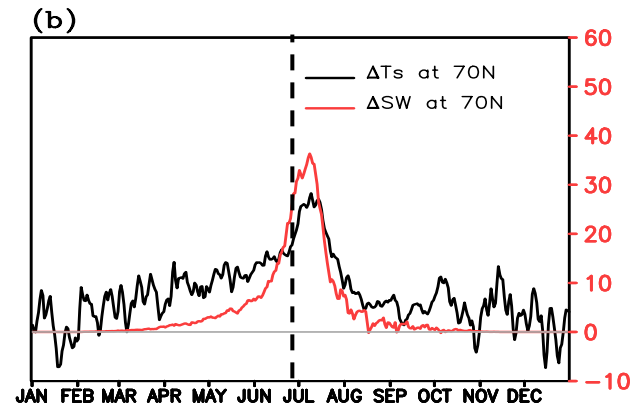
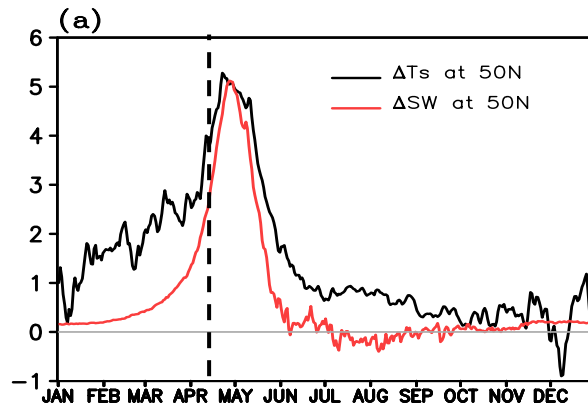
Accelerated snowmelt, warming, soil wetting and drying following the seasonal migration of the snowline over Eurasian land (0-140 E).



Wet-First
Dry-Later
(WFDL)

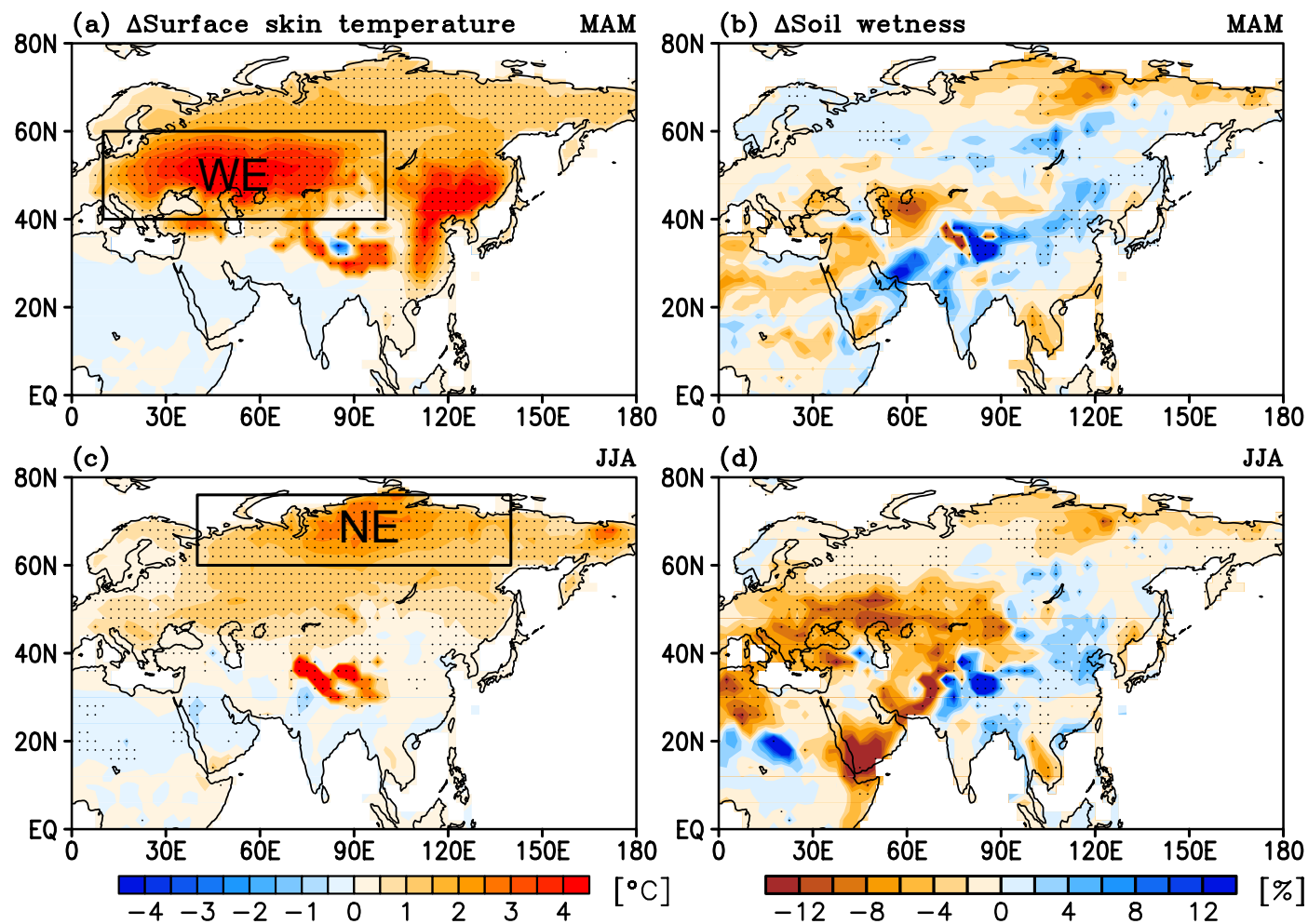
50 N

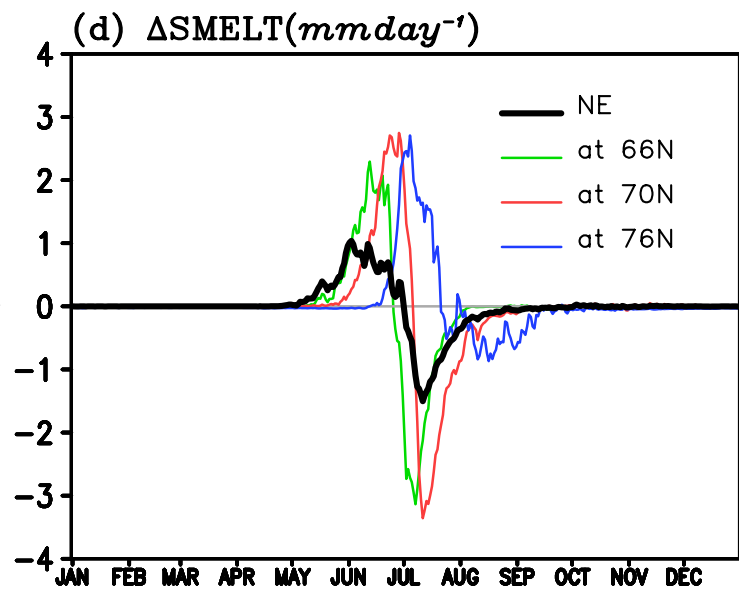
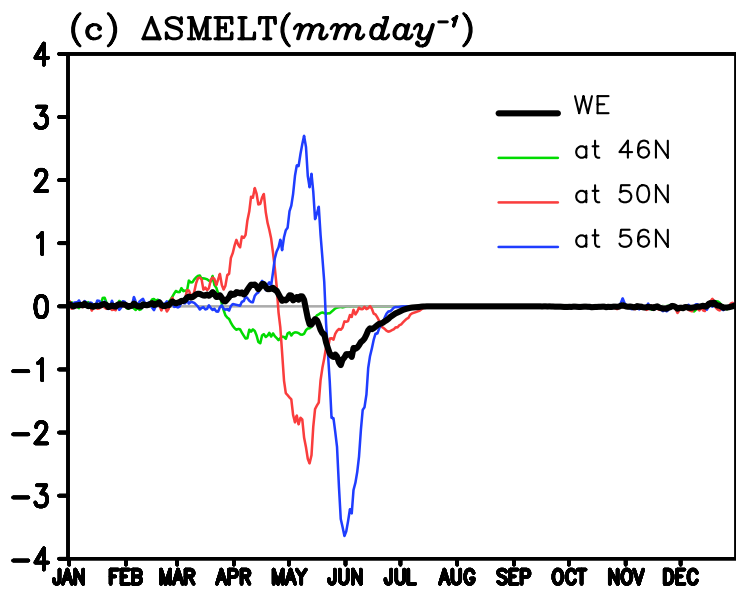
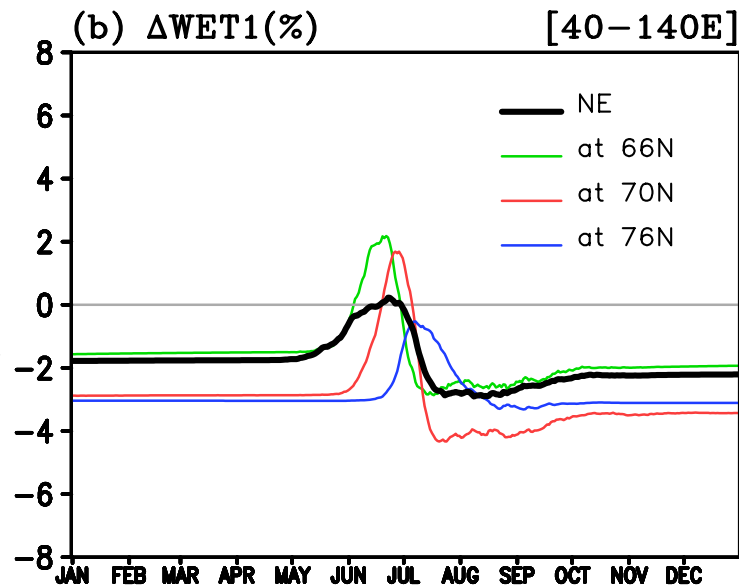
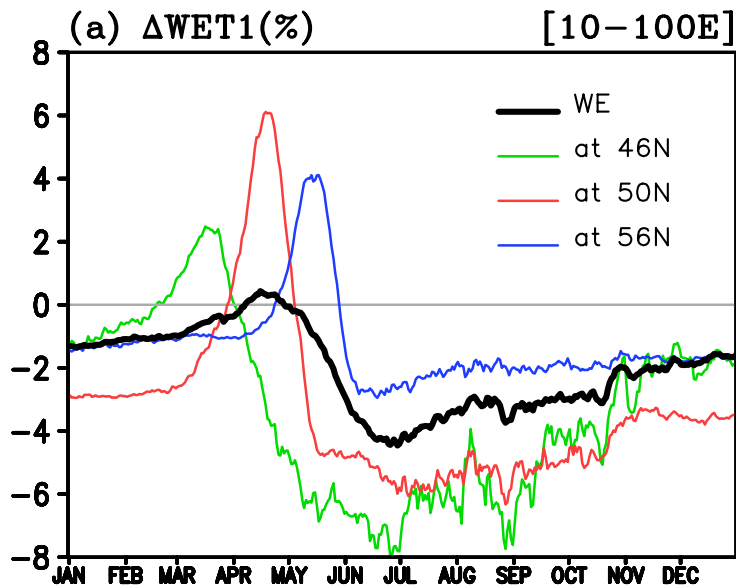
70 N



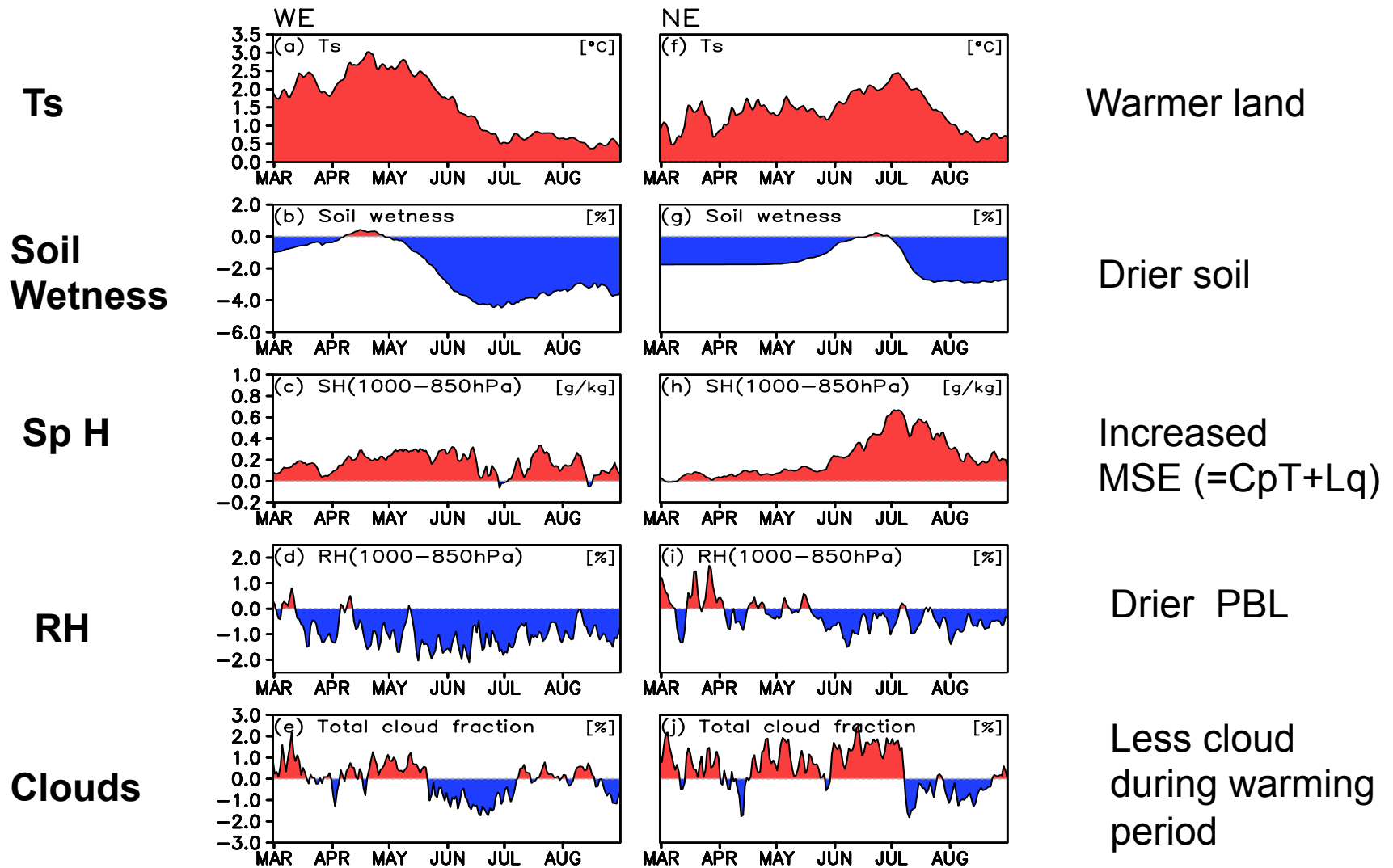
The Wet-First- Dry-Later (WFDL) effect:

Fast and brief soil wetting followed by fast and prolonged drying at progressively later time from following the movement of the snowline from high to low latitudes





SDE induced hydroclimate anomalies



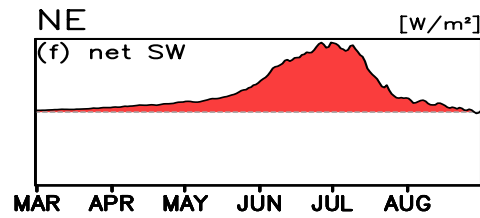
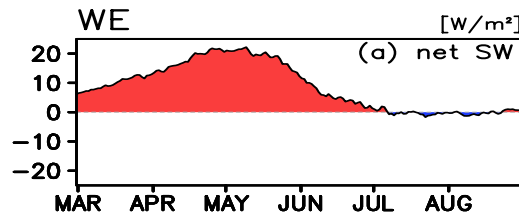
$$\delta RH = \delta q / q_s - \alpha R_h \delta T$$

$$\alpha = L(R_v T^2)^{-1} \sim 6\% K^{-1}$$

CC relationship,

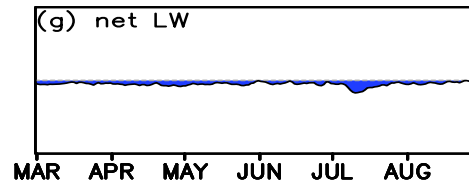
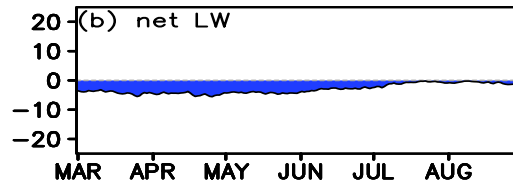
Land surface energy balance

SW

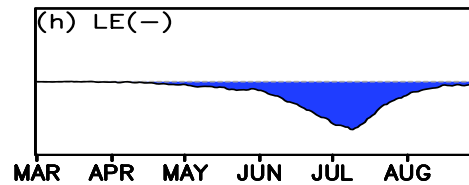
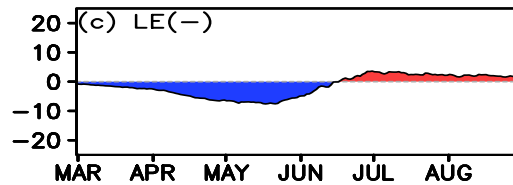


Land gain
Land lost

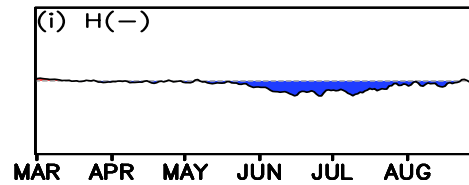
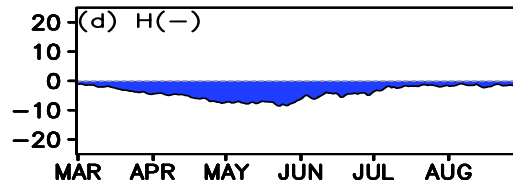
LW



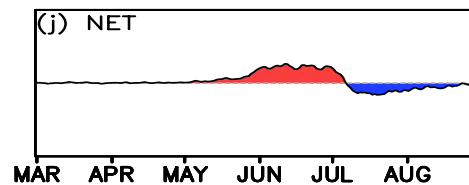
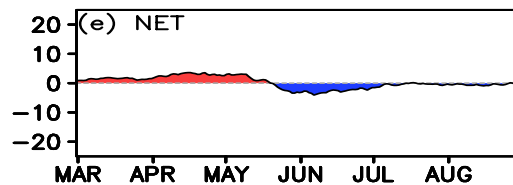
LE



H



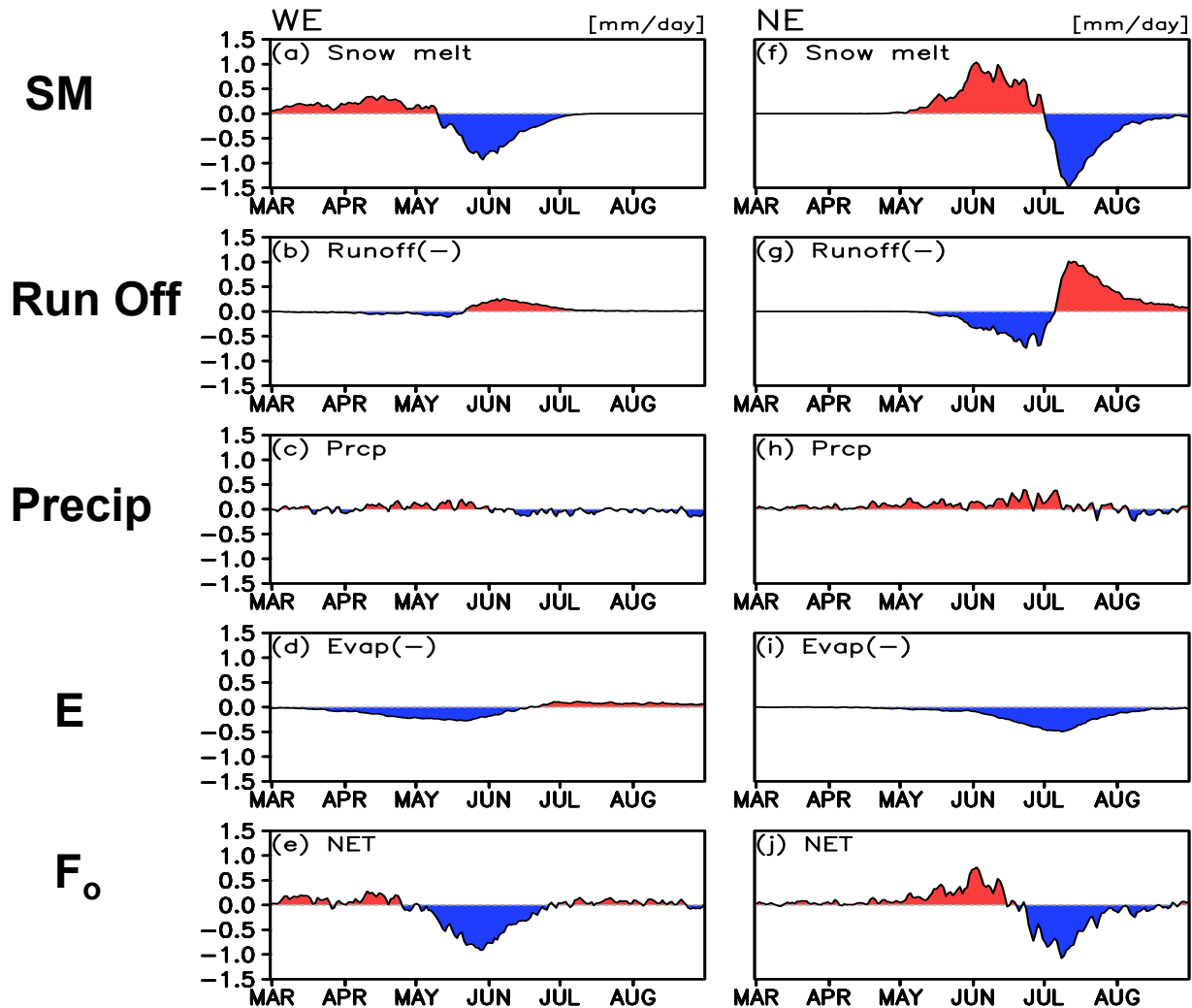
Net



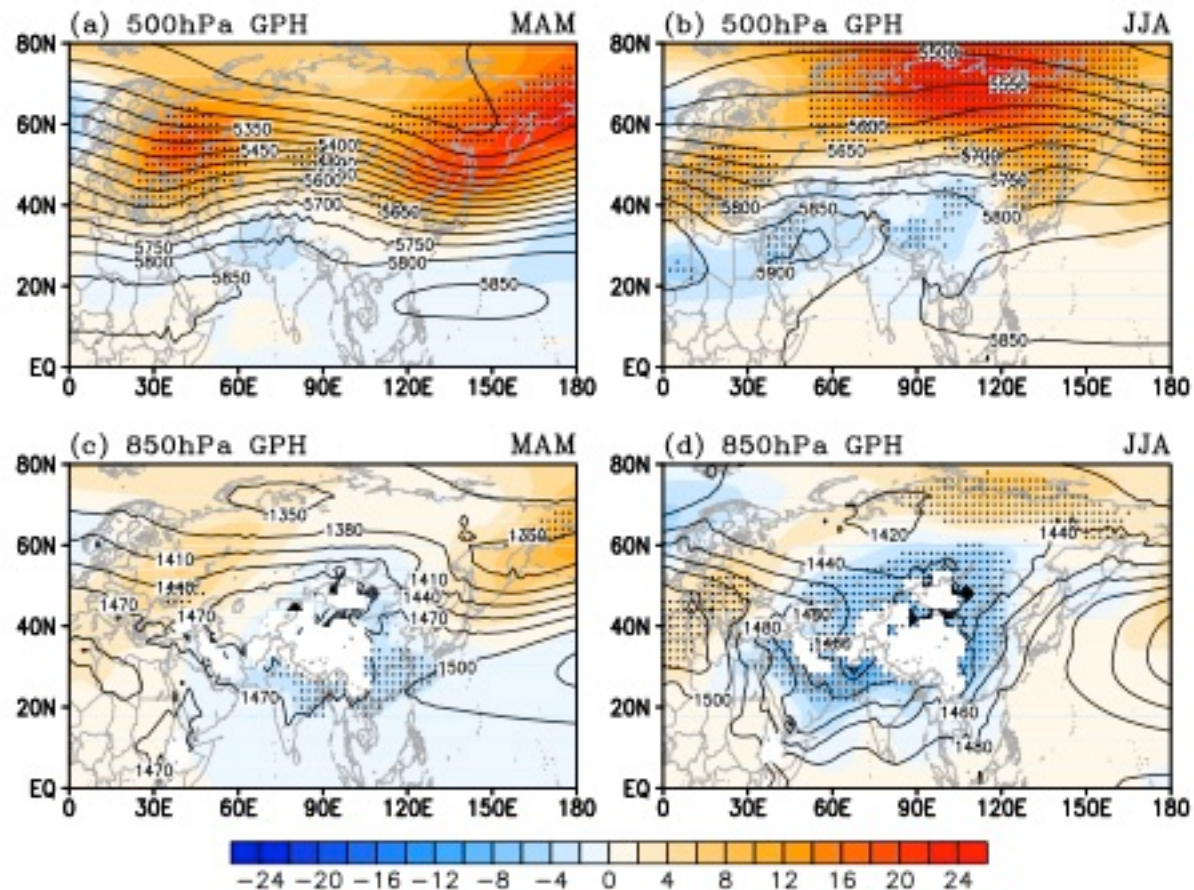
Land Surface Water Balance:

$$F_o = SM - R + P - E$$

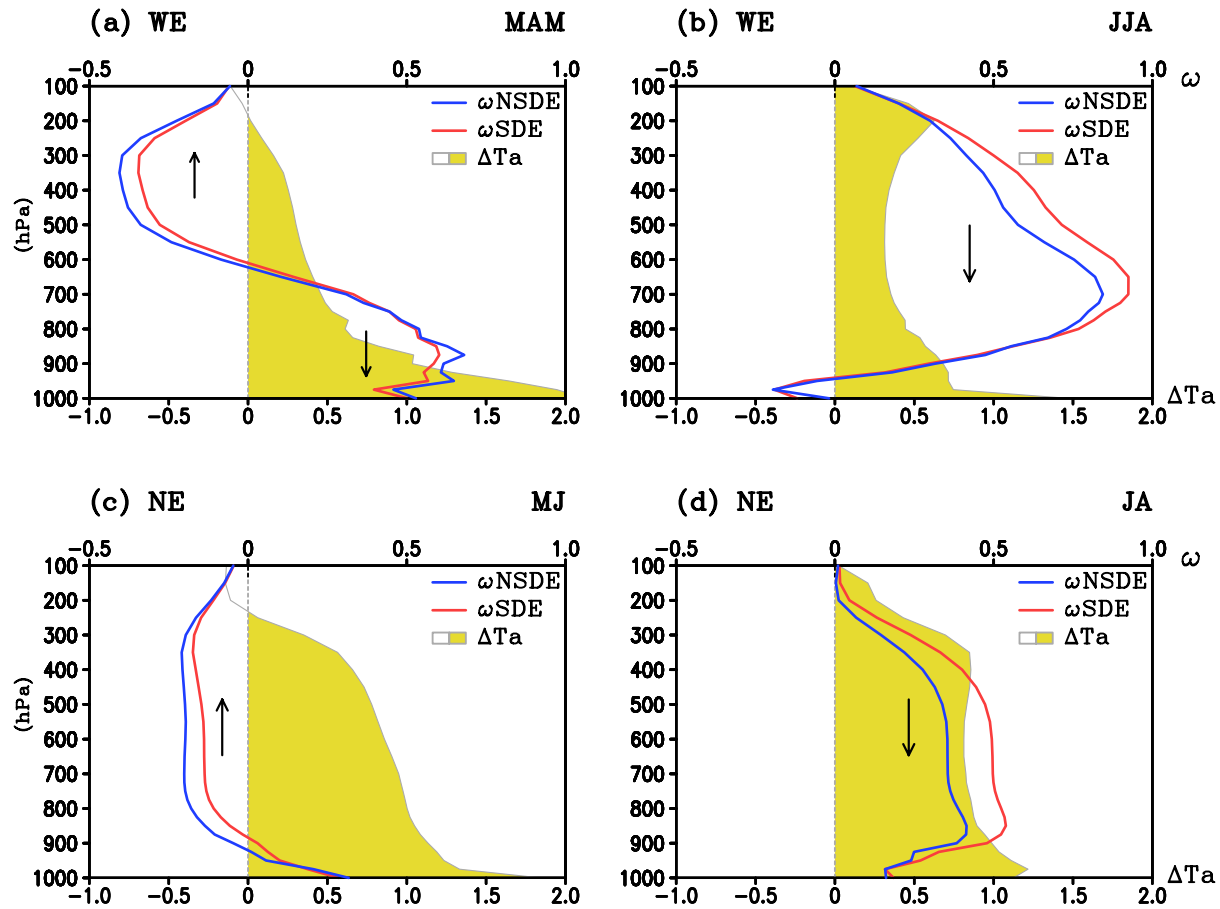
Land Gain
Land Lost



Favorable conditions for heatwaves:
High seasonal mean GPH over WE and NE,



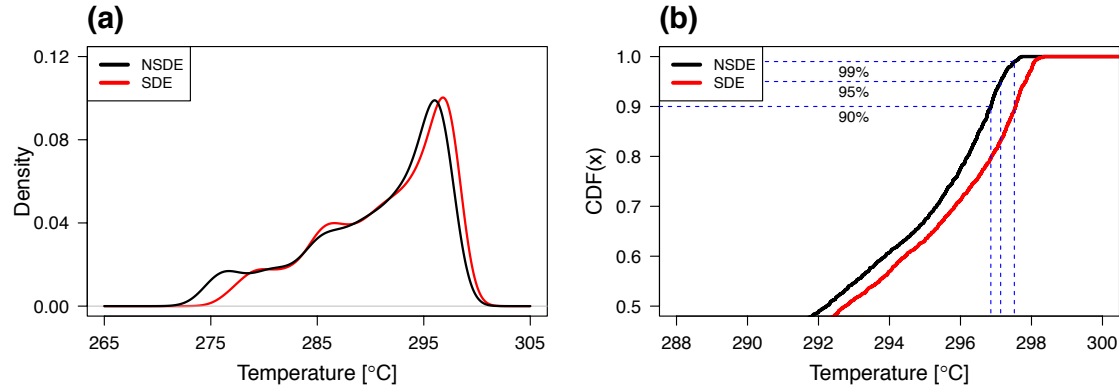
SDE induces warmer surface and troposphere, and stronger subsidence



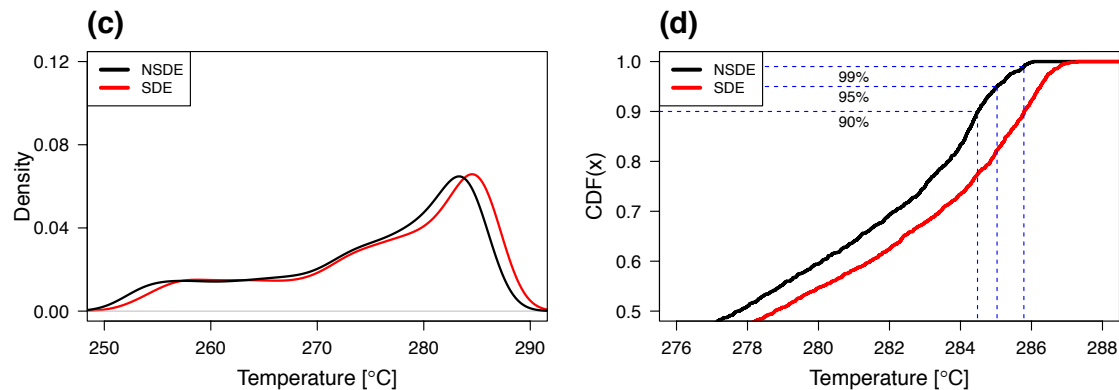
→ Increased subsidence →

SDE increases frequency of 1% (~10%) extreme heating days (NSDE) by 10 (~20) fold)

WE

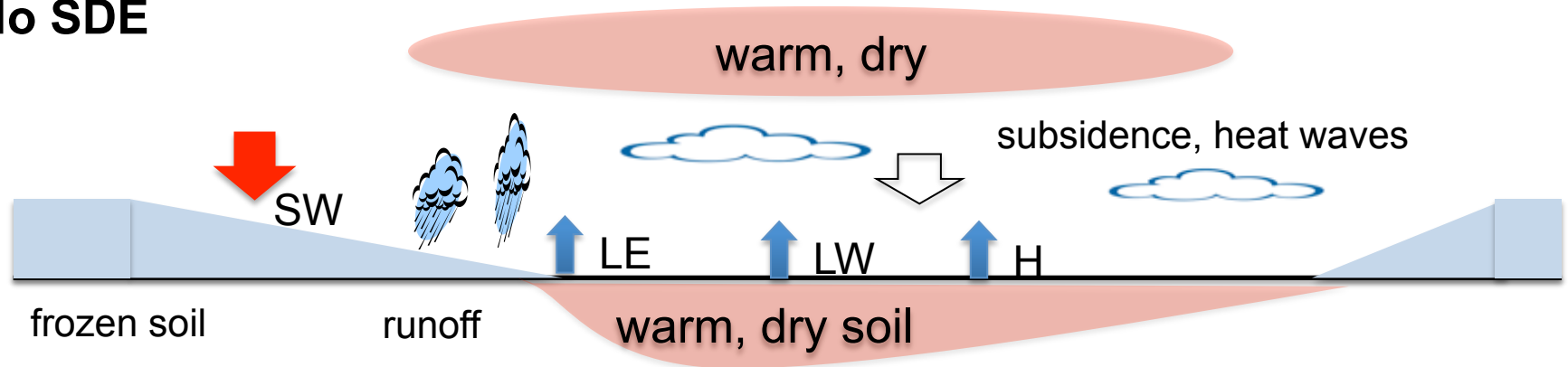


NE

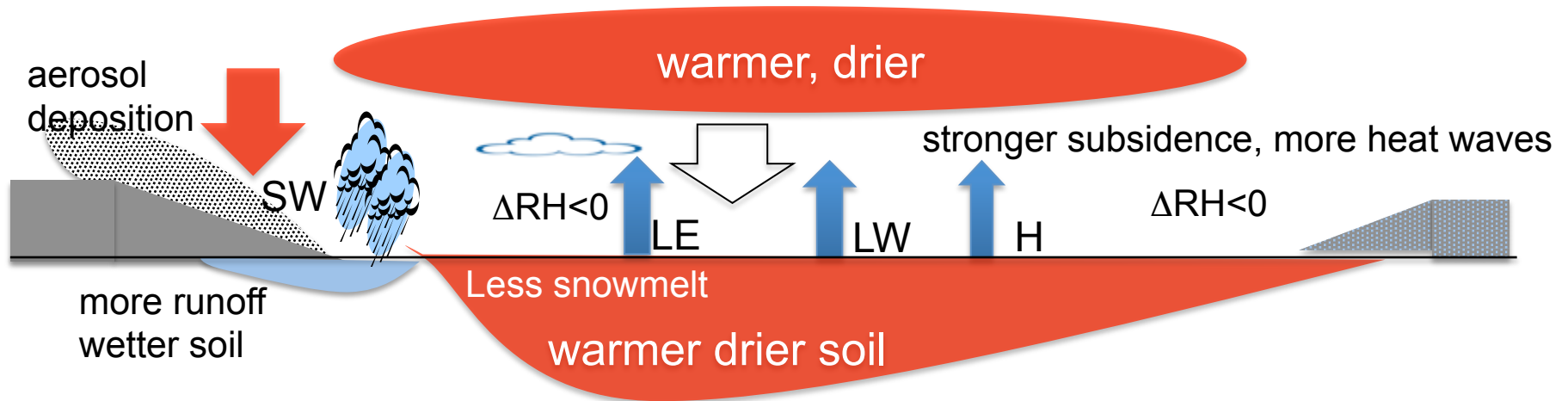


| NSDE | | 10% | 5% | 1% |
|------|----|-----|----|----|
| SDE | WE | 20 | 17 | 11 |
| | NE | 23 | 18 | 10 |

No SDE



SDE



Spring

Summer

Fall

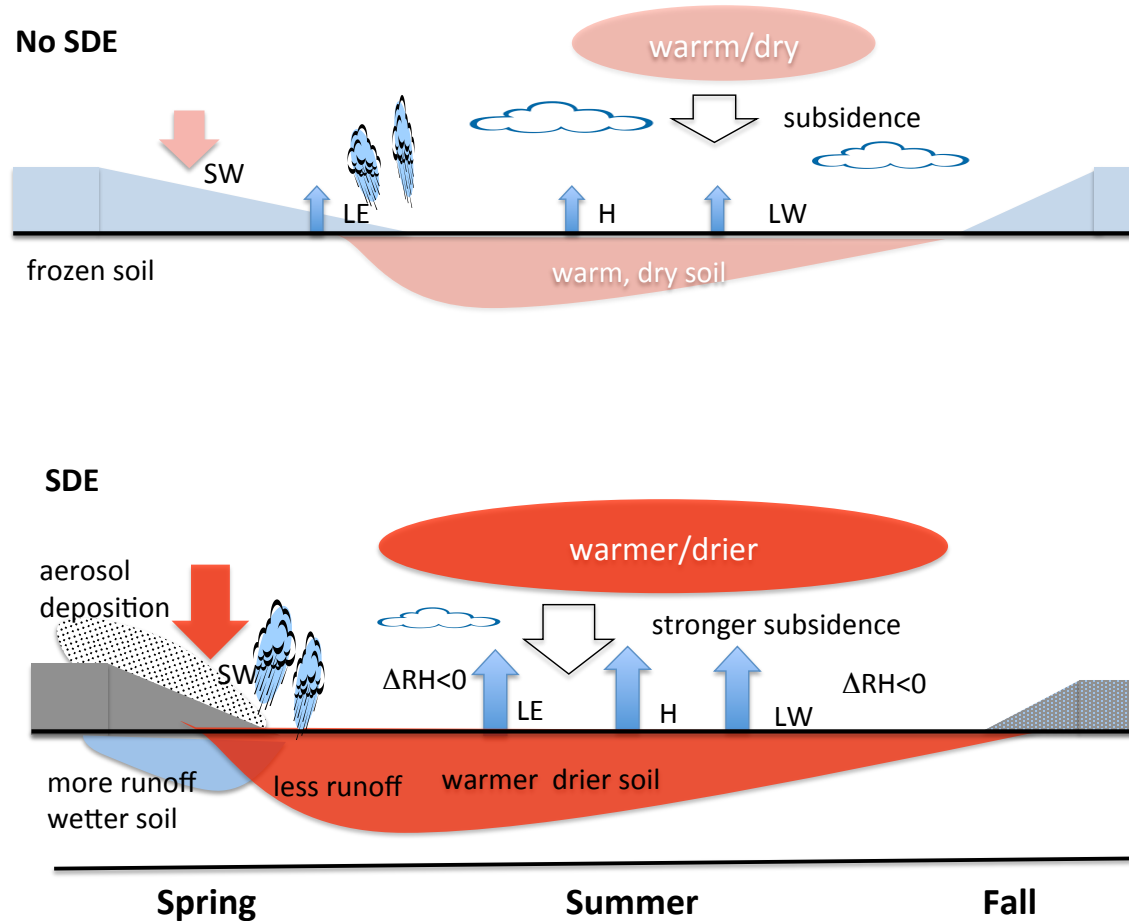
Thank you

Conclusions

Snow-darkening effects by deposition of absorbing aerosols (dust, BC, and OC) leads to:

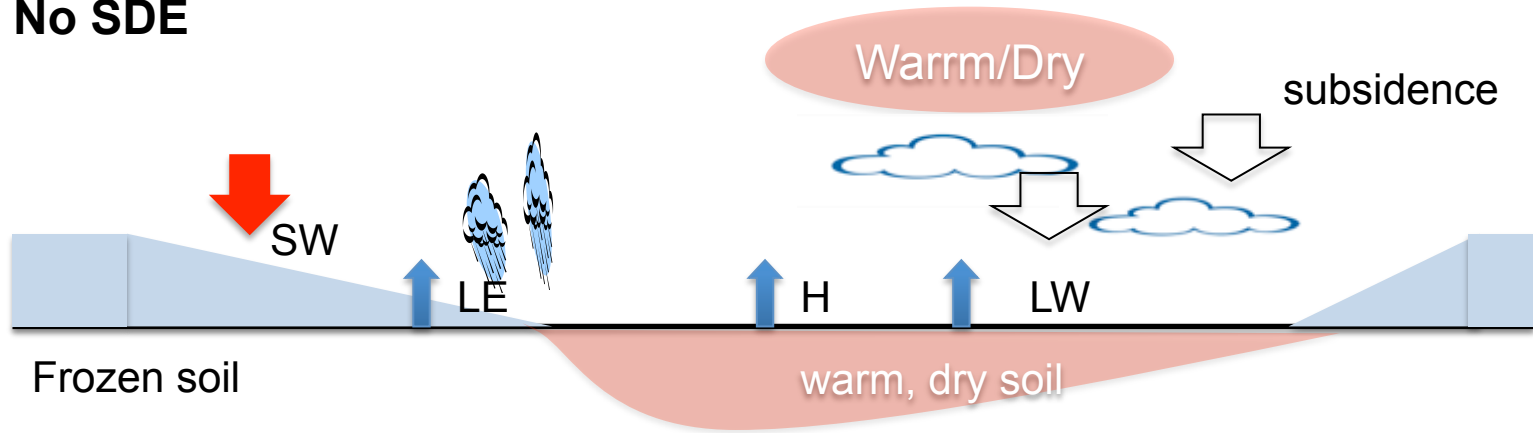
- **a warmer and drier climate over Eurasia** :an increase in annual mean surface skin temperature of up 2.5°C, with a substantial reduction (up to 10% in soil wetness over vast regions of the Europe and increased Asian summer monsoon precipitation.
- **A continental scale hydroclimate and land-atmosphere WFDL feedback**, with fast and earlier snowmelt near seasonal migrating snowline, followed by prolonged and widespread over the snow-covered region of extratropical Eurasia land and high mountain regions
- **Increased the top1% (10%) extreme heating days (based on NSDE) over Western Europe, and northern Eurasia by more than 10 (20) times**

The Wet-First-Dry-Later (WFDL) hydroclimate feedback induced by snowdarkening effect (SDE)

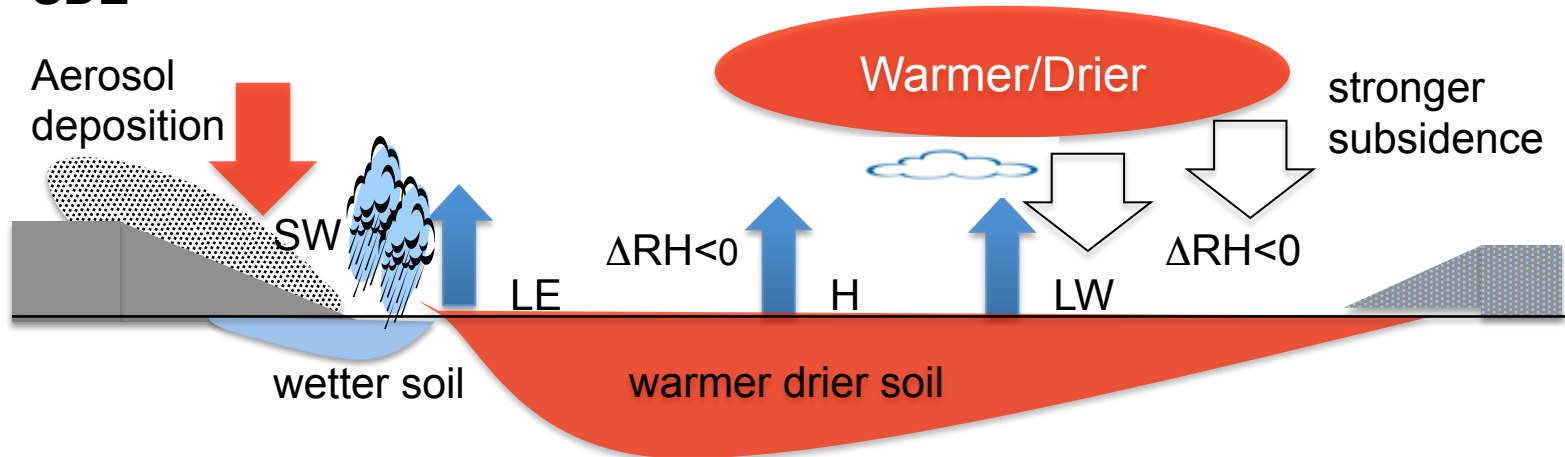


A Wet-First-Dry-Later (WFDL) hydro-climate feedback induced by aerosol snow-darkening (SDE)

No SDE



SDE

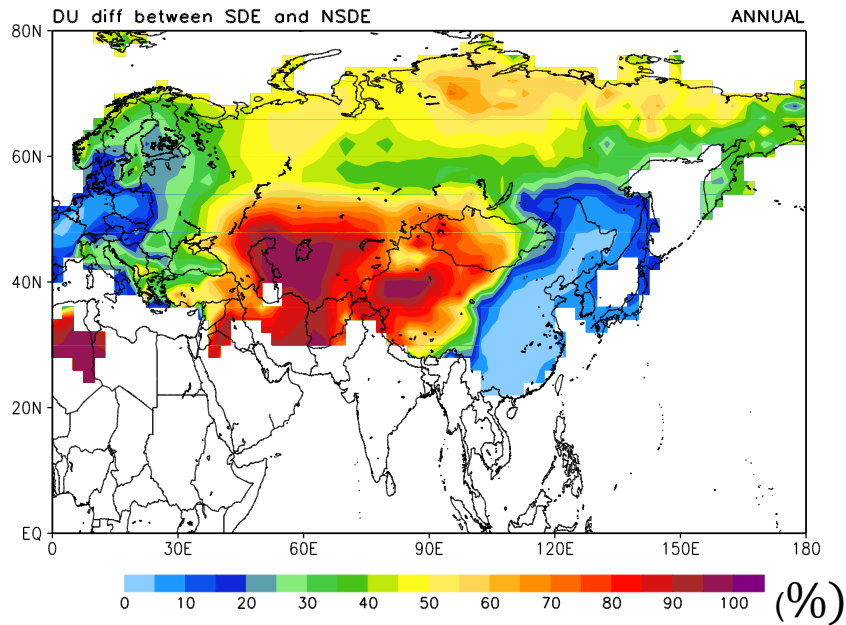


Spring

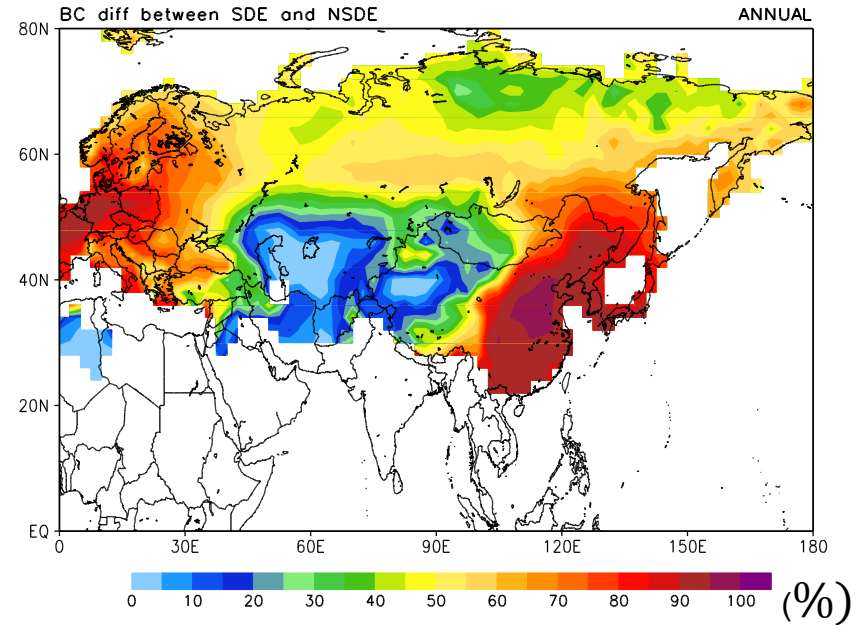
Summer

Fall

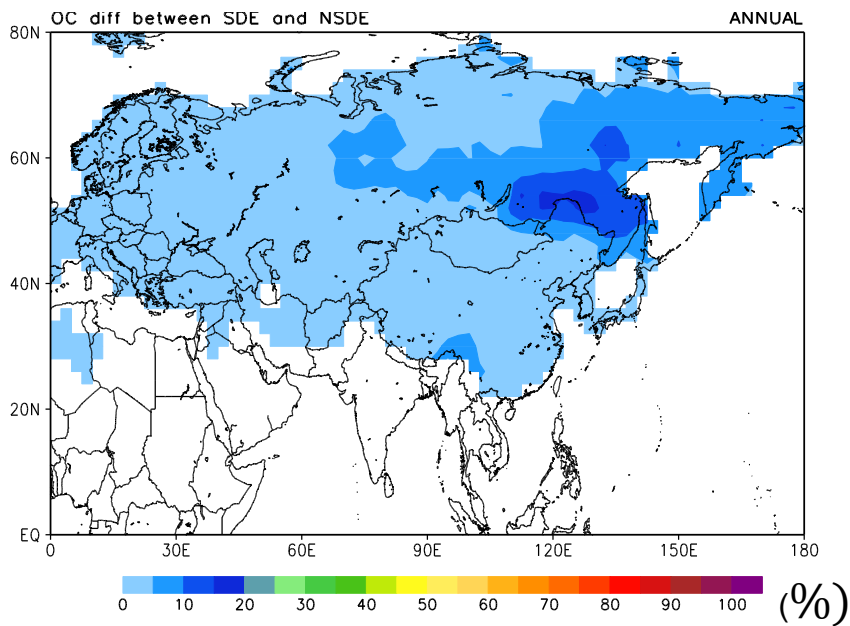
Dust



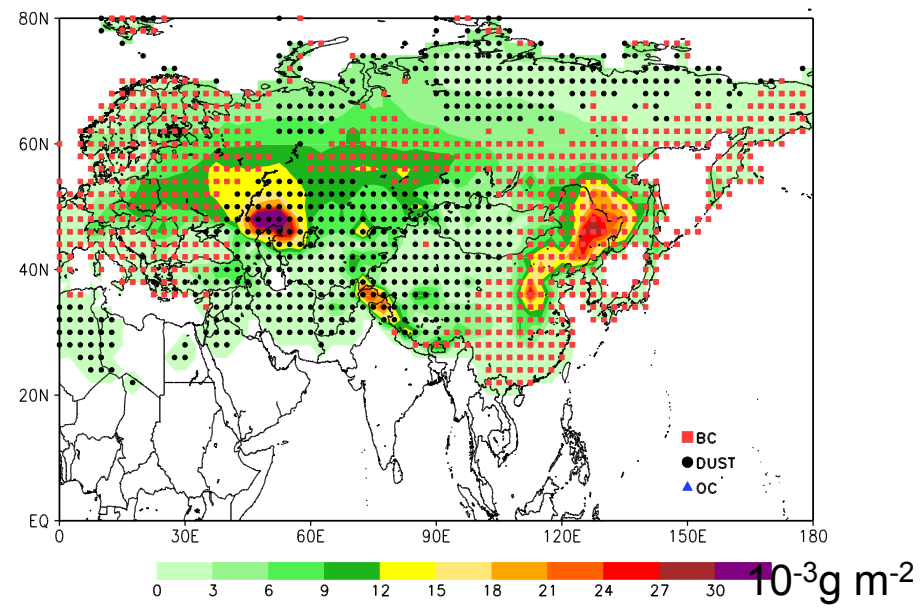
BC



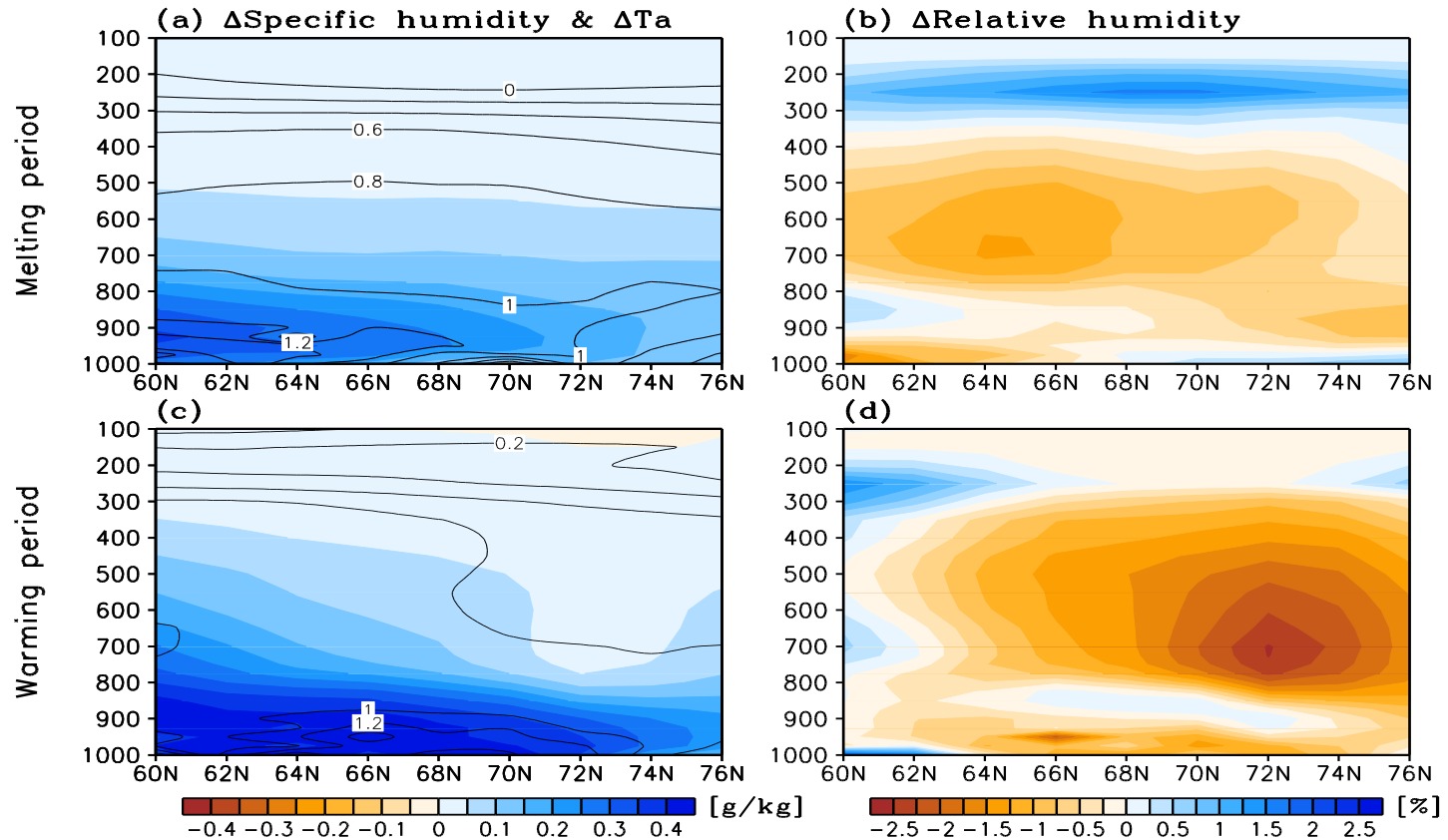
OC



Total snow impurity absorption efficiency



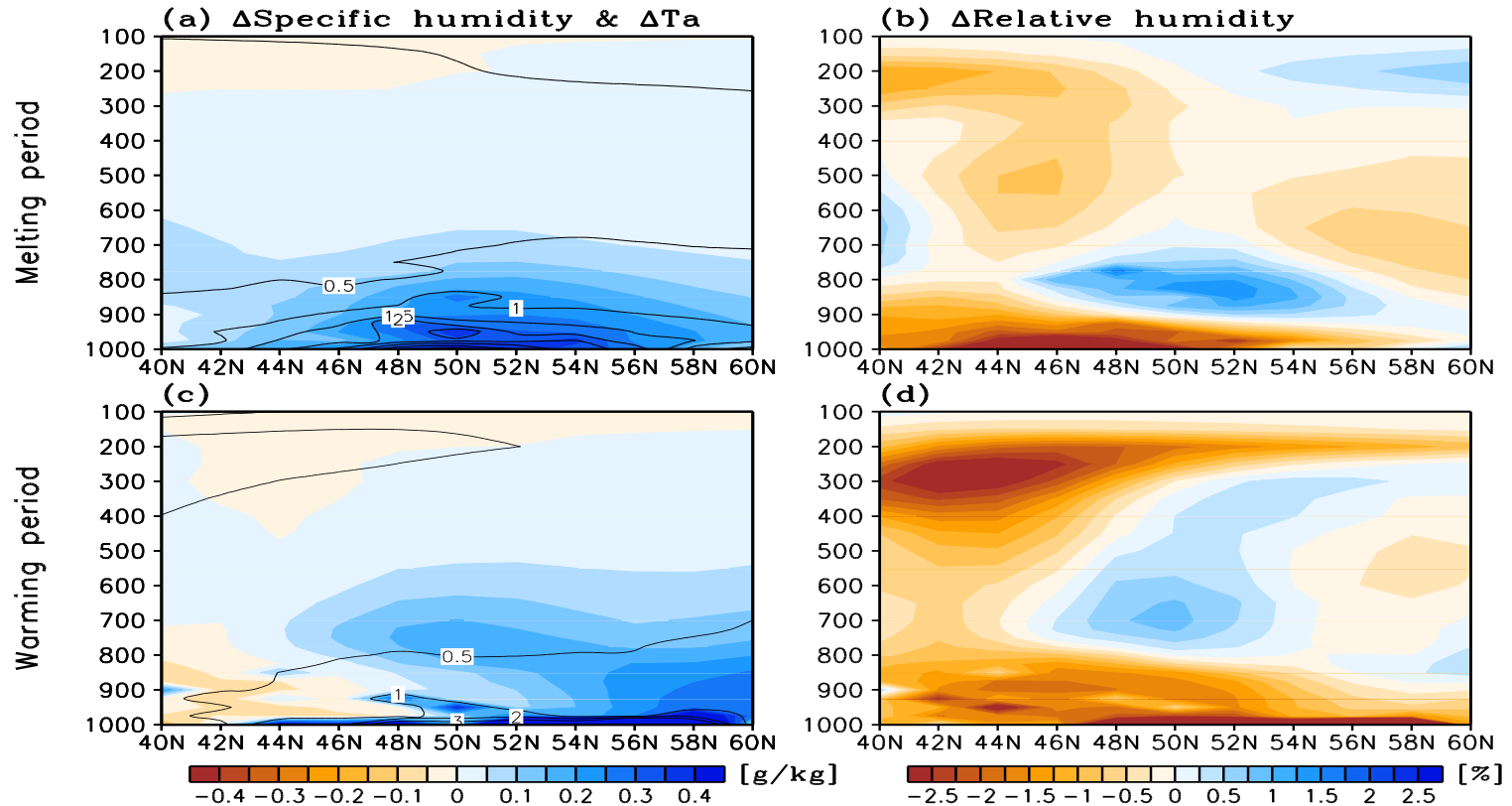
Northern Eurasia



Increase near surface and
mid-troposphere
MSE ($C_p T + q$)

Drier ($\Delta RH < 0$) surface and upper
troposphere

Western Europe



Increase near surface and
mid-troposphere
MSE ($C_p T + q$)

Drier ($\Delta RH < 0$) surface and
upper troposphere